

Theoretical and Computational Aspects of Nonlinear Surface Waves, Banff, Canada, 30.10-4.11.2016

Ship generated tsunamis: linearity vs. nonlinearity

John Grue, Dept. of Mathematics, Univ. of Oslo, Norway

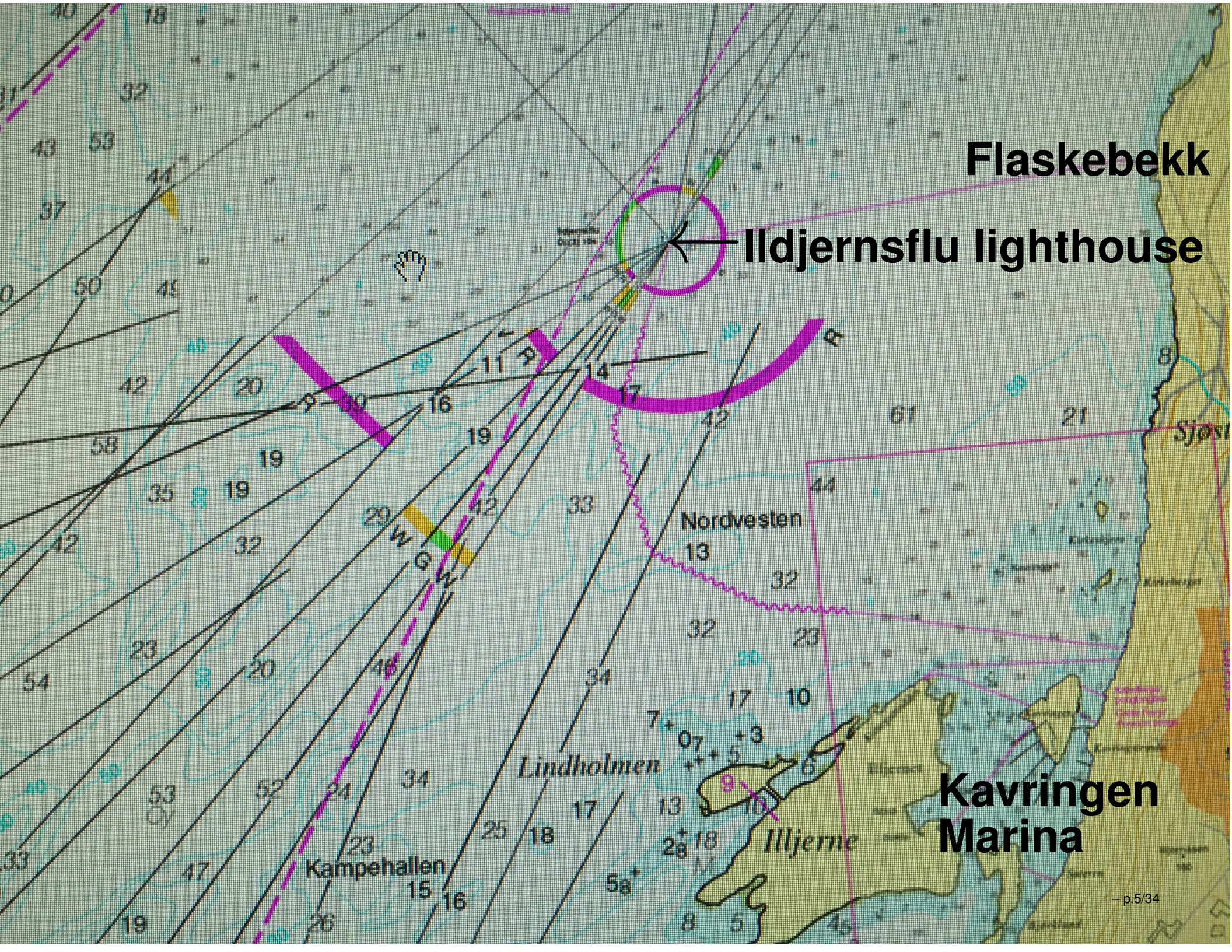


Outline

1. observations of a new phenomenon
2. linear dispersive analysis
3. asymptotics
4. full calculations
- (5. potential effect of nonlinearity)
6. conclusions
7. movie!







Flaskebekk

Ildjærnsflua lighthouse

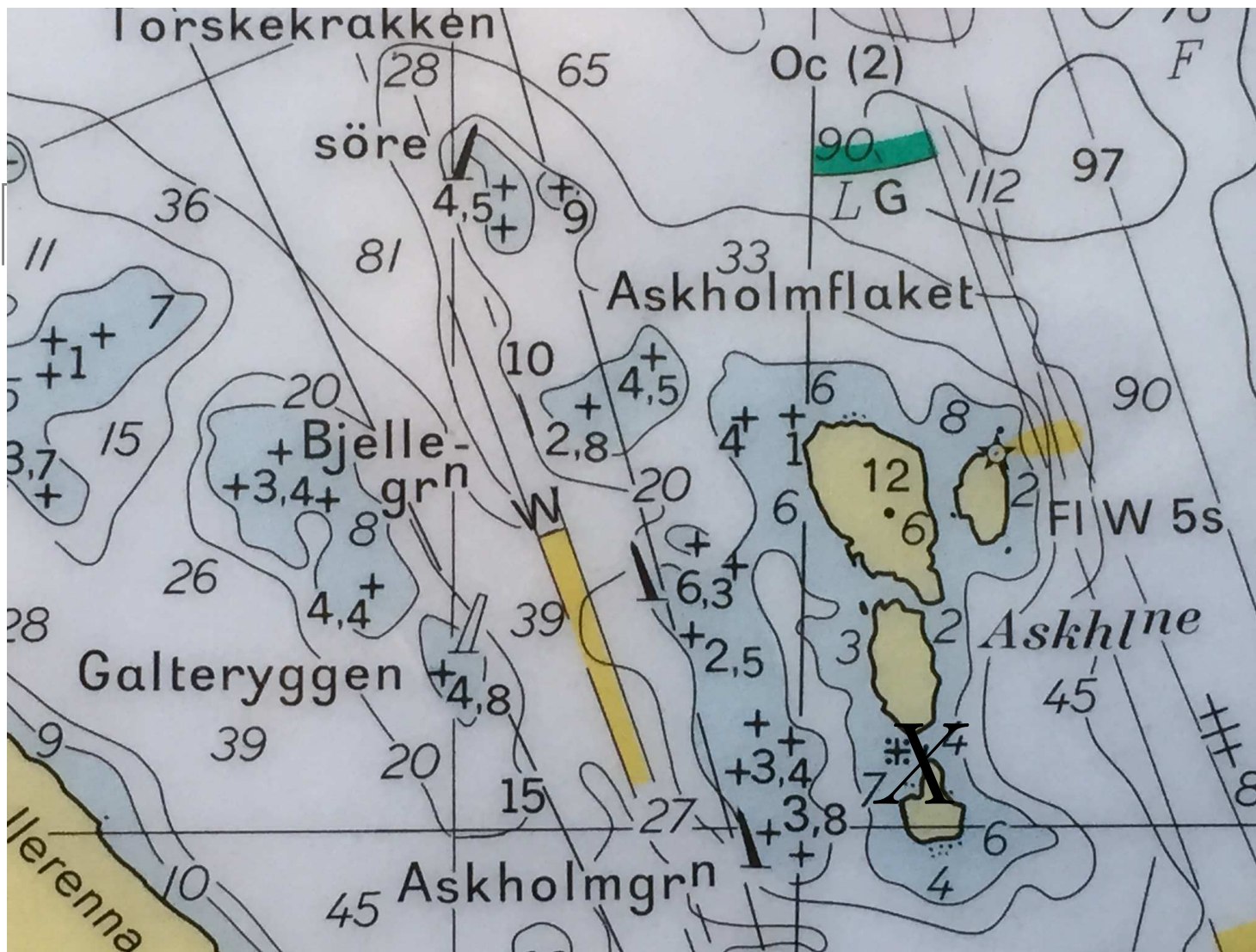
Nordvesten

Lindholmen

Kampehallen

Illjerne

**Kavringen
Marina**



The Oslofjord stretches 100 km north-south

Wave periods between two consecutive inflows (in-in) and outflows (out-out) in harbors or small bays due to three different ships observed at three different positions.

Fantasy, 10.3 ms^{-1}

DFDS, 8.6 ms^{-1}

Magic, $U = 7.8 \text{ ms}^{-1}$

Magic, $U = 10 \text{ ms}^{-1}$

Position 1

Position 1

Position 2

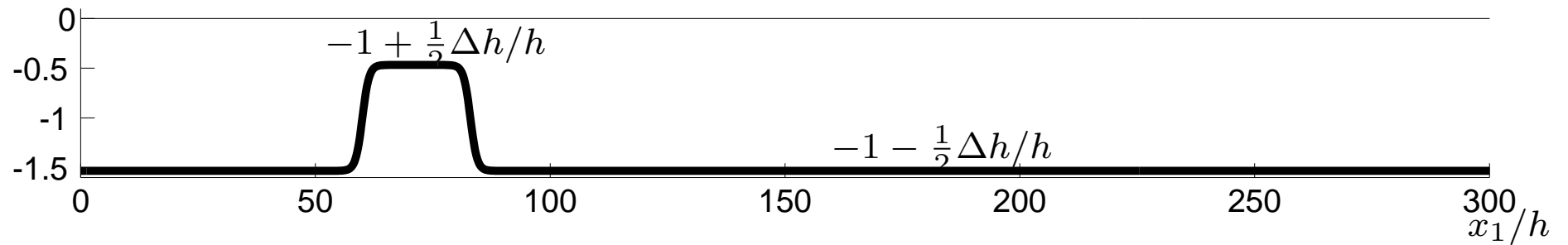
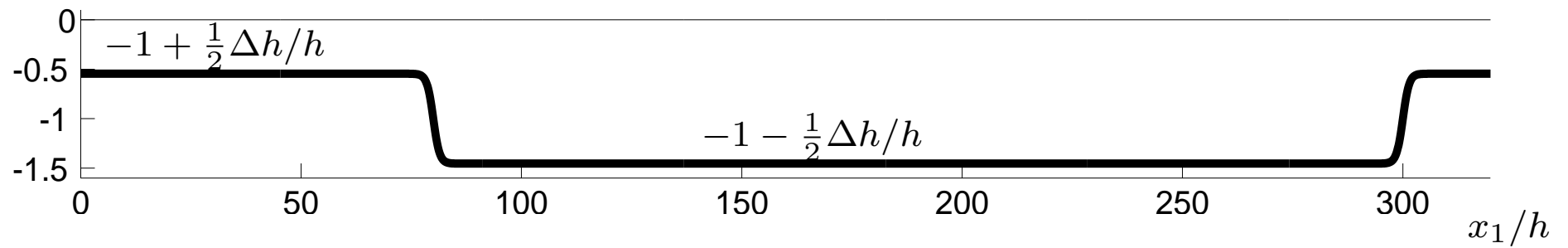
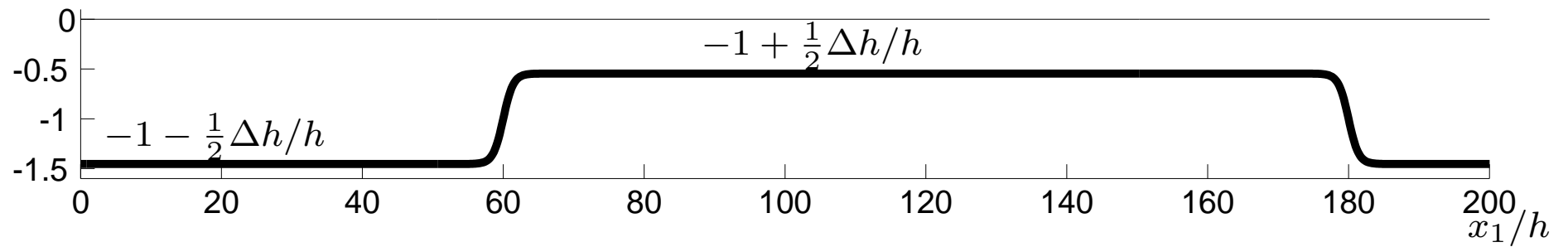
Position 3

wave no.	in-in	out-out	in-in	out-out	in-in	out-out	in-in	out-out
1	34 s	36 s	26 s	34 s	58 s	57 s	35 s	30 s
2	32 s	-	34 s	-	49 s	40 s	29 s	34 s
3	13 s	-	-	-	64 s	-	46 s	36 s



1. Very large cruise ferries moving along the Oslo fjord generate very long upstream waves: Color Line (36 100 m³, since 2004) vs. DFDS (24 000 m³, older).
2. Periods $\sim 30 - 40 / \sim 50 - 60$ sec.; shorter waves follow. .
3. Crests extend across the 2-3 km wide fjord.
4. Wave height of 1 m (1.4 m) / currents of 1 ms⁻¹ documented in small harbors / bays.
5. Generation: ship running over a depth change of Δh which is comparable to the average depth h , i.e. $\Delta h/h \sim 1$.
6. Depth Froude number $U/\sqrt{gh} \sim 0.5 - 0.7$. Not related to nonlinear transcritical generation of upstream solitons on flat bottom (Ertekin et al. 1986, Wu, 1987; Pedersen, 1988).
7. Problem: Considerable erosion along the shoreline.

Bottom shapes $y = -h + \beta(x)$



Fully dispersive linear theory. Kinematic and dynamic free surface condition (Fourier transformed):

$$\frac{\partial \hat{\eta}}{\partial t} - \hat{V} = 0, \quad \frac{\partial \hat{\phi}_F}{\partial t} + g\hat{\eta} = -\frac{\hat{p}}{\rho},$$

A set of integral equations connect the vertical surface velocity V , surface potential ϕ_F and the bottom potential $\phi_B = \phi(y = -h + \beta(x_1, x_2))$ (Clamond and Grue, 2001; Fructus and Grue, 2007; Grue, 2015):

$$\begin{aligned} & \int_F V \left(\frac{1}{r} + \frac{1}{r_1} \right) dS \\ &= 2\pi\phi'_F + \int_F V \left(\frac{1}{r} + \frac{1}{r_1} \right) \phi_F dS + \int_B \phi_B \frac{\partial}{\partial n} \left(\frac{1}{r} + \frac{1}{r_1} \right) dS \\ \\ & 2\pi\phi'_B = \int_F V \left(\frac{1}{r} + \frac{1}{r_{1B}} \right) dS - \int_B \phi_B \frac{\partial}{\partial n} \left(\frac{1}{r} + \frac{1}{r_{1B}} \right) dS \end{aligned}$$

Green fu. expanded in the bottom variation β obtaining

$$\hat{V} = k\hat{\phi}_F T_1 + \frac{i\mathbf{k}}{C_1} \cdot [\mathcal{F}(\beta\nabla_1\phi_B) + \frac{k^2}{3!}\mathcal{F}(\beta^3\nabla_1\phi_B) + \dots]$$

$$\begin{aligned} k\hat{\phi}_B &= \frac{k\hat{\phi}_F}{C_1} - iT_1\mathbf{k} \cdot \mathcal{F}(\beta\nabla_1\phi_B) + \frac{k}{2-2e_1}\mathcal{F}(\beta^2\mathcal{F}^{-1}(k\hat{\phi}_B)) \\ &+ \frac{k}{1-e_1}\mathcal{F}(\beta\mathcal{F}^{-1}[i\mathbf{k} \cdot \mathcal{F}(\beta\nabla_1\phi_B)]) - \frac{ke_1}{2-2e_1}i\mathbf{k} \cdot \mathcal{F}(\beta^2\nabla_1\phi_B) \\ &+ \dots \end{aligned}$$

$$C_1 = \cosh kh, T_1 = \tanh kh, e_1 = e^{-2kh}.$$

The elevation:

$$\hat{\eta} = \int_{t_0}^t \cos \omega(s - t) \hat{h}_1(s) ds + \int_{t_0}^t \sin \omega(s - t) \frac{\omega \hat{p}}{\rho g} ds,$$

$$h_1 = \frac{\mathbf{i}k}{C_1} \cdot [\mathcal{F}(\beta \nabla_1 \phi_B) + \frac{k^2}{3!} \mathcal{F}(\beta^3 \nabla_1 \phi_B) + \dots]$$

Term 1: effect of the bottom variation;

Term 2: effect of the moving pressure distribution.

Asymptotic analysis of the waves due to a small bottom step at $x_1 = 0$, where $y = -h$ ($x_1 < 0$) and $y = -h + \Delta h$ ($x_1 > 0$) with $\Delta h/h \ll 1$; waves are assumed long (small kh):

$$\hat{\eta}_0 = \int_{t_0}^t \cos \omega(s - t) \hat{h}_1(s) ds$$

$$\hat{h}_1 = \frac{i\mathbf{k}}{C_1} \cdot [\mathcal{F}(\beta \nabla_1 \phi_B) + \dots] = i\mathbf{k} \cdot \mathcal{F}(\beta \nabla_1 \phi_F) + O(k^2)$$

$$\hat{h}_1 = ik_1 \Delta h \int_0^\infty \int_{-\infty}^\infty \frac{\partial \phi_F}{\partial x_1} e^{-i\mathbf{k} \cdot \mathbf{x}} d\mathbf{x} + O(k^2)$$

$$= -ik_1 \Delta h \int_{-\infty}^\infty \phi_F(x_1 = 0) dx_2 + O(k^2)$$

The time derivative:

$$\frac{\partial \hat{h}_1}{\partial t} = -ik_1 \Delta h \int_{-\infty}^{\infty} \frac{\partial \phi_F(x_1 = 0)}{\partial t} dx_2 \simeq \frac{ik_1 \Delta h}{\rho} \int_{-\infty}^{\infty} p(x_1 = 0) dx_2$$

Obtaining the elevation:

$$\begin{aligned} \hat{\eta}_0 &= - \int_{t_0}^t \frac{\sin \omega(s - t)}{\omega} \frac{\partial \hat{h}_1}{\partial s} ds \\ &\simeq - \frac{ik_1 \Delta h}{\rho} \int_{t_0}^t \frac{\sin \omega(s - t)}{\omega} \int_{-\infty}^{\infty} p(x_1 = 0) dx_2 ds, \\ \omega &= \sqrt{gkT_1} \end{aligned}$$

Pressure distribution: $p(x_1 - Ut, x_2, t) = \rho g V_0 \delta(x_1 - Ut) \delta(x_2)$:

$$\hat{\eta}_0 = \frac{ik_1 \Delta h V_0}{U\omega/g} \sin \omega t,$$

Method of stationary phase gives upstream waves, in 2D,
($t \rightarrow \infty$)

$$\eta_0(x_1, t) = -\frac{k_{1,0} \Delta h V_0}{(U \omega_0 / g)(2\pi |\omega''| t)^{1/2}} \cos(k_{1,0} x_1 - \omega_0 t + \frac{\pi}{4}) + O(t^{-1}).$$

$k_{1,0}$: $\partial \omega / \partial k_1 = x_1 / t < \sqrt{gh} = c_0$, $\omega_0 = \omega(k_{1,0})$, double prime means double derivative. Near the wave front ($x_1 / t > \sim \sqrt{gh}$)

$$\eta_0(x_1, t) \sim -\frac{\Delta h V_0}{(U c_0 / g)(4c_0 h^2 t)^{1/3}} \mathbf{Ai}(Z),$$

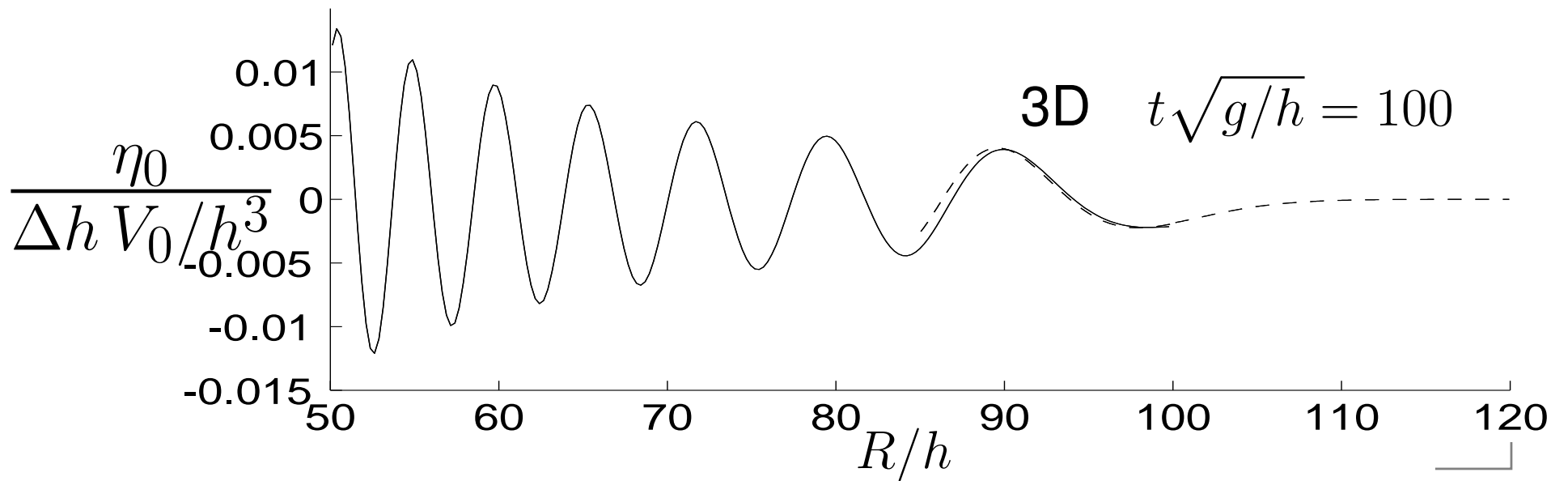
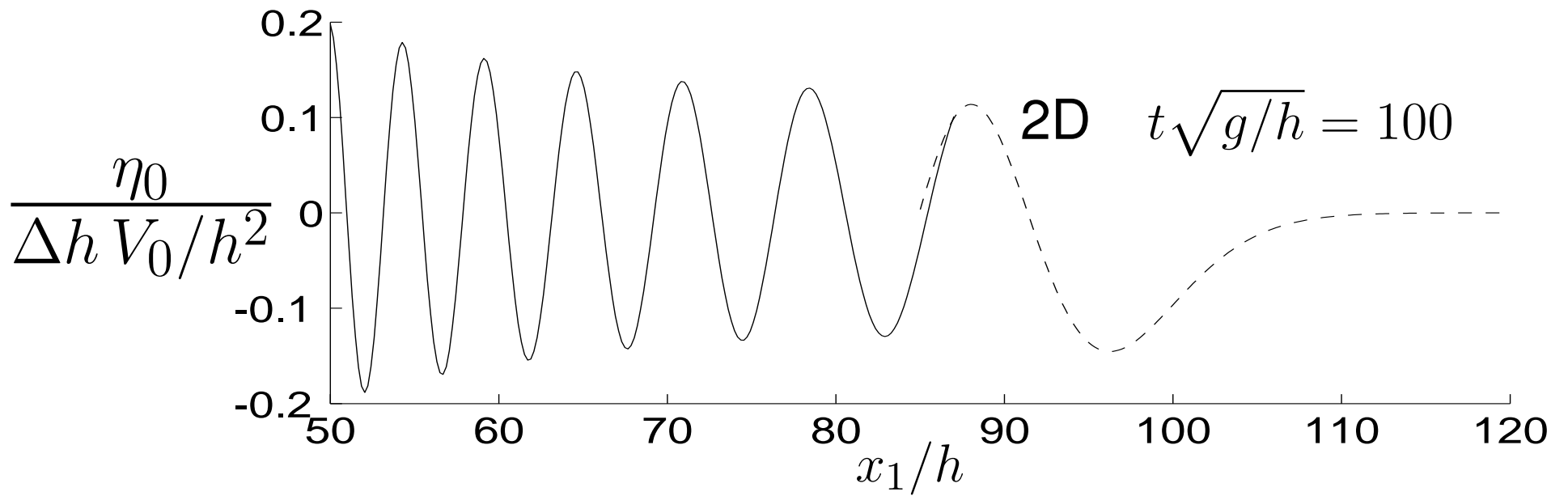
$\mathbf{Ai}(Z)$ – the Airy function; $Z = [2(x_1 - c_0 t)^3 / c_0 h^2 t]^{1/3}$.

Inverse transform in the three-dimensional case gives:

$$\eta_0(R, t) \sim -\frac{\Delta h V_0 k_0}{2\pi U \omega_0 / g} \frac{\partial}{\partial x_1} \frac{\sin(k_0 R - \omega_0 t)}{(k_0 R |\omega_0''| t)^{1/2}},$$

where k_0 is determined by $R/t = \partial\omega/\partial k < \sqrt{gh}$, $\omega_0 = \omega(k_0)$.
Near the wave front, following Clarisse, Newman and Ursell (1995)

$$\eta_0(x_1, x_2, t) \sim \frac{\Delta h V_0 / h^2}{2\pi U / \sqrt{gh}} \frac{\partial}{\partial x_1} \frac{2^{1/3} \pi}{(R^* / t^*)^{1/2} t^{*2/3}} \times \\ \times [E_0 \mathbf{Ai}^2 - t^{*-2/3} C_0 \mathbf{Ai}'^2 + 2t^{*-4/3} A_1 \mathbf{Ai} \mathbf{Ai}'].$$



Numerical wave tank:

$$(L_1, L_2) = (200h_0, 13h), \quad h = 55 \text{ m (location 1)}.$$

$$(L_1, L_2) = (300h_0, 80h), \quad h = 30 \text{ m (location 2)}.$$

$$(N_1, N_2) = (1200, 432) \text{ or } (2090, 248).$$

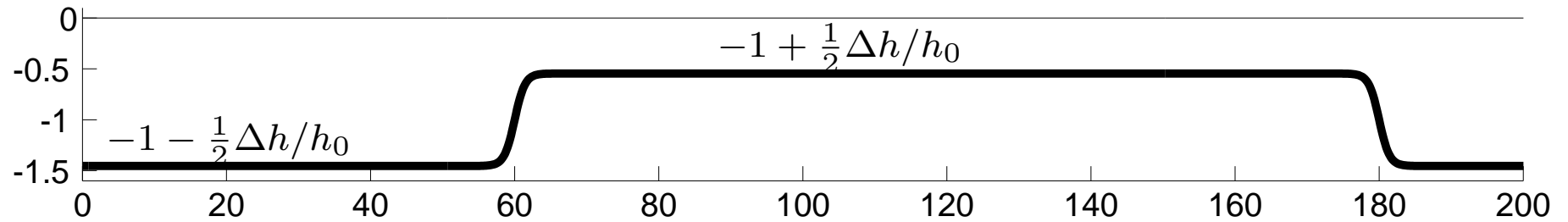
Ship replaced by a hovercraft/pressure distribution of
(length, width, draught) = (l, w, d) ,

$$l/h_0 = 210 \text{ m} / h \sim 4 - 7,$$

$$w/h_0 = 30 \text{ m} / h \sim \frac{1}{2} - 1,$$

draught such that $\int p/(\rho g) dx_1 dx_2 = V_{ship}$.

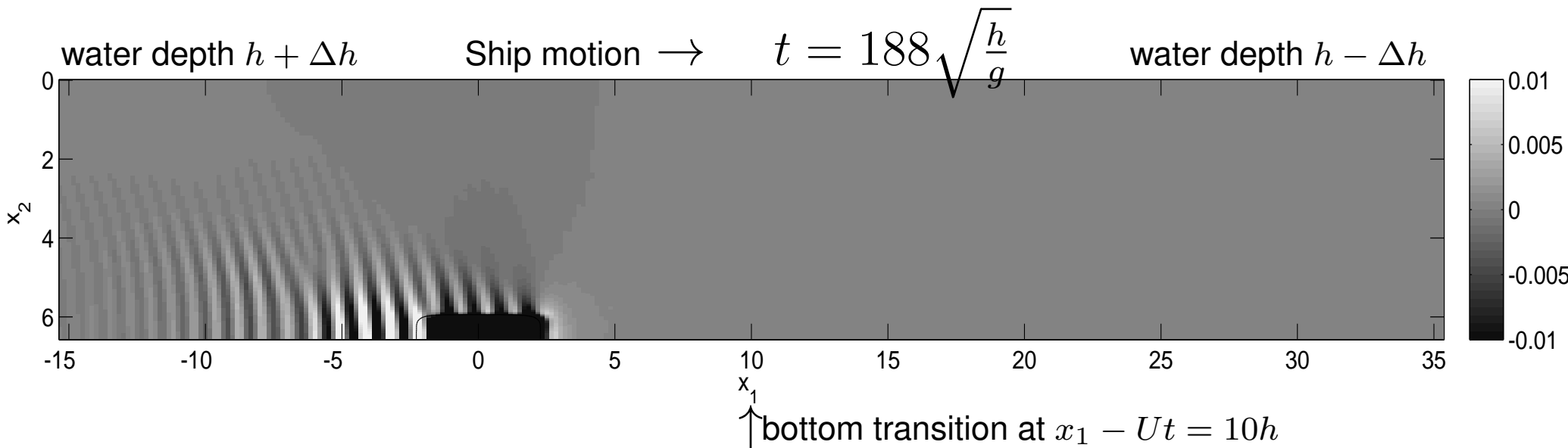
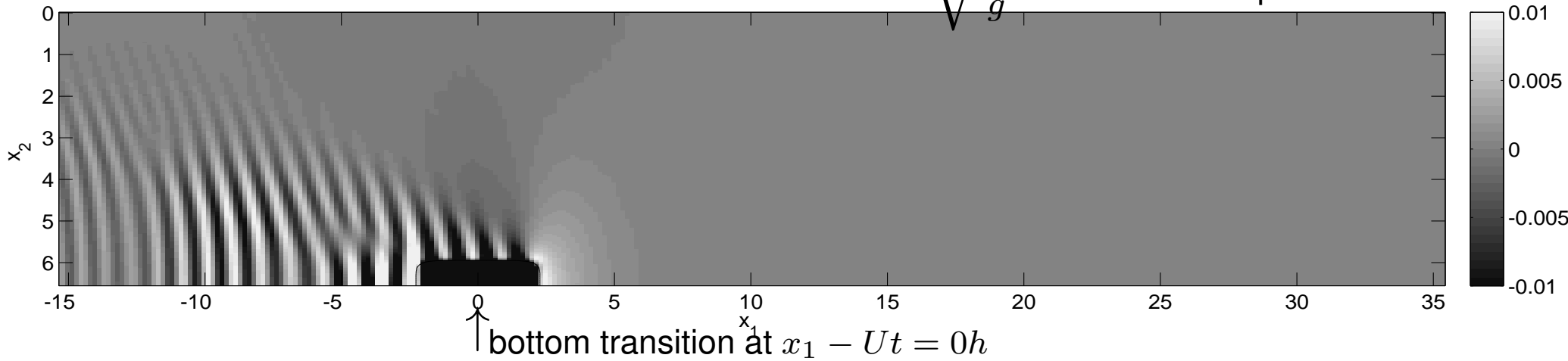
deep2shallow



$$\Delta h/h = 0.909, h = 55 \text{ m}, l_{ship} = 210 \text{ m}$$

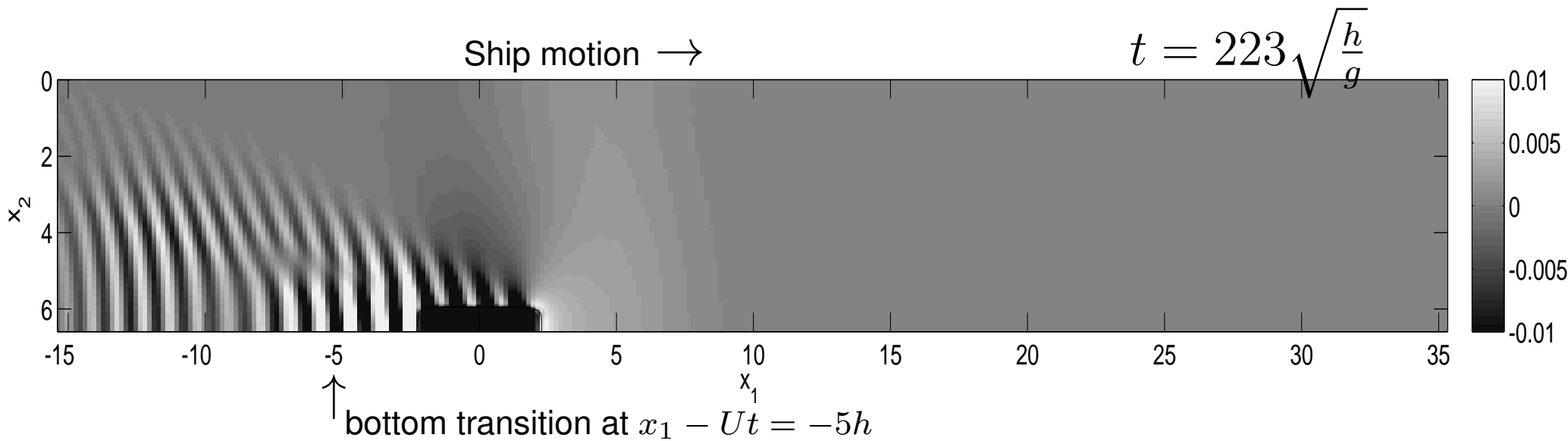
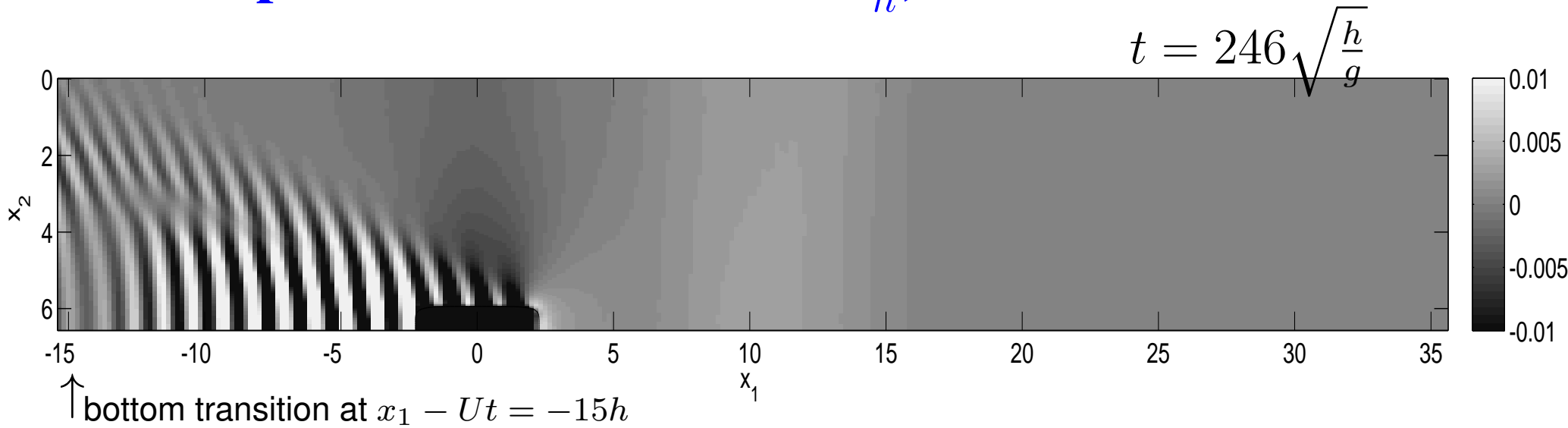
Upstream wave elevation $\frac{\eta}{h}$, $Fr = .43$

$t = 211 \sqrt{\frac{h}{g}}$ water depth $h - \Delta h$



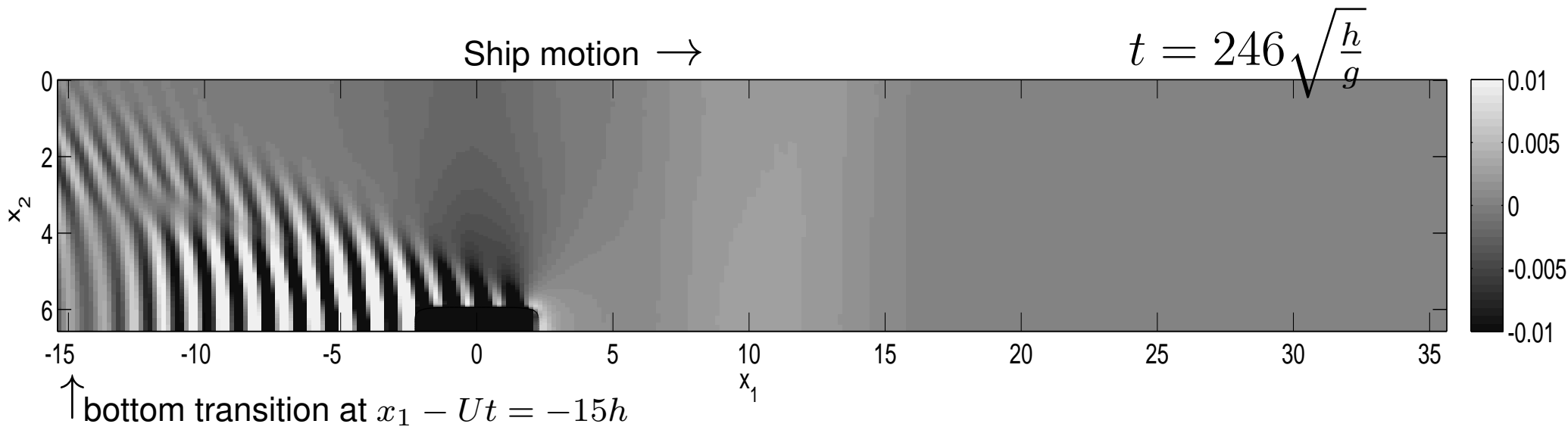
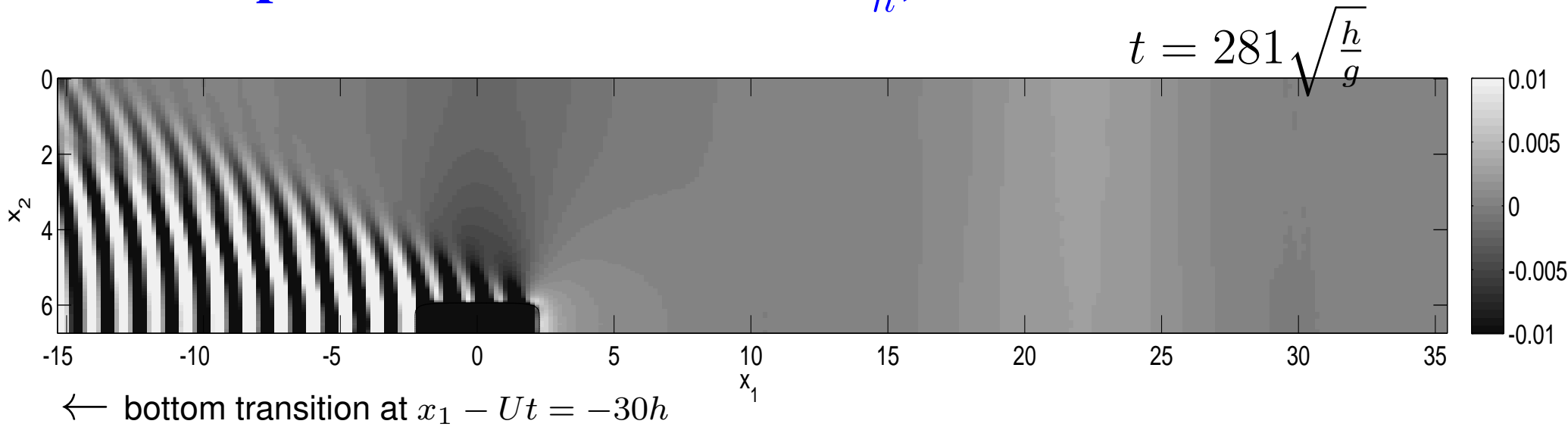
Frame of reference fixed to ship, $\frac{\Delta h}{h} = 0.909$, $h = 55$ m.

Upstream wave elevation $\frac{\eta}{h}$, $Fr = .43$



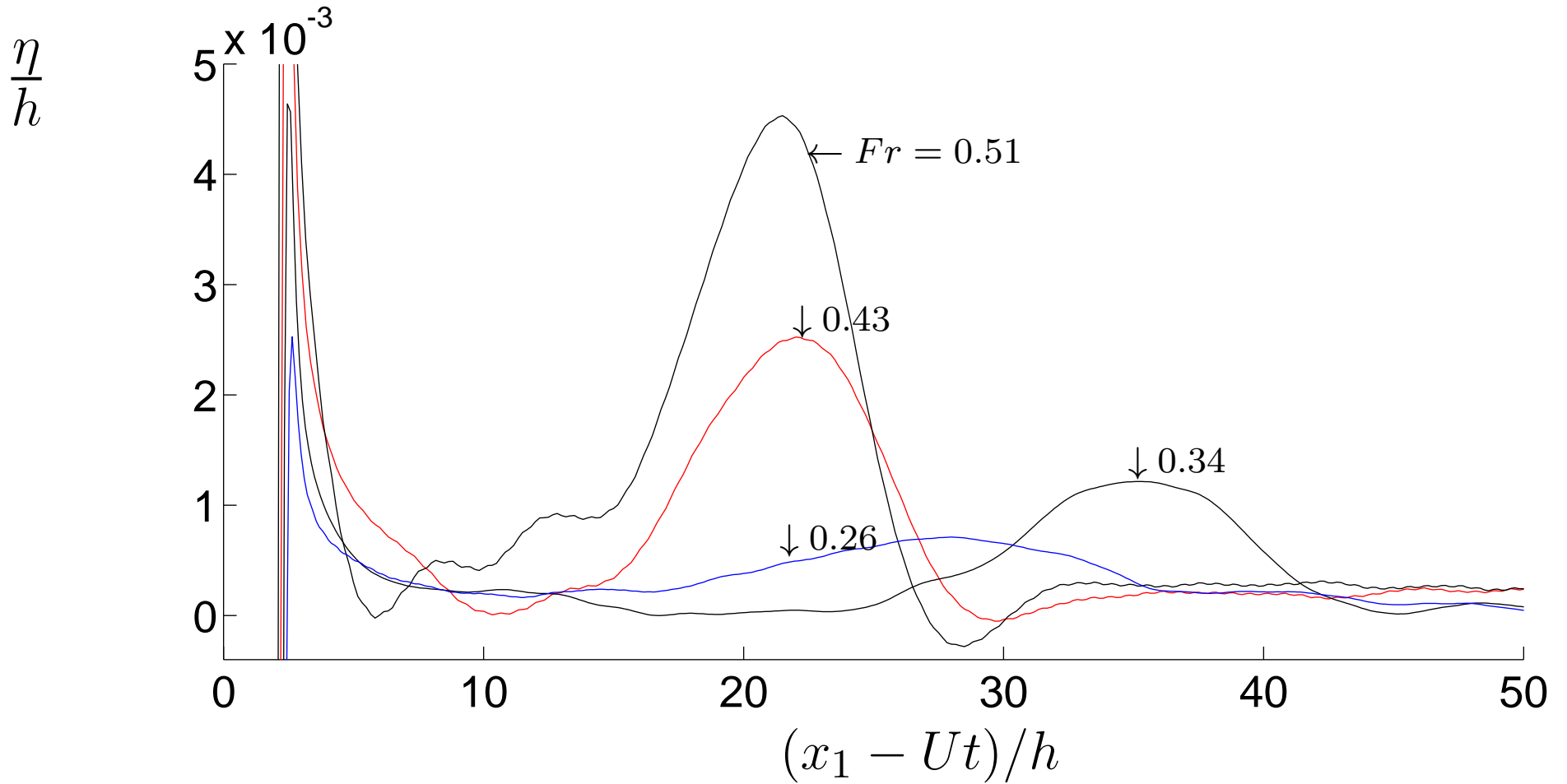
Frame of reference fixed to ship, $\frac{\Delta h}{h} = 0.909$.

Upstream wave elevation $\frac{\eta}{h}$, $Fr = .43$

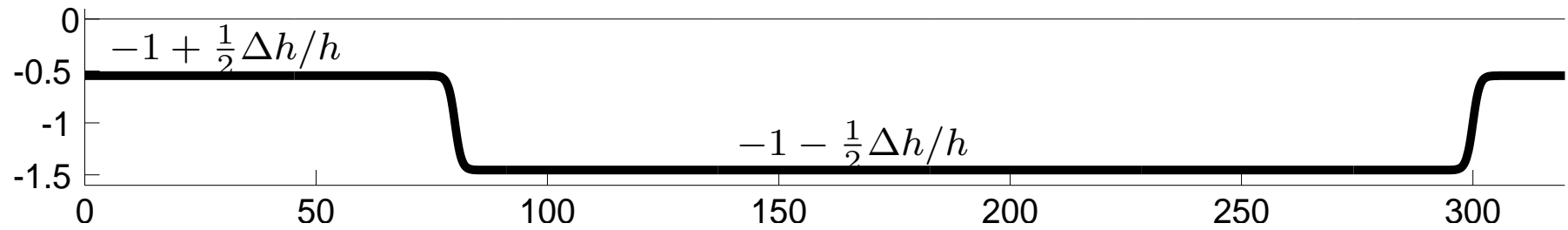


Frame of reference fixed to ship, $\frac{\Delta h}{h} = 0.909$.

Elevation ahead of ship, $\frac{\Delta h}{h} = 0.909$.

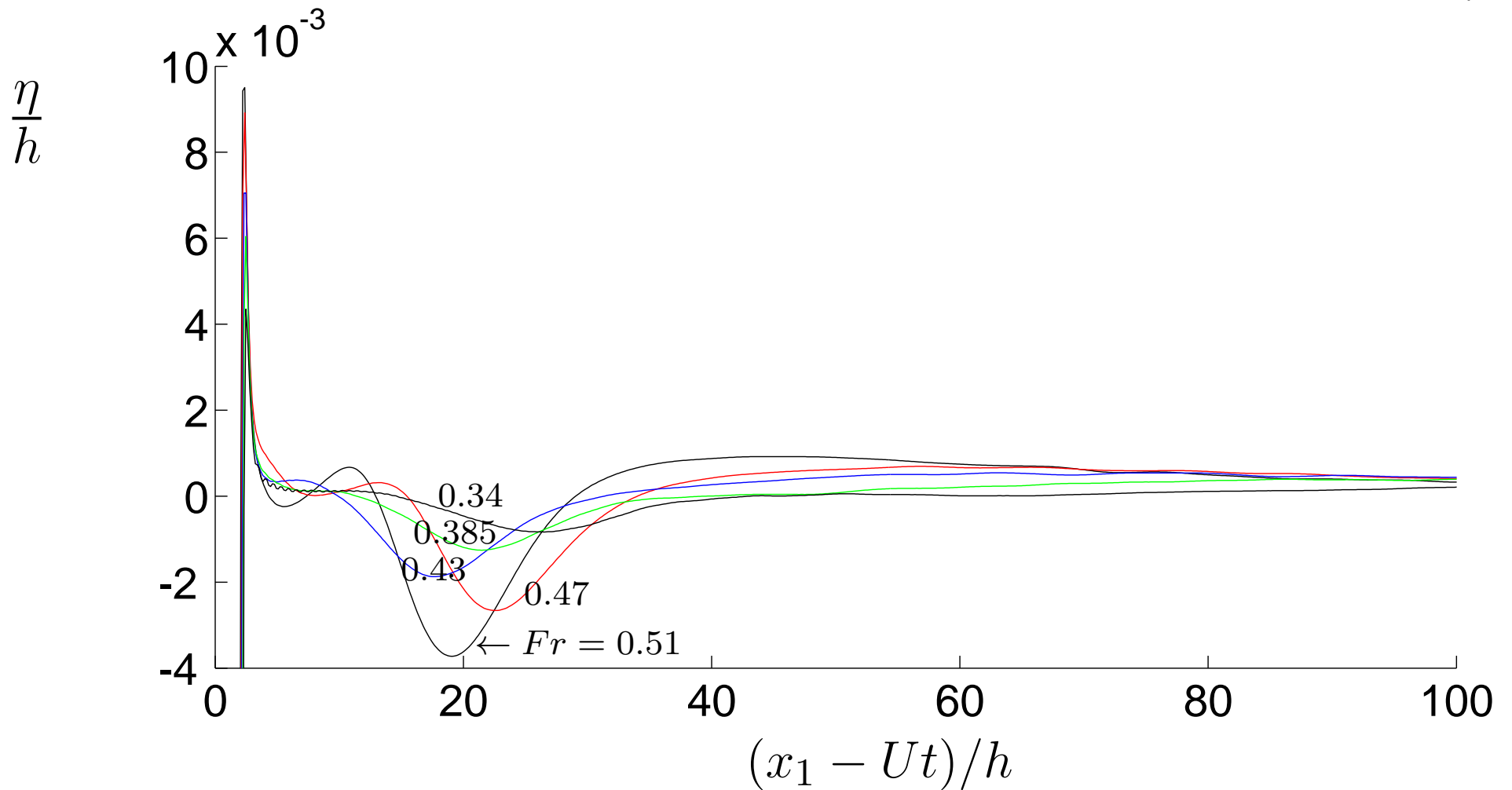


shallow2deep



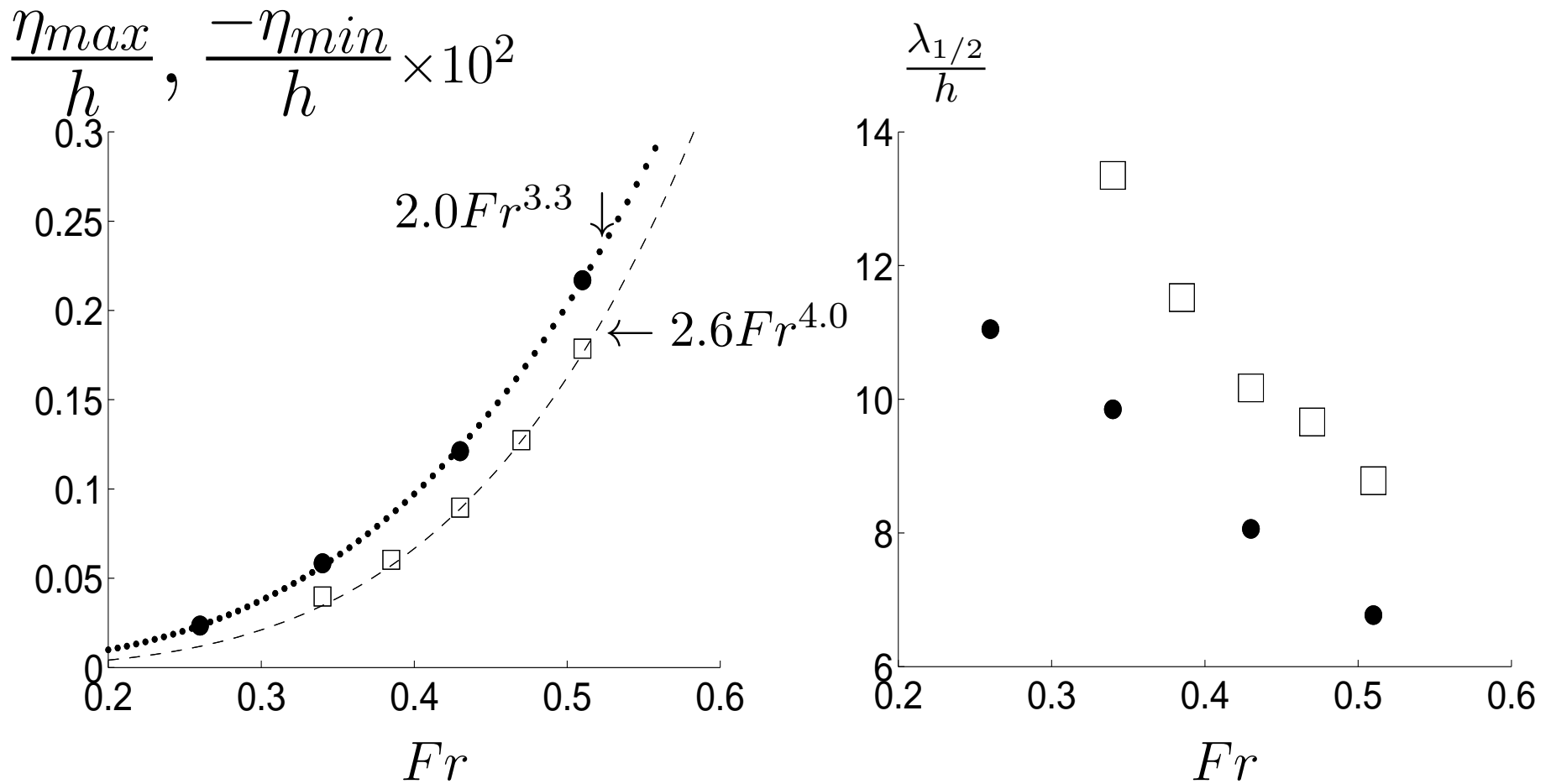
$$\frac{1}{2} \Delta h / h = 0.909, h = 55 \text{ m}, l_{ship} = 210 \text{ m}$$

Depression ahead of ship, $\frac{\Delta h}{h} = 0.909$.



Ship moving from depth $h - \frac{1}{2}\Delta h$ to depth $h + \frac{1}{2}\Delta h$.

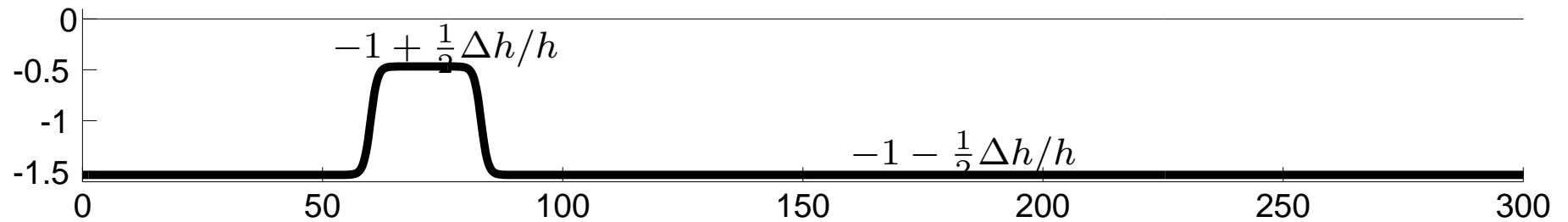
deep2shallow (●); shallow2deep (□)



Elevation/depression (left); wavelength at half height (right) vs. Fr .

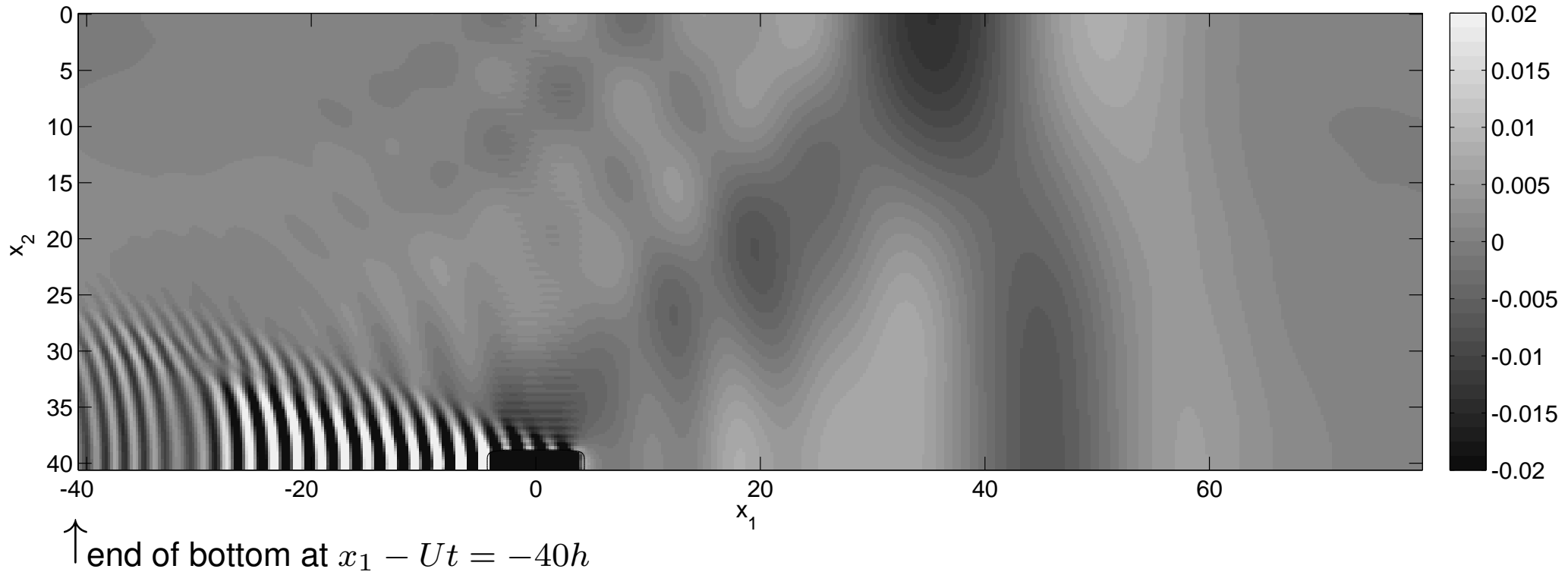
$$yy = 0.042 * fr^{3.3},$$

deep2shallow2deep



$$\Delta h/h = 1.067, h = 30 \text{ m}, l_{ship} = 210 \text{ m}$$

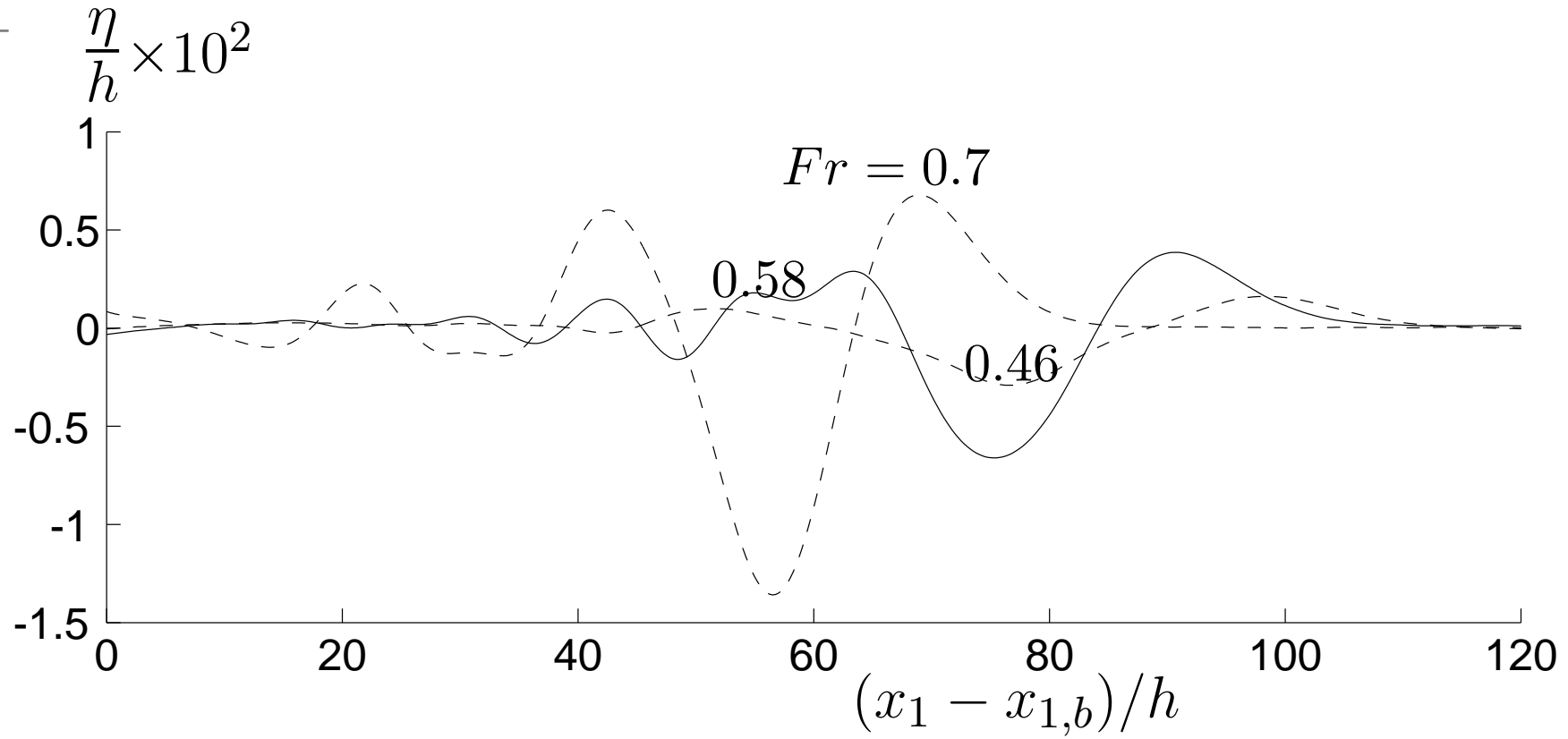
Elevation $\frac{\eta}{h}$, $Fr = 0.58$.

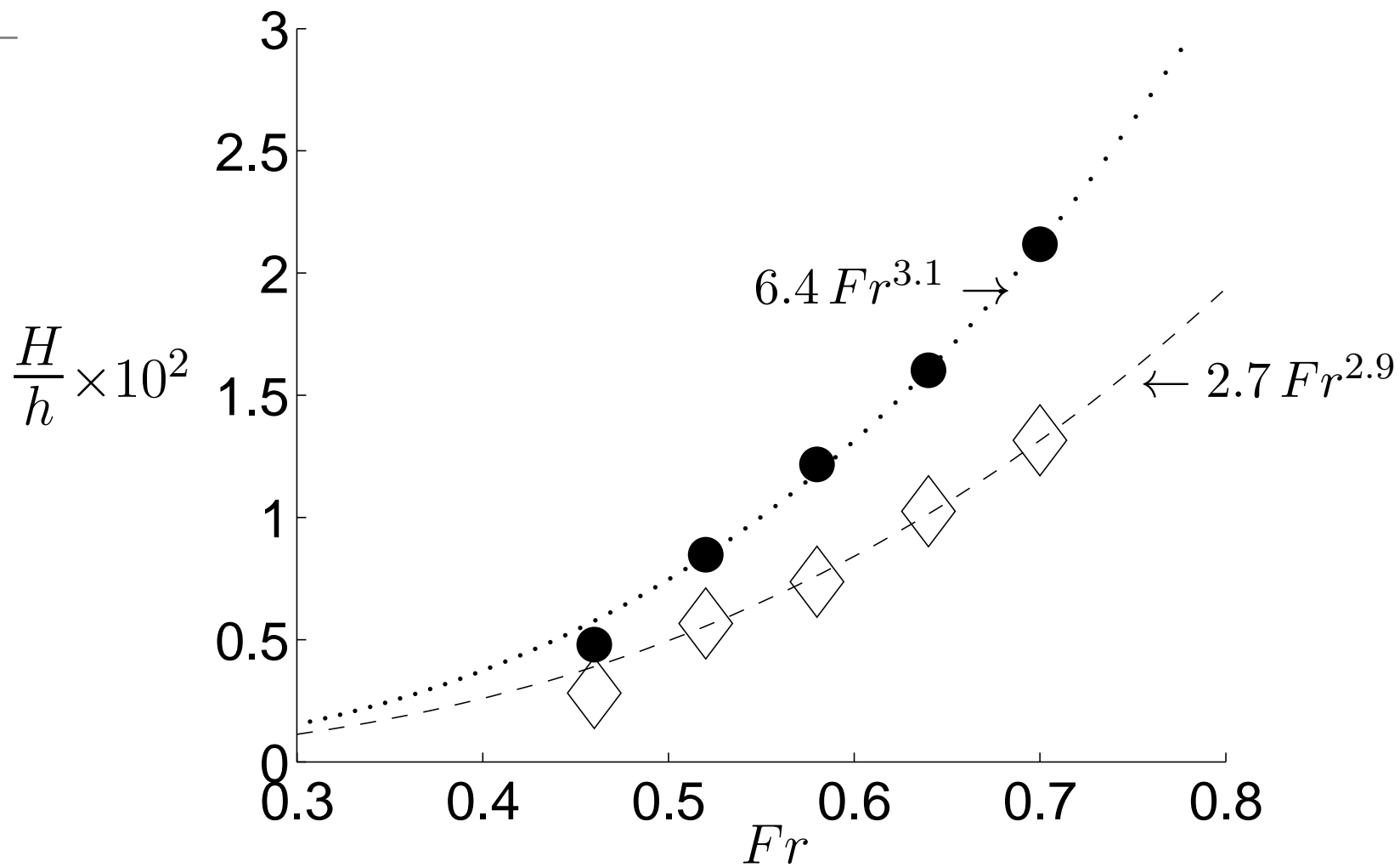


Position of ship: $x_b + 40h$.

$\Delta h/h = 1.067$, $x_a = 60h$, $x_b = 83h$, $L_2 = 80h$, $h = 30$ m.

Elevation at the shore





Max. wave height at channel wall (●) and ahead of ship (◇).

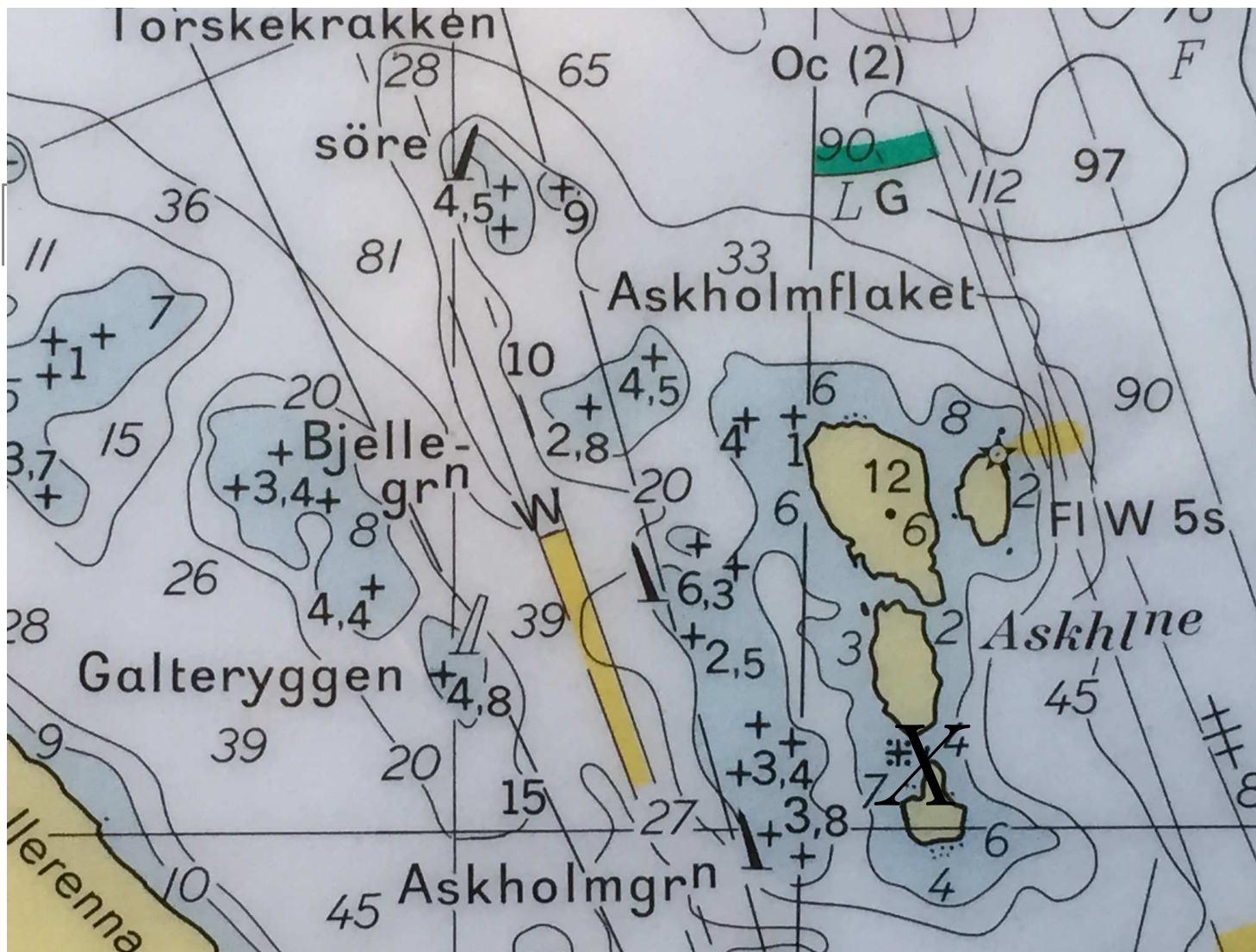
Conclusions

1. A ship at subcritical speed moving across a depth change $\Delta h/h_0 \sim 1$ generates upstream long waves in the form of a small tsunami.
2. An asymptotic analysis provides a qualitative explanation where the upstream waves are expressed in terms of a pressure impulse acting at the depth change, obtaining a leading depression if the depth is reduced, in accordance with observations.
3. A ship moving from deep to shallow to deep water ($\Delta h/h_0 \sim 1$) produces a short upstream wave train: a crest one, **a deep trough**, a second crest, a second trough, and then a train of shorter waves. Reflections at the channel walls magnify the elevation.
4. The wave height ($\sim \frac{1}{2}$ m) grows according to Fr^n , $n \simeq 3$.

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5. More details: Grue (JFM, 2016, in revision).

MOVIE!



The Oslofjord stretches 100 km north-south