

STATISTICAL PREDICTION OF SURFACE WIND COMPONENTS

YIWEN MAO

SUPERVISOR: ADAM MONAHAN

SCHOOL OF EARTH AND OCEAN SCIENCES

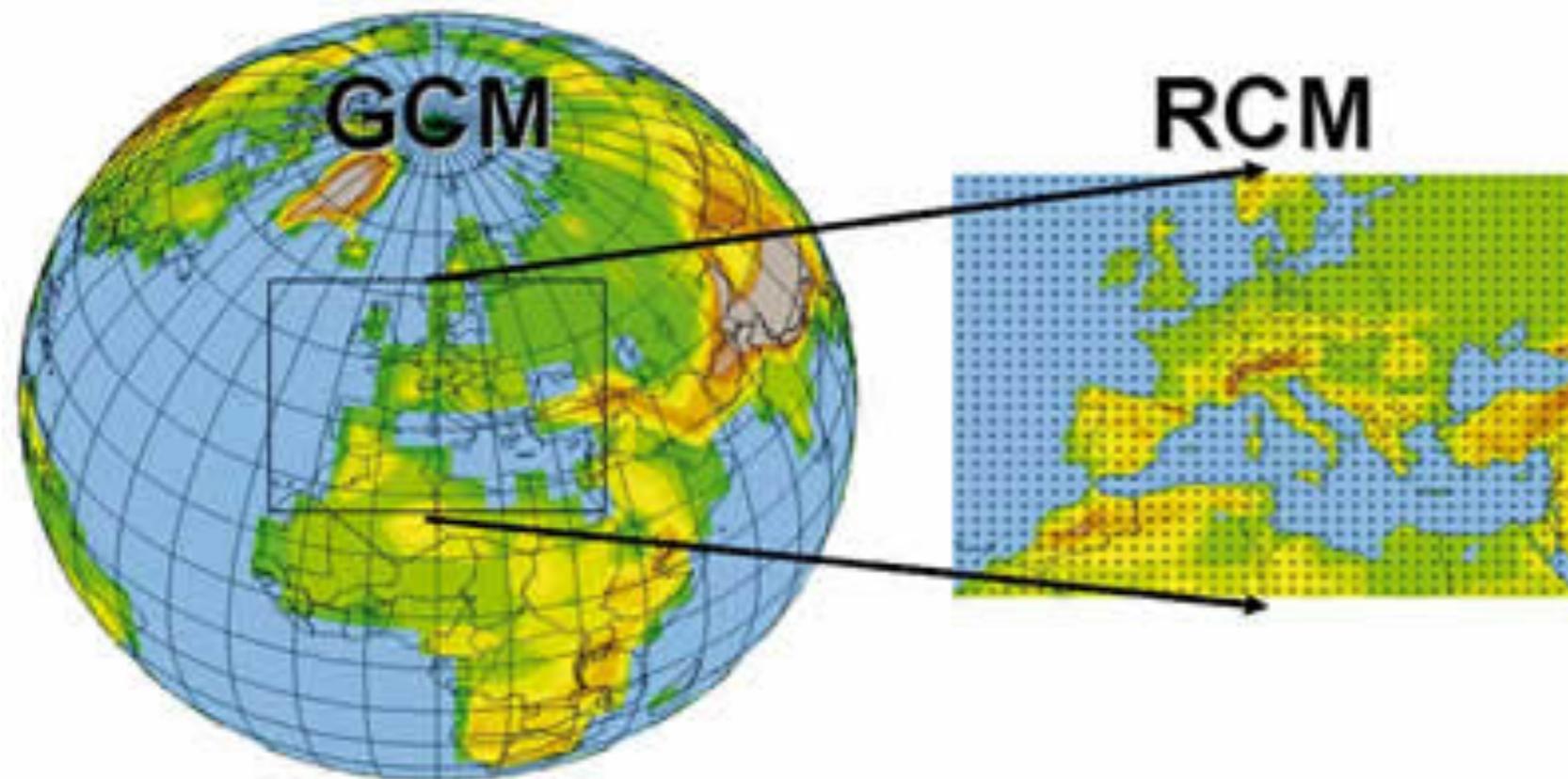
UNIVERSITY OF VICTORIA, BC, CANADA

Introduction

- High-resolution, site-specific **surfaces winds** (e.g. transport of airborne particles, wind energy production)
- Global Climate Models (GCMs): coarse resolution (>100km)  **Cannot model surface winds**
- Downscaling:
 - **Methods** used to infer local-scale climate information (**predictands**) from coarsely resolved climate models (**predictors**)
 - **Dynamical vs Statistical** downscaling

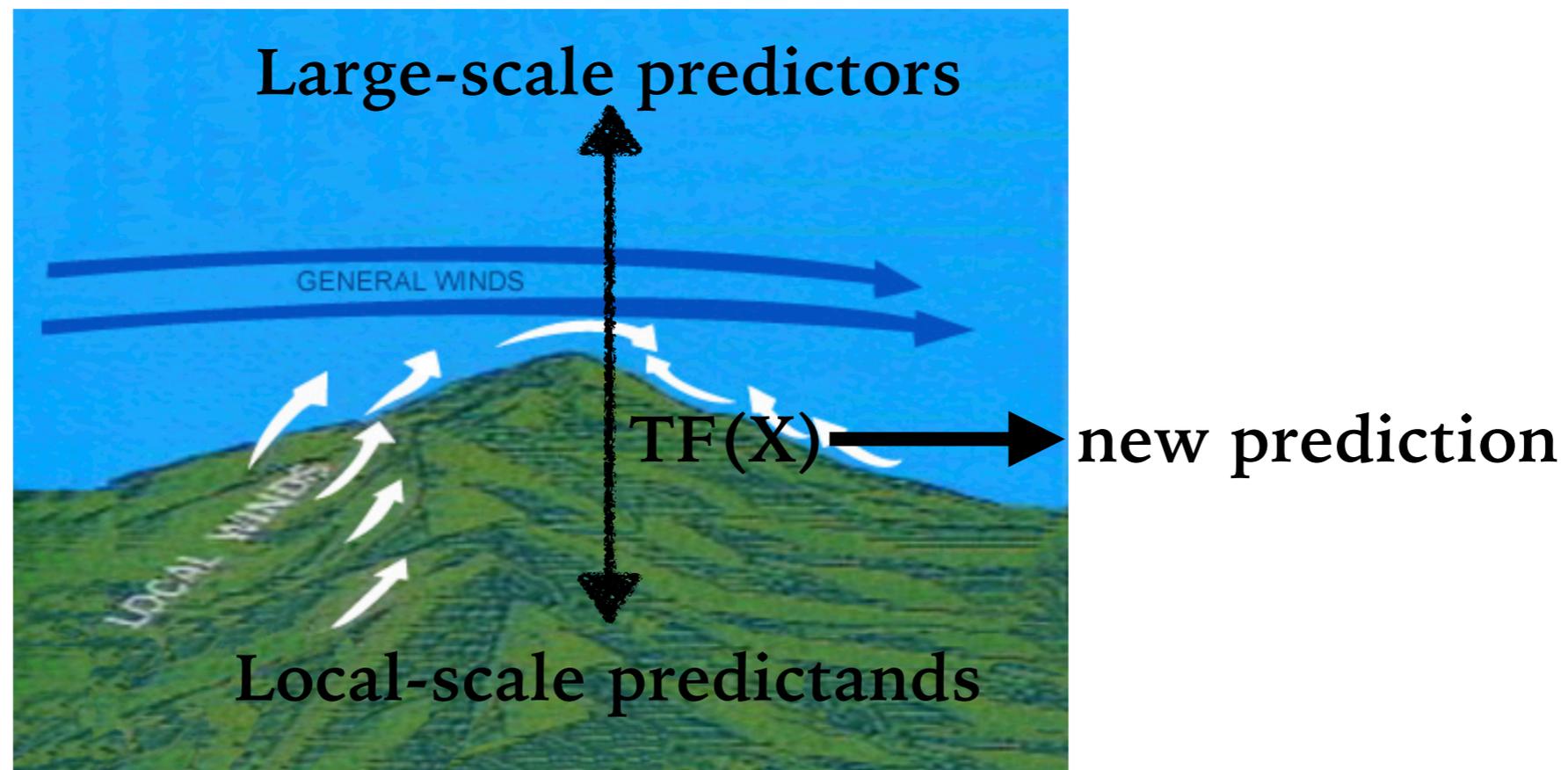
Dynamical downscaling (DD)

- * Nest Regional climate models (RCMs) in the grids of GCMs
- * Physically based ✓
- * Computationally **expensive**



Statistical downscaling (SD)

- * Derive a transfer function (TF) from empirical relationships between predictors and predictands
- * **Flexible** functional form of TFs (**linear, nonlinear**)
- * **Cheap** computational cost ✓



* DD and SD: Comparative skills in predicting historical data

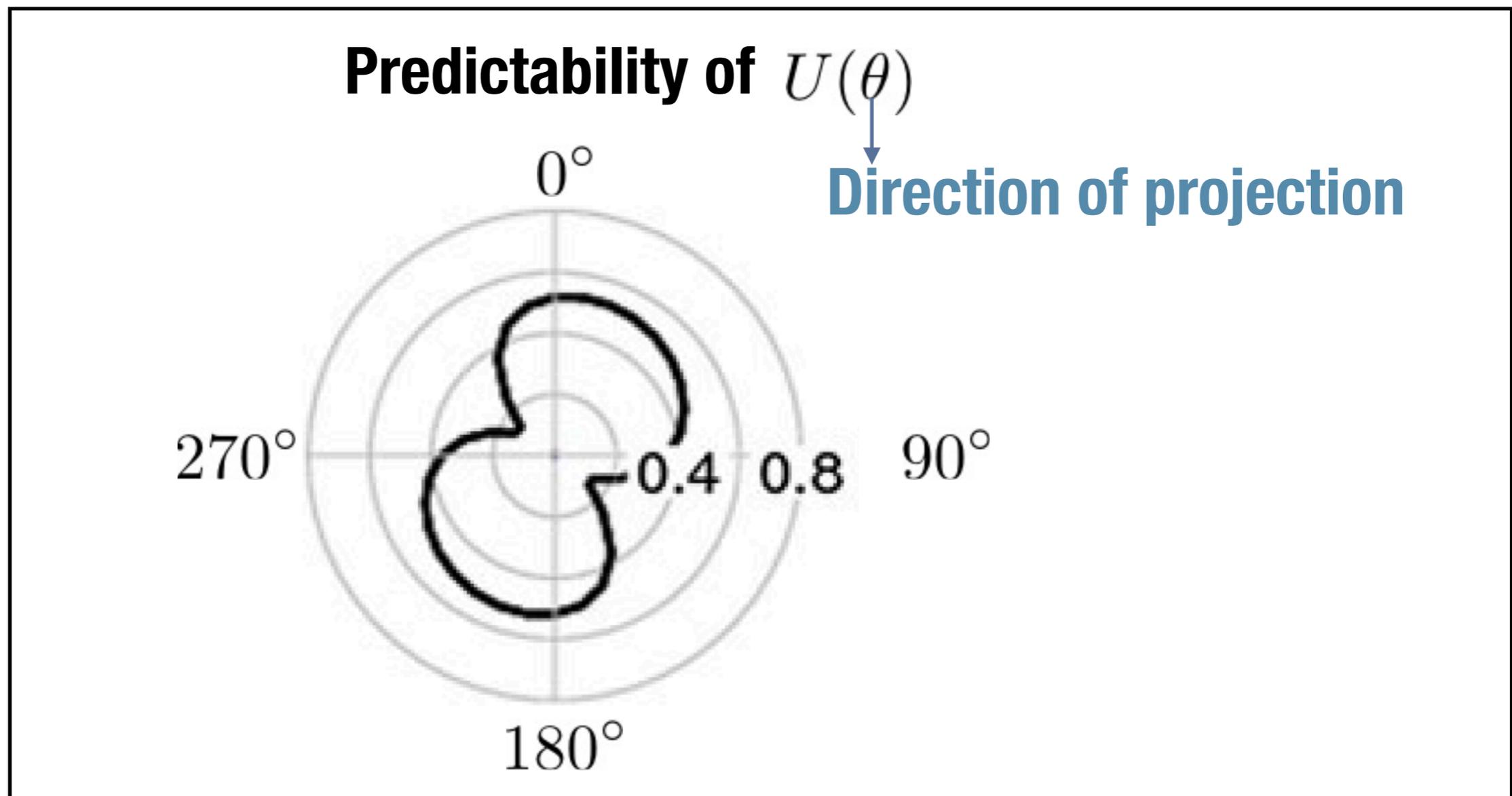


* SD: cheap computational cost ★

**SD: most commonly applied in predicting scalar variables
(Temperature, Precipitation)**

Research Focus

- * Assess how well SD can do in predicting surface wind components (**vector**)  **Directional Characteristics**
- * **Predictive anisotropy:** predictability of surface wind components varies with the direction of projection.



Research Objectives

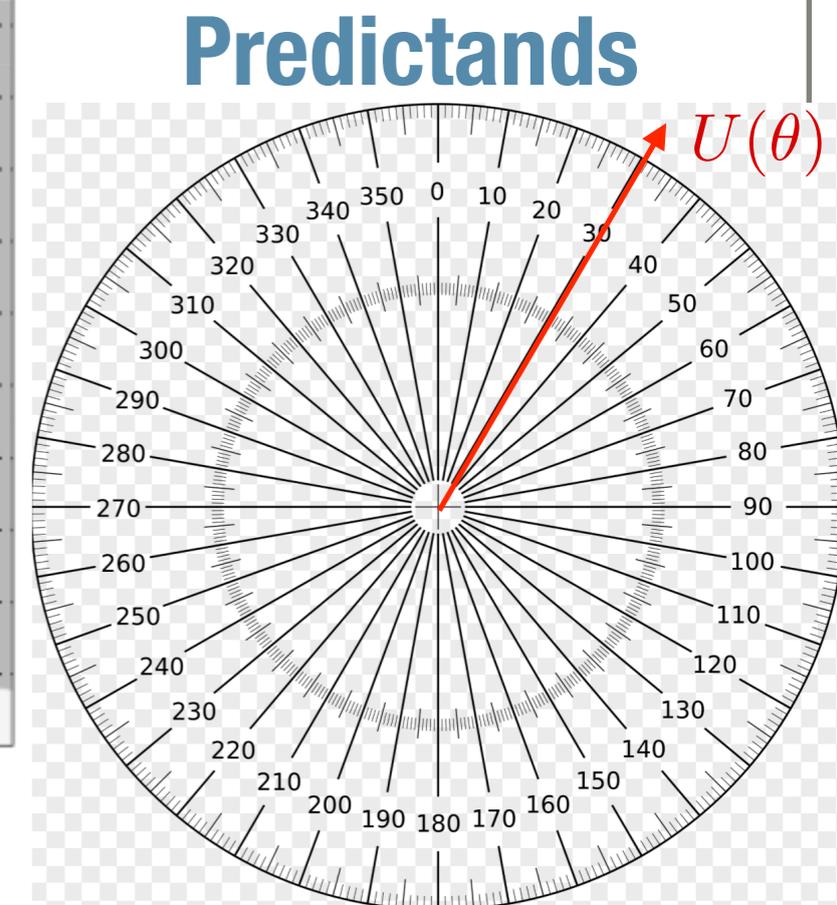
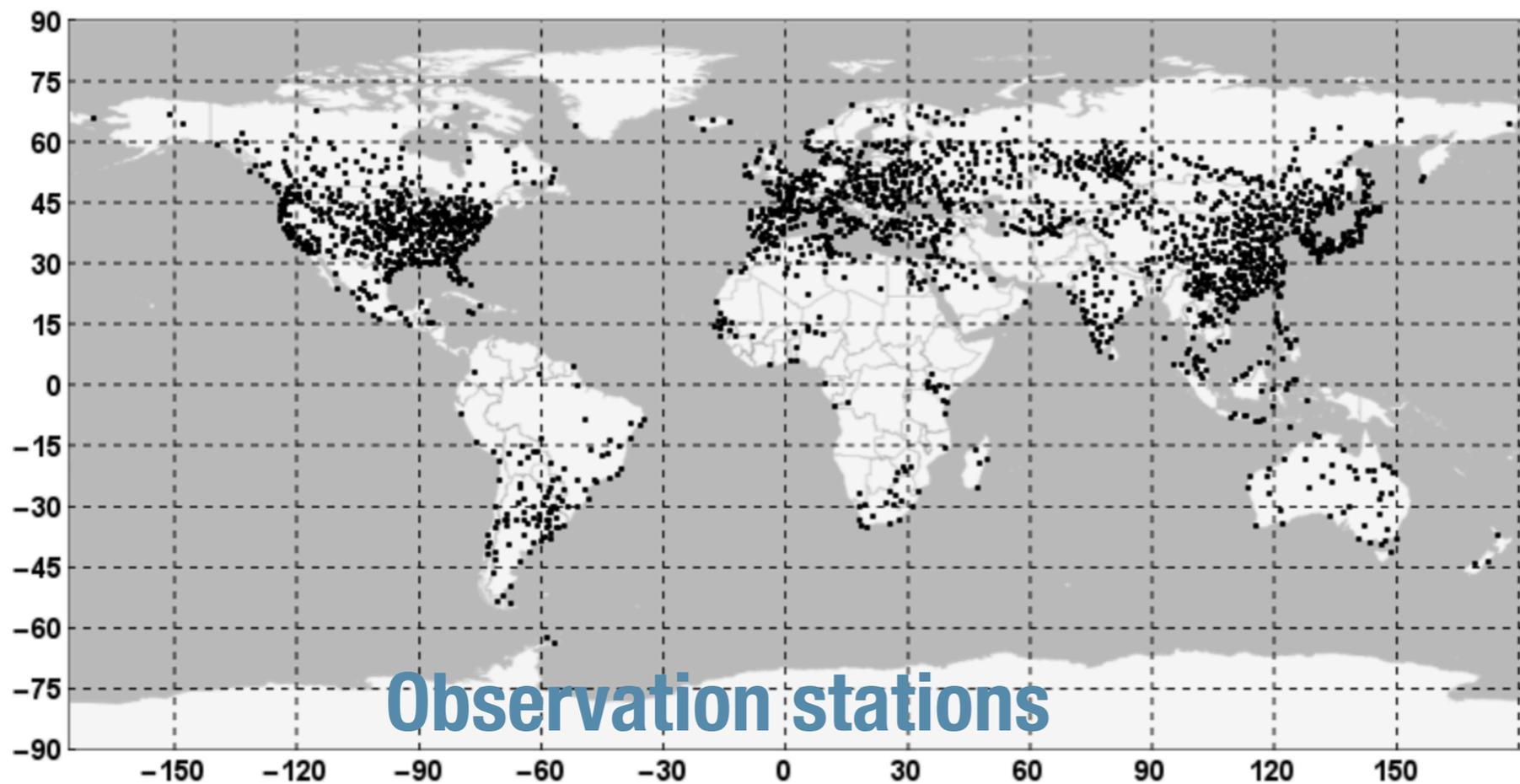
1. to provide a **global characterization** of statistical predictability of surface wind components
2. to compare the **efficiency** of **linear** and **nonlinear** TFs
3. to build **a general framework** to explain characteristics of statistical predictability with an emphasis on **predictive anisotropy (contributing factors)**



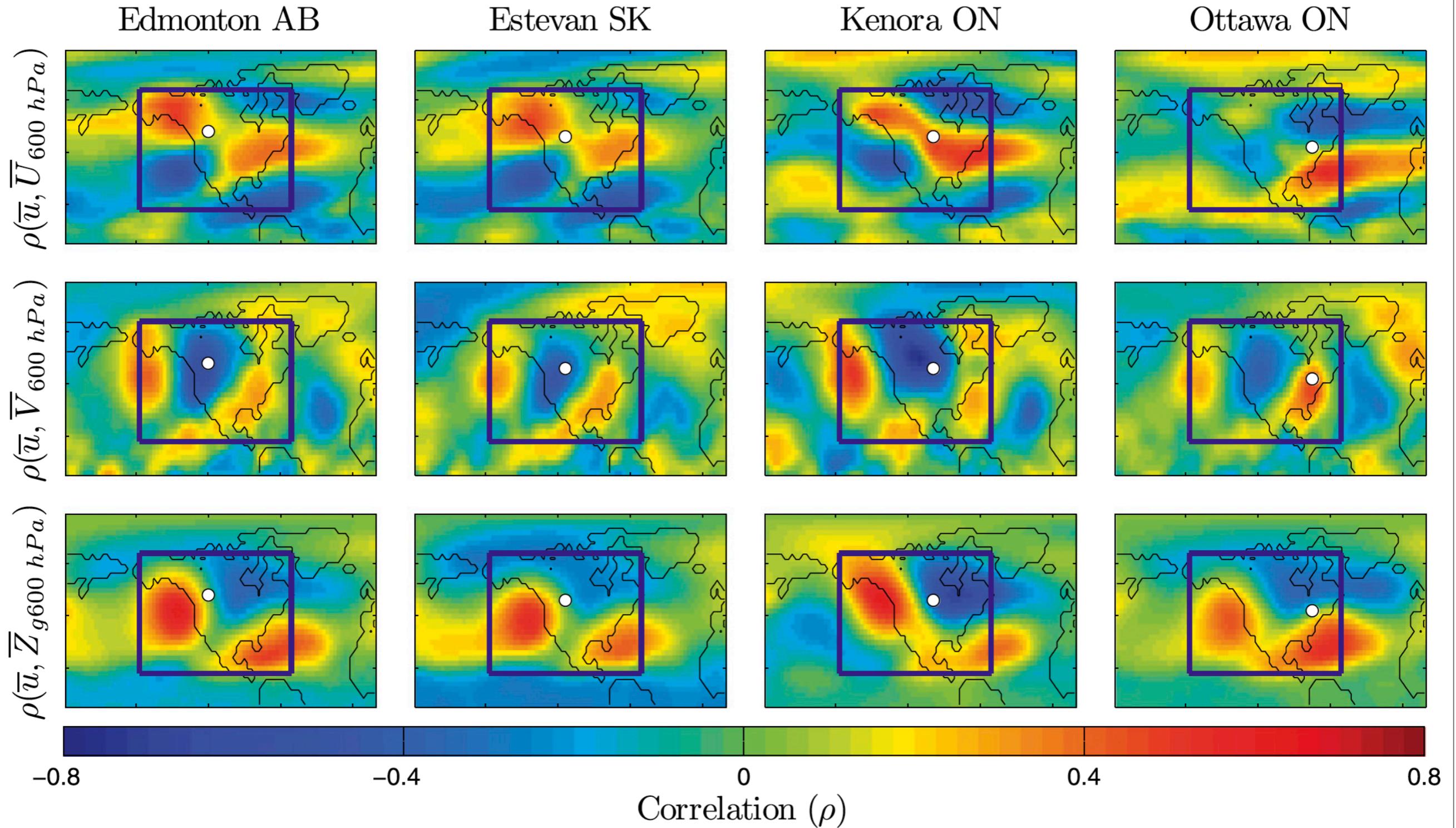
- **wrong functional form of TFs (i.e. linear)?**
- **physical factors?**

Methodology

- * **Predictands:** surface wind components projected onto 0, 10, 20,...360 deg at 2109 land stations
- * **Predictors:** Temperature (T), Geopotential height (Z), zonal (U), meridional (V) wind components at 500 mb from NECP2 reanalysis
- * **Prediction period:** 1980-2012, Summer/Winter, Daily/Monthly
- * **Predictability:** $R^2 = \text{corr}^2(\text{Obs}, \text{Pred})$

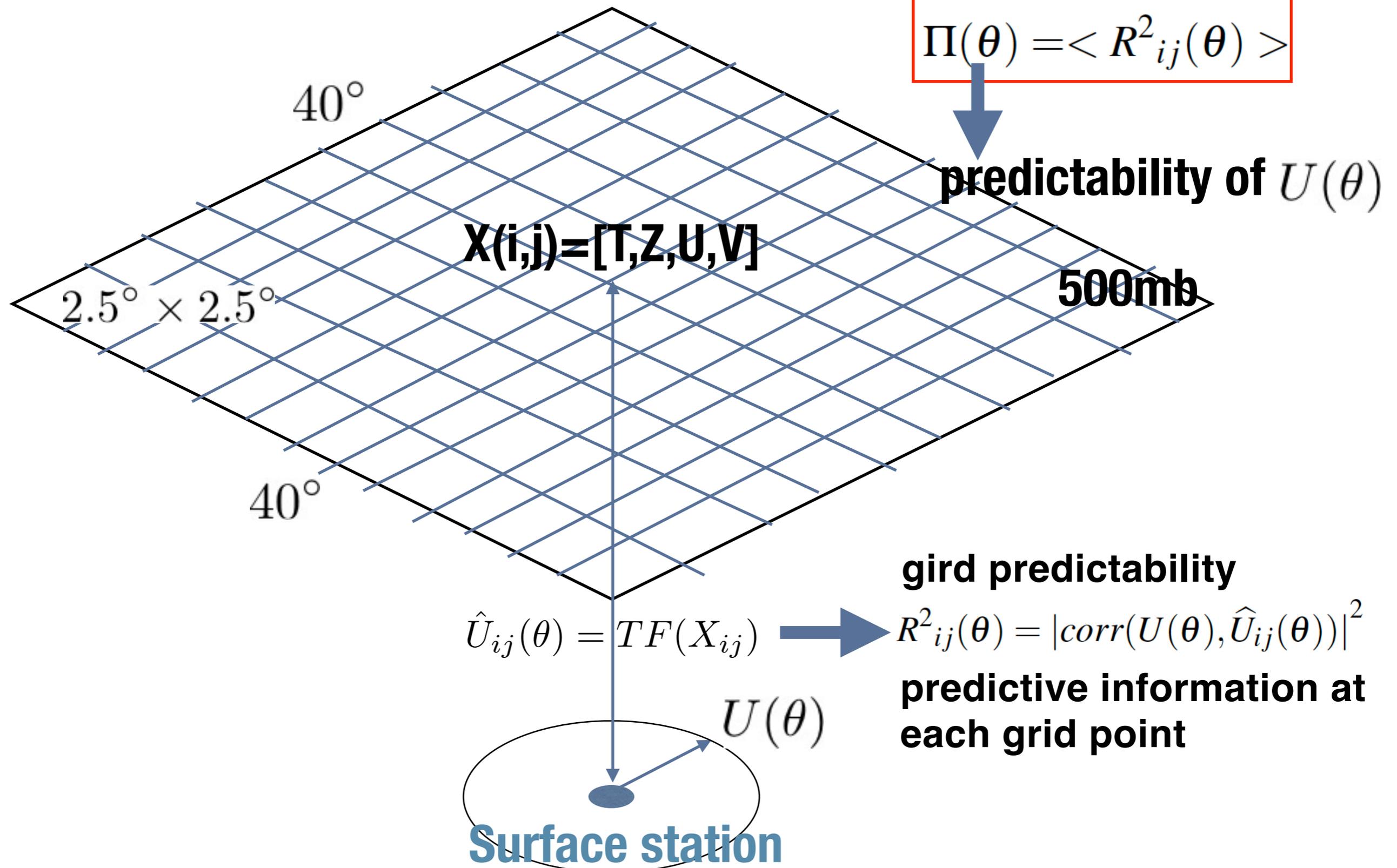


Predictive Information



Culver AM, Monahan AH. The statistical predictability of surface winds over western and central Canada. *Journal of Climate*. 2013 Nov;26(21):8305-22.

Prediction at each station



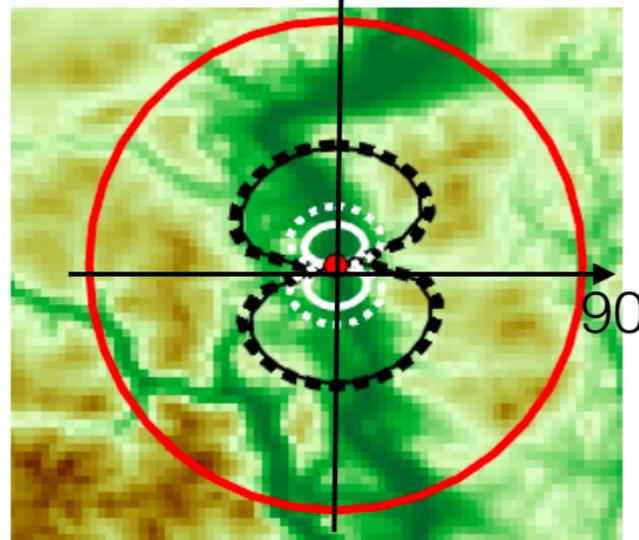
TF: linear regression

Metrics of Predictability: $\min(\Pi)$, $\max(\Pi)$, $\alpha(\Pi)$

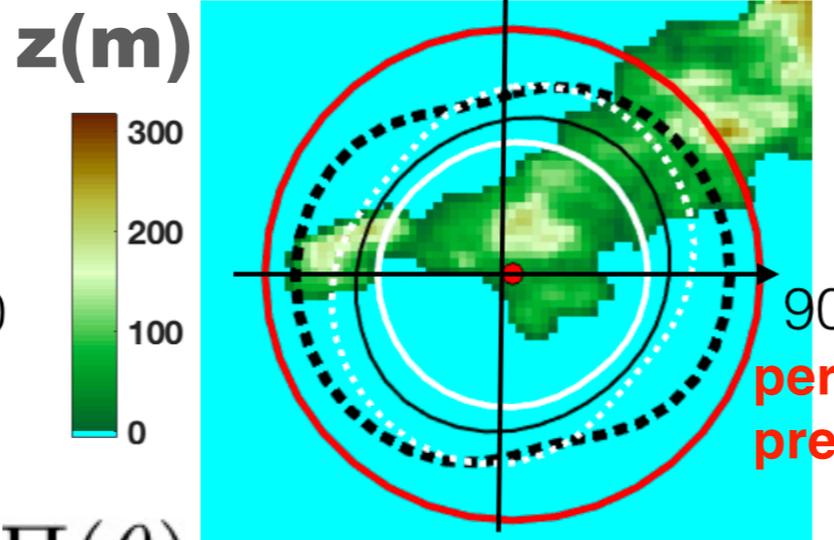
Predictive anisotropy

$$\alpha(\Pi) = \frac{\min(\Pi)}{\max(\Pi)}$$

Strong $\alpha \rightarrow 0$



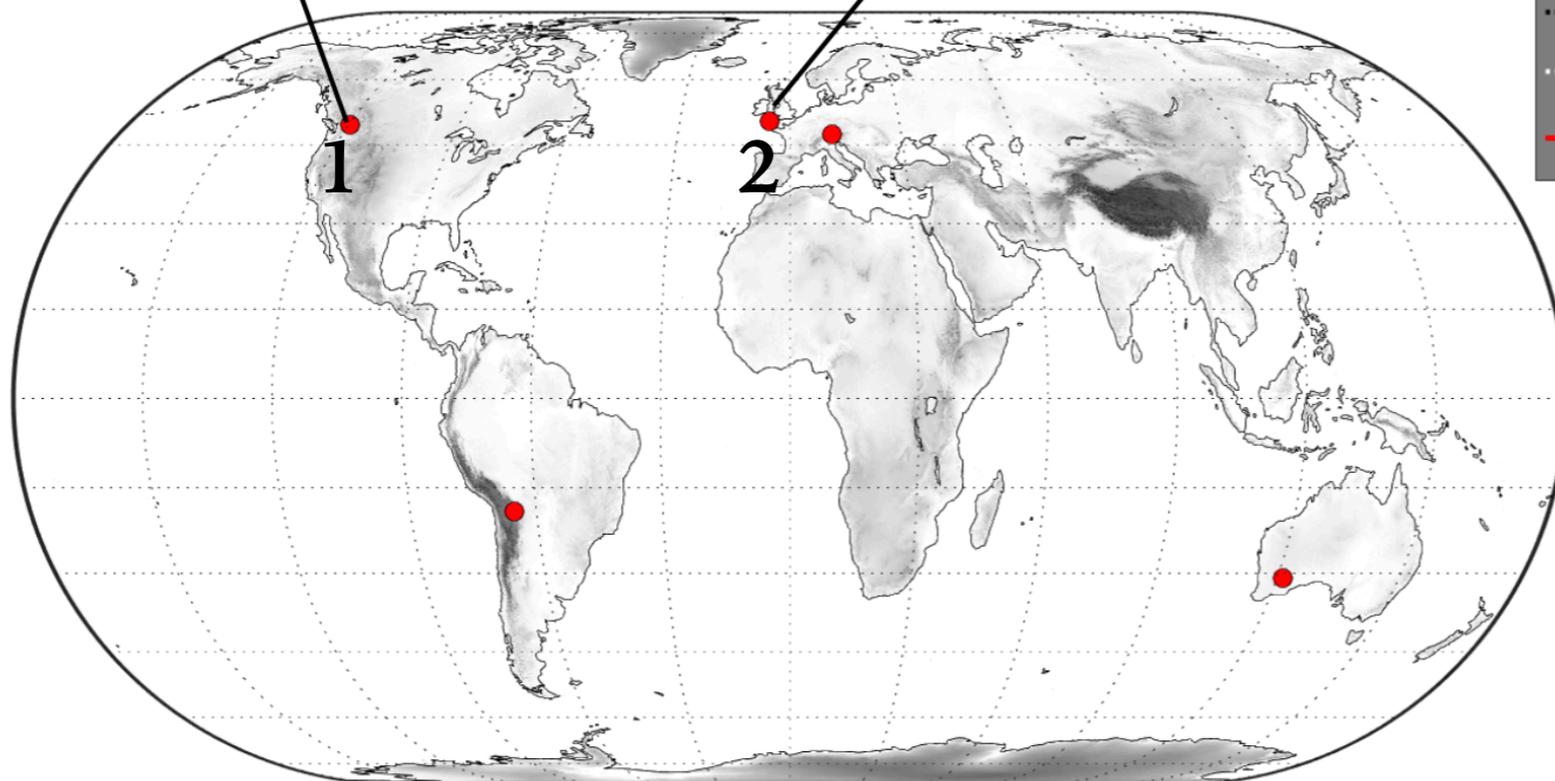
Weak $\alpha \rightarrow 1$



perfect prediction

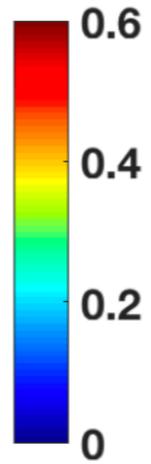
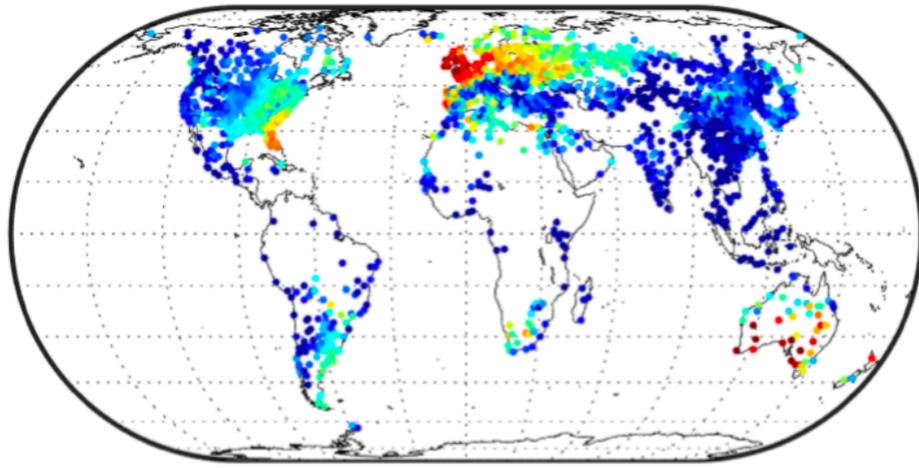
$\Pi(\theta)$

- Daily winter
- Daily summer
- Monthly winter
- Monthly summer
- $R^2=1$

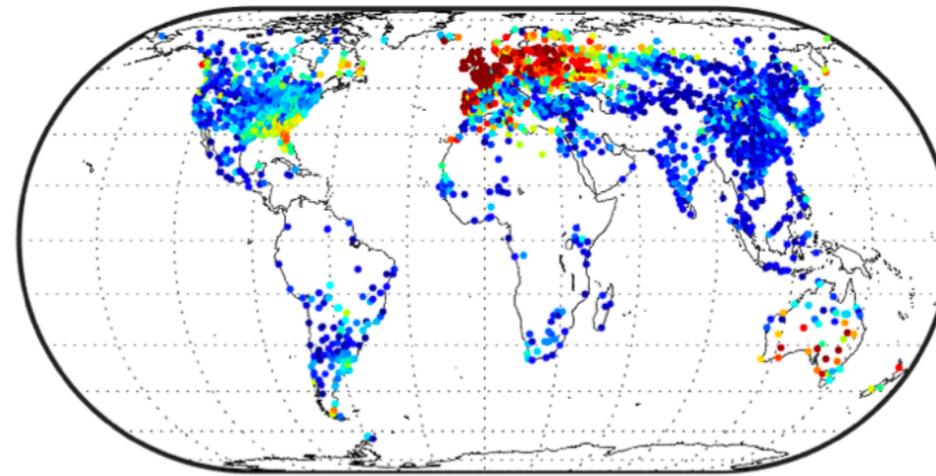


Predictability

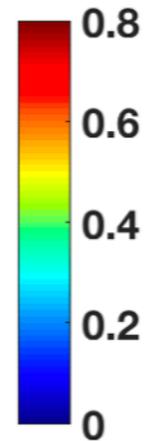
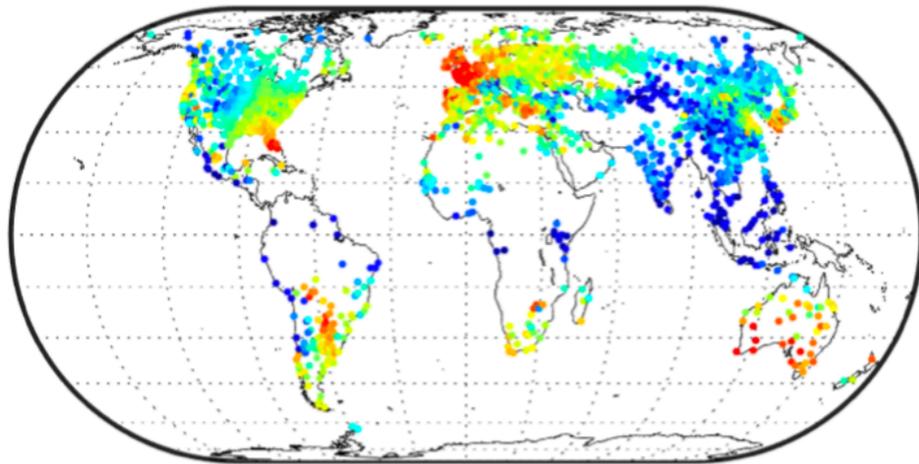
Daily winter min(Π)



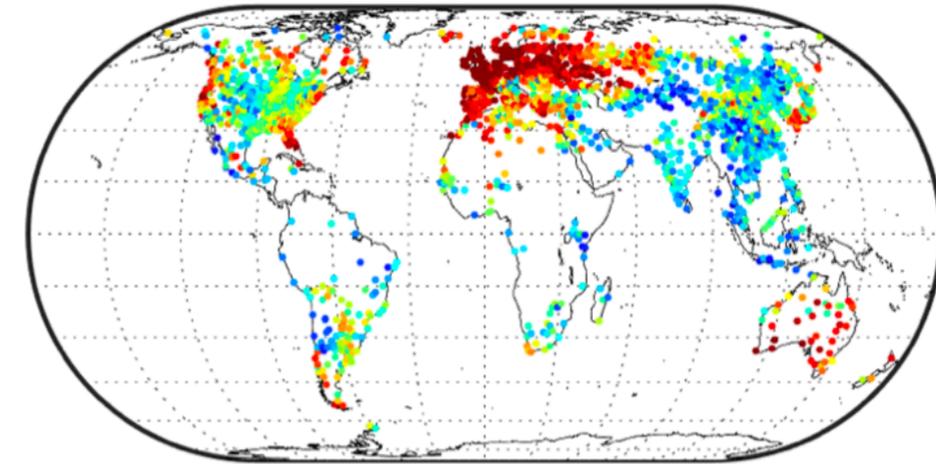
Monthly winter min(Π)



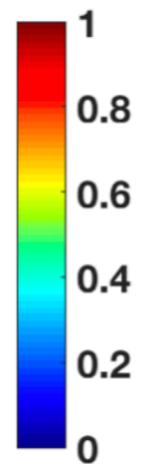
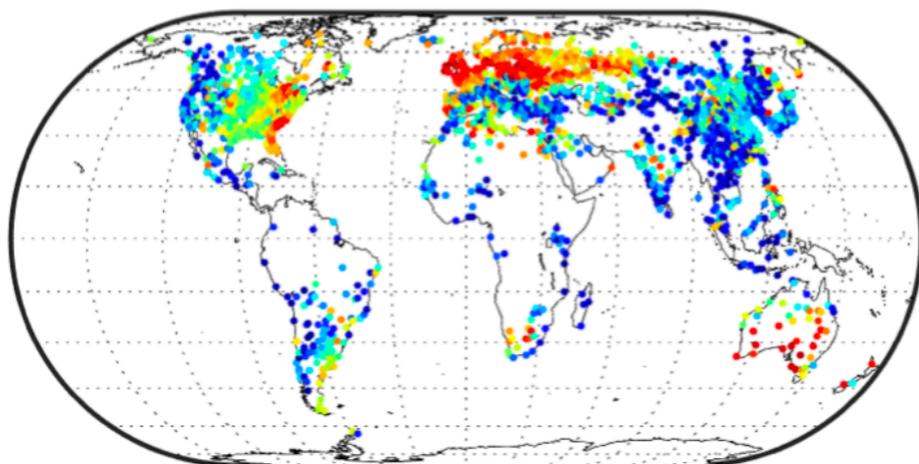
Daily winter max(Π)



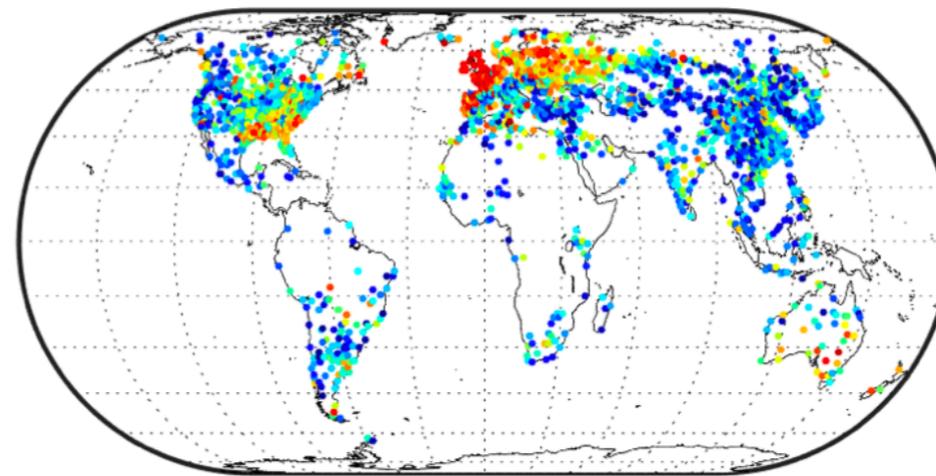
Monthly winter max(Π)



Daily winter $\alpha(\Pi)$



Monthly winter $\alpha(\Pi)$



Predictive anisotropy

Potential factors that can influence predictability

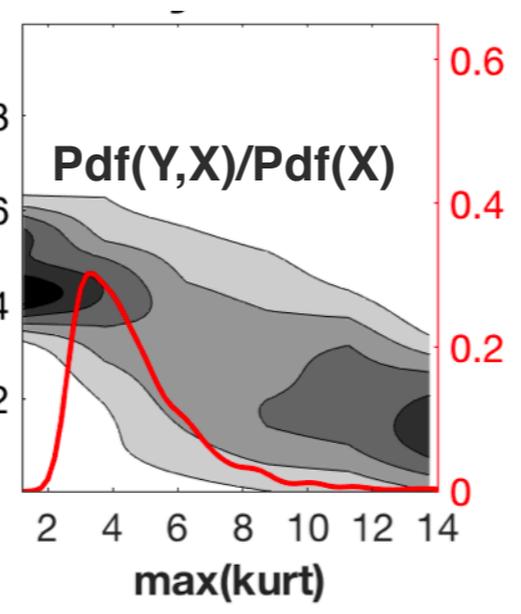
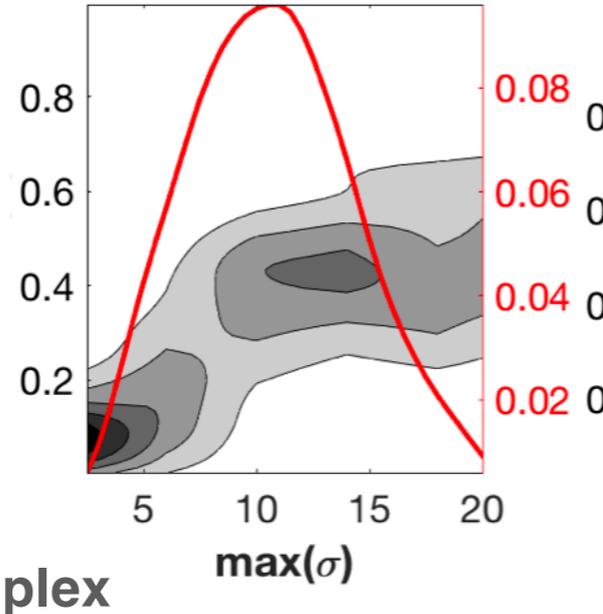
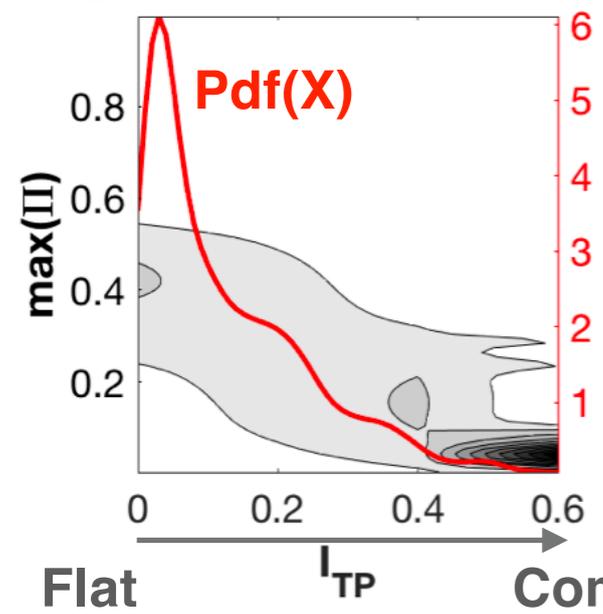
- * Topographic complexity
- * Variability and shape of fluctuations of surface wind components

Topographic Complexity

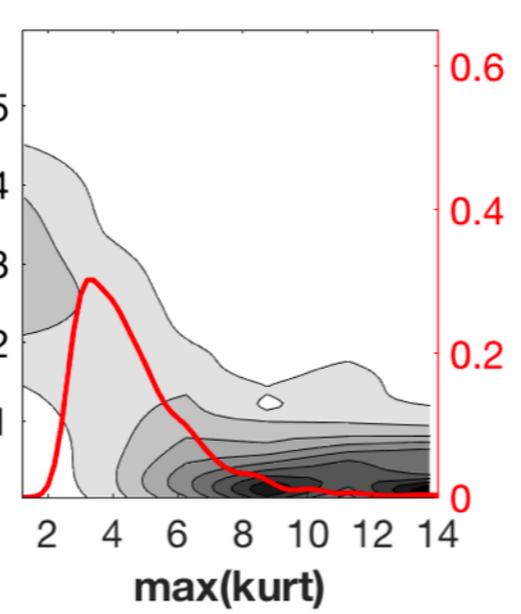
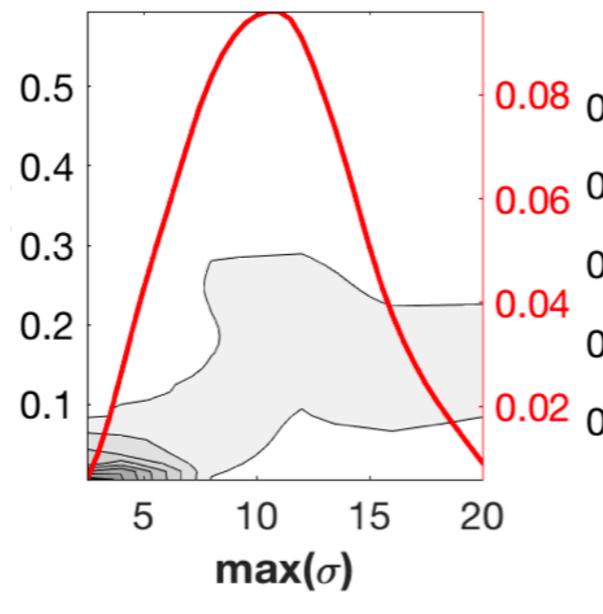
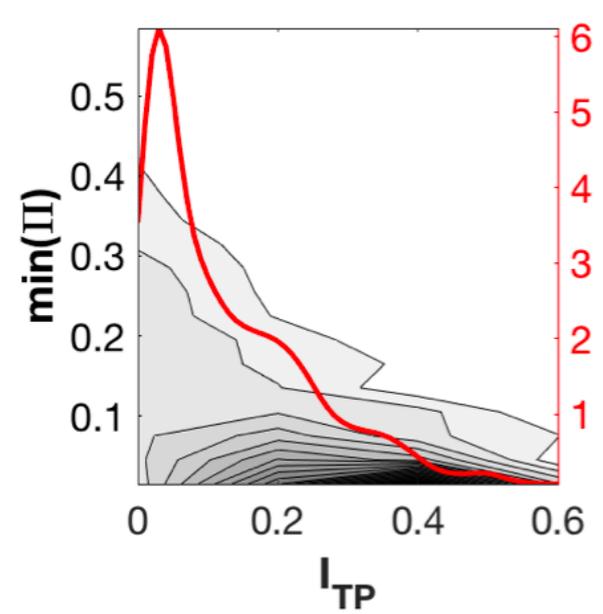
Std of U

Kurtosis of U

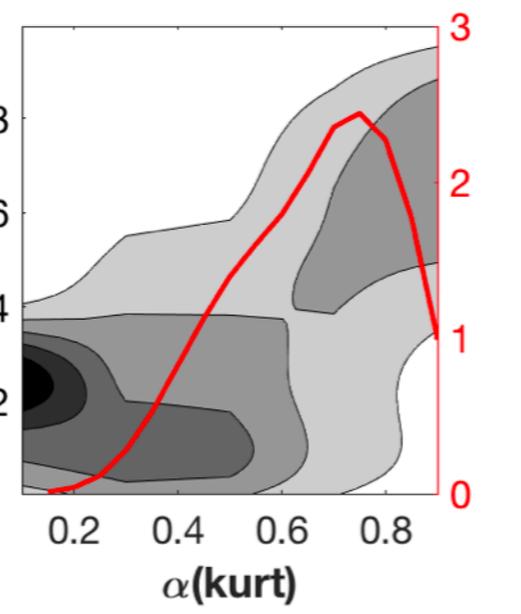
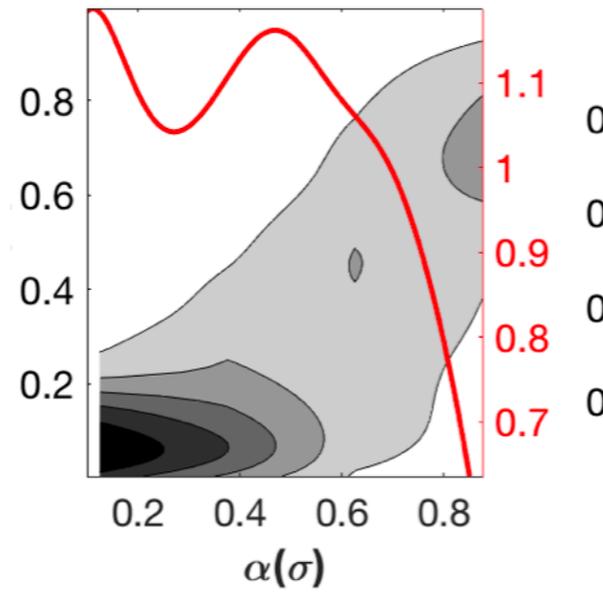
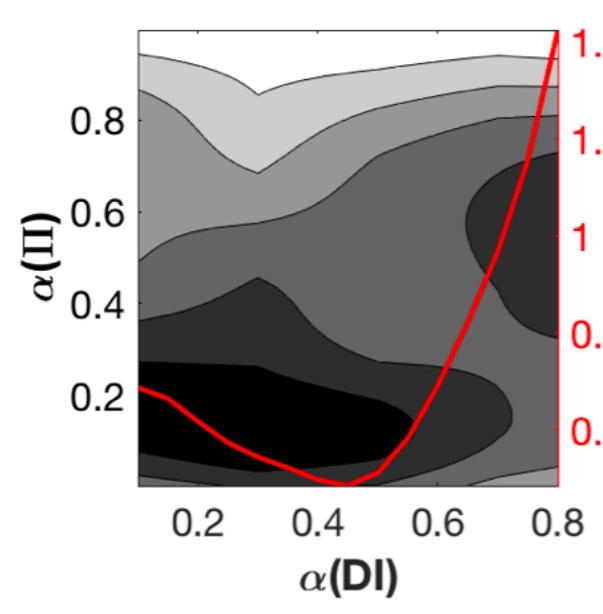
Daily Winter



More/less complex terrain
 less/more variable U
 non-Gaussian/near-Gaussian



Low/High Predictability



directional variation of Std, Kurtosis of U

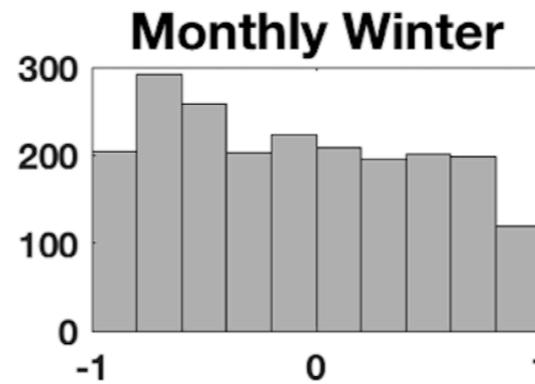
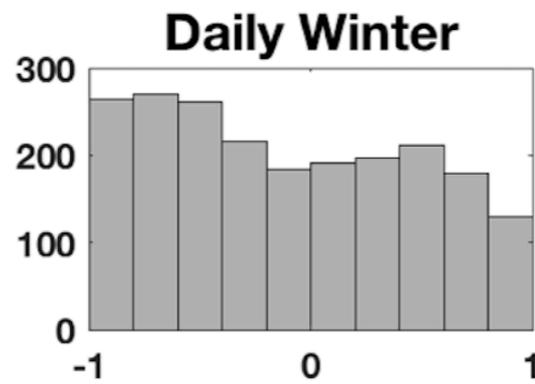
Predictive anisotropy

Weak

directional variation of terrain

$$\rho(\Pi(\theta), DI(\theta))$$

Terrain

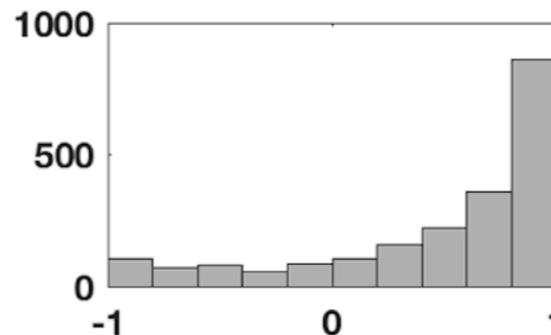
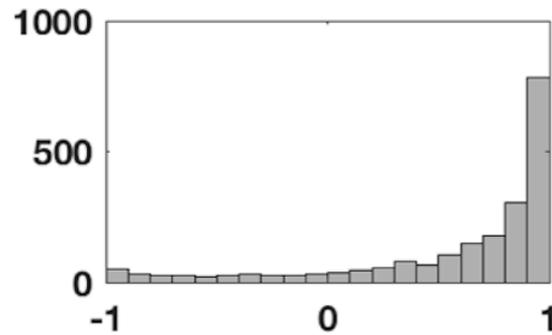


direction of **Low/High** Predictability



$$\rho(\Pi(\theta), \sigma(\theta))$$

U Variability

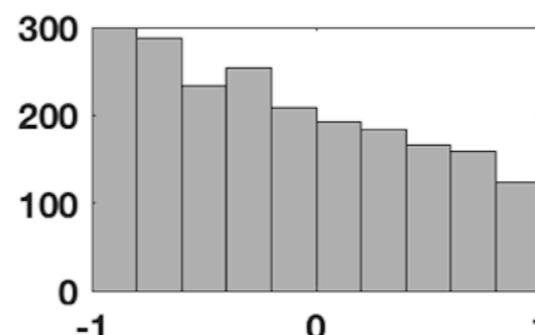
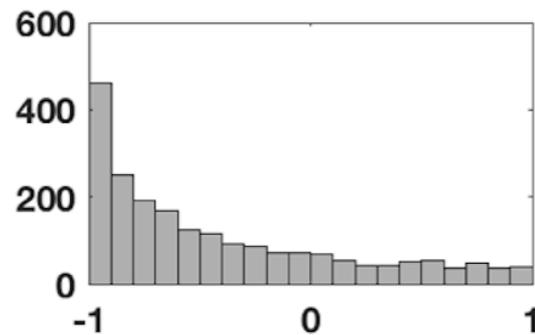


direction of

- **more/less** complex terrain
- **less/more** variable fluctuation
- **non-Gaussian/near-Gaussian**

$$\rho(\Pi(\theta), kurt(\theta))$$

U Distribution



Rank correlation coefficient

$\rho \rightarrow 1$ Directional maxima (minima) tend to align

$\rho \rightarrow -1$ Directional maxima (minima) tend to be orthogonal

$\rho \rightarrow 0$ No directional relationships

Idealized model

$$U(\theta) = \text{Noise} + \text{Signal}$$

- zero linear correlation with predictors
- Isotropic
- non-Gaussian

- Perfect linear correlation with predictors
- vary with direction
- Gaussian

$$\Pi(\theta) = \text{corr}^2(U(\theta), \text{Signal})$$

Idealized model

$$U(\theta) = U_x + g(\theta)U_y,$$

Noise

Signal

$$\Pi(\theta) = \text{corr}^2(U, U_y) = \frac{G(\theta)}{\beta + G(\theta)}$$

$$\sigma(\theta) = \sqrt{V_x + V_y G(\theta)},$$

$$\text{kurt}(\theta) = \frac{K_x \beta^2 + G(\theta)^2 K_y + 6G(\theta)\beta}{\beta^2 + 2G(\theta)\beta + G(\theta)^2},$$

where,

$$G(\theta) = g^2(\theta) \quad 0 \leq G(\theta) \leq 1$$

V_x = Variance of U_x (Noise)

V_y = Variance of U_y (Signal)

$$\beta = \frac{V_x}{V_y}$$

$K_x > 3$ Kurtosis of U_x (Noise)

$K_y = 3$ Kurtosis of U_y (Signal)

Simulate metrics:

$$\min(\sigma) = \sqrt{V_x + V_y G_{\min}},$$

$$\max(\sigma) = \sqrt{V_x + V_y},$$

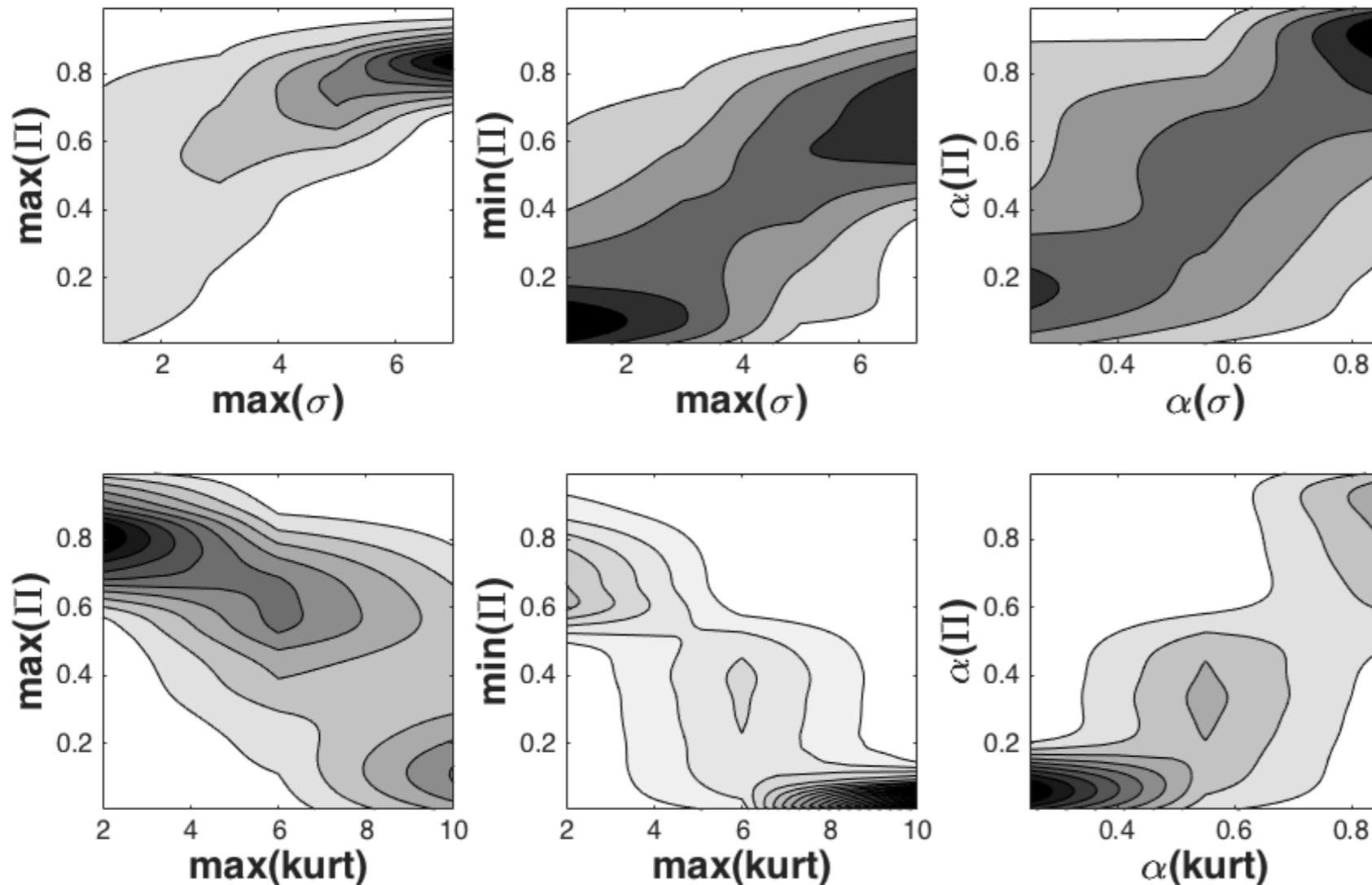
$$\min(\text{kurt}) = \frac{K_x \beta^2 + K_y + 6\beta}{\beta^2 + 2\beta + 1},$$

$$\max(\text{kurt}) = \frac{K_x \beta^2 + G_{\min}^2 K_y + 6G_{\min} \beta}{\beta^2 + 2G_{\min} \beta + G_{\min}^2},$$

$$\min(\Pi) = \frac{G_{\min}}{\beta + G_{\min}},$$

$$\max(\Pi) = \frac{1}{\beta + 1};$$

Idealized model



Signal-to-noise ratio:

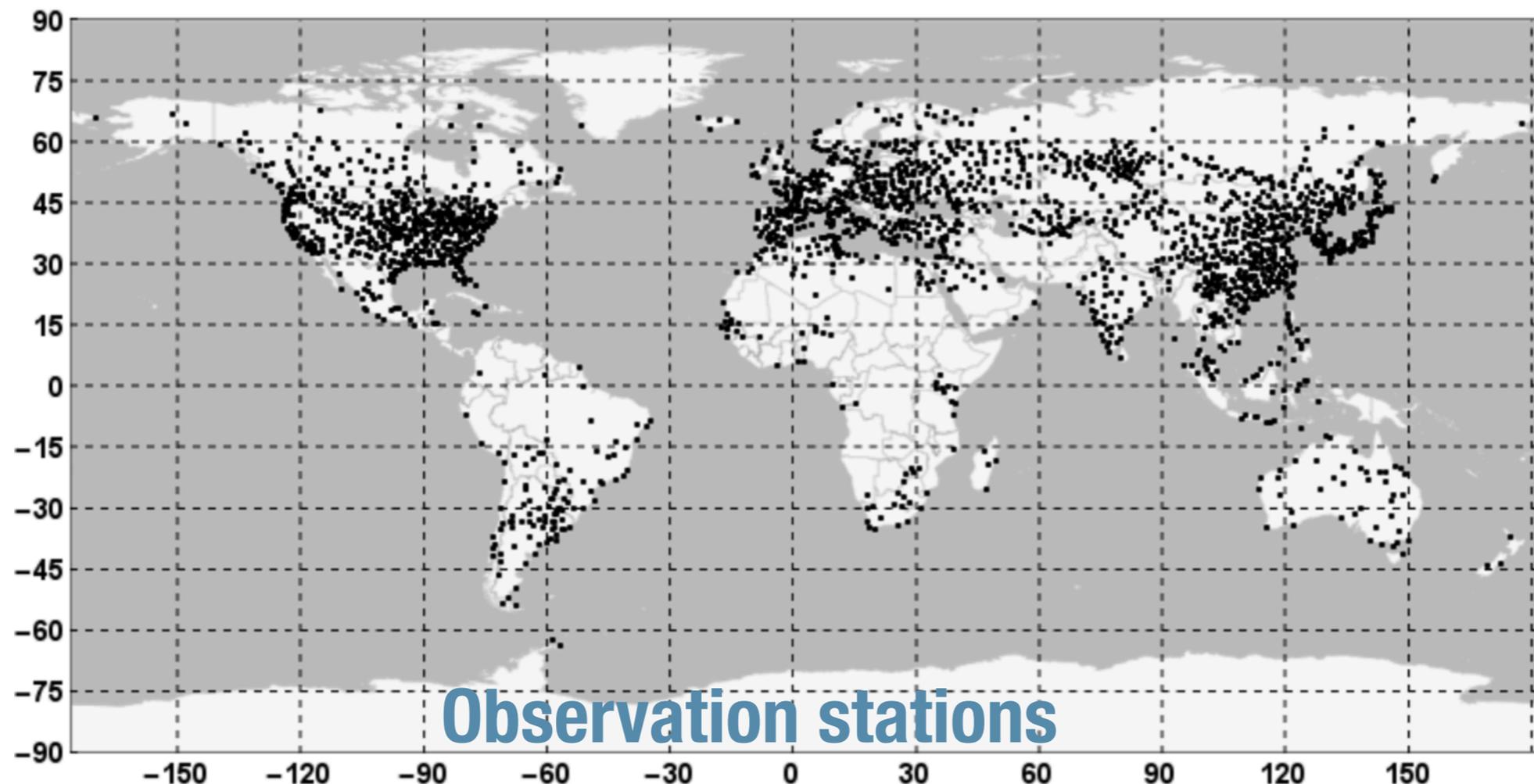
accounts for observed characteristics of linear predictability (e.g. predictive anisotropy)

Origin of the noise:

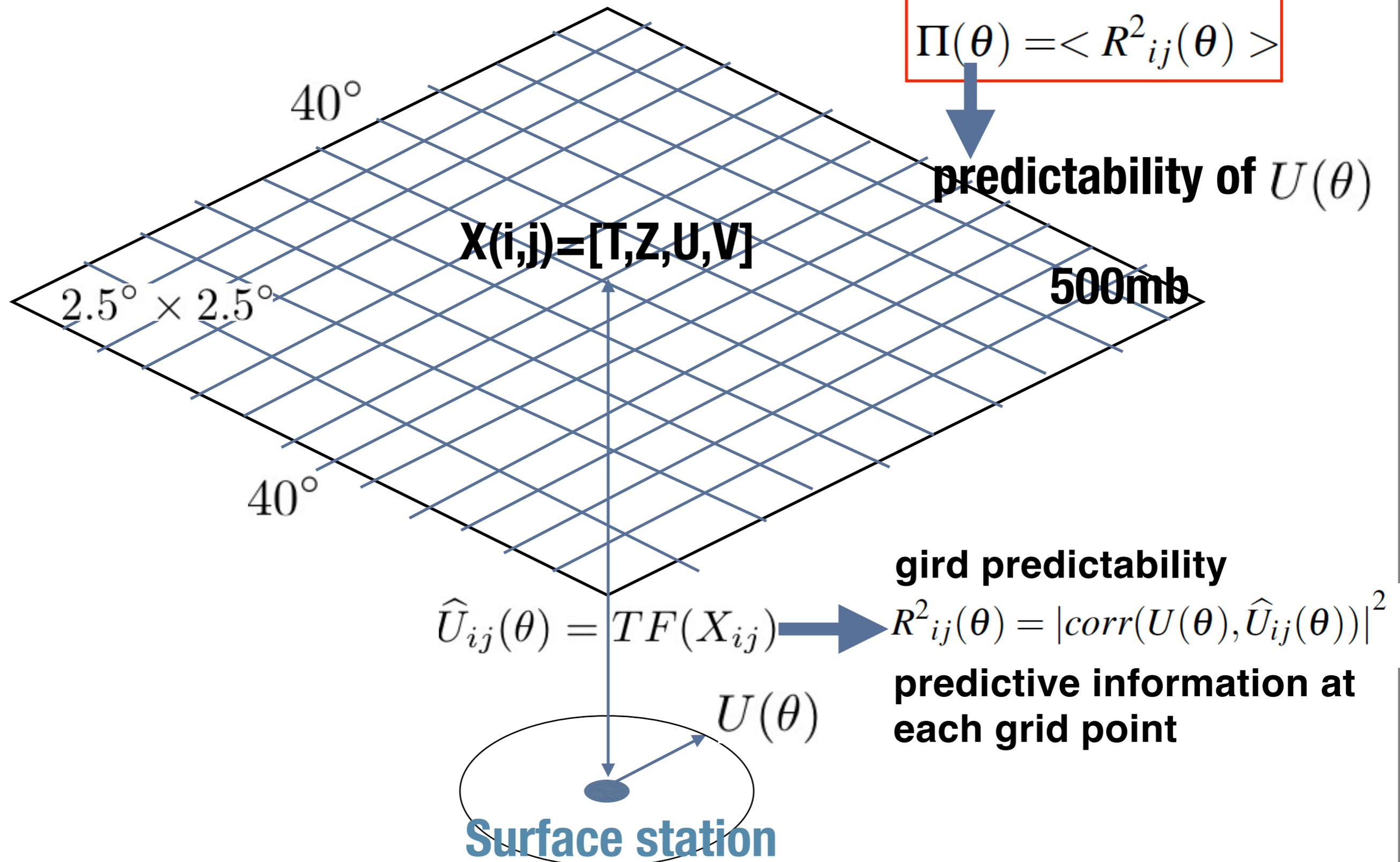
- physical factors?
- nonlinear predictor-predictand relationship?

Linear vs nonlinear TFs

- * Nonlinear prediction of surface wind components is carried out at the same 2109 stations
- * Neural network (NN), Support vector machine (SVM), Random Forest (RF)

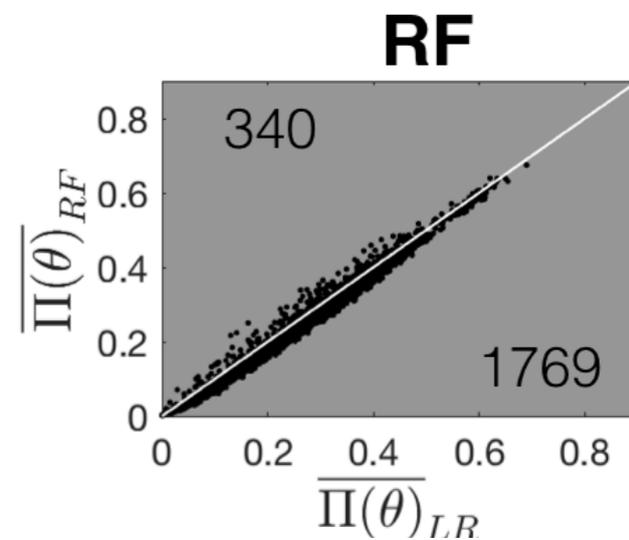
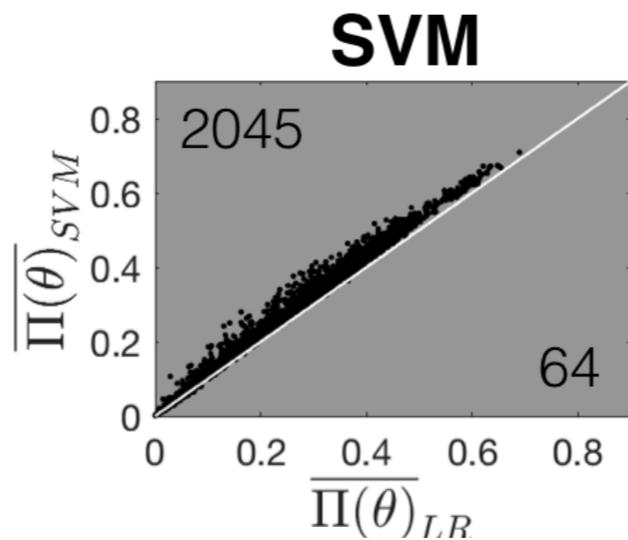
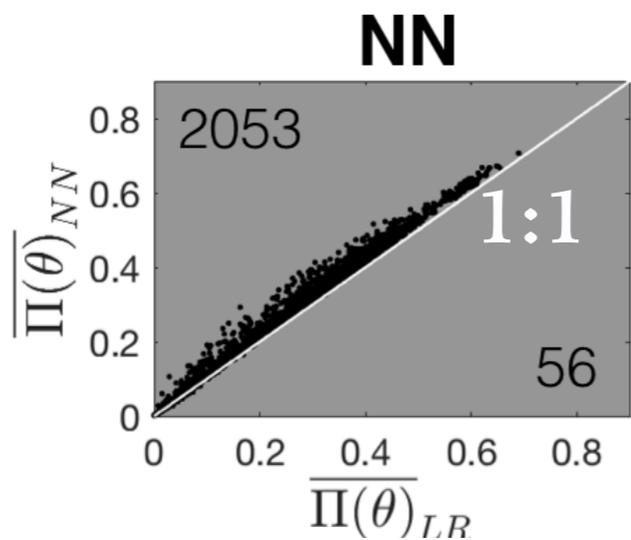


Prediction at each station

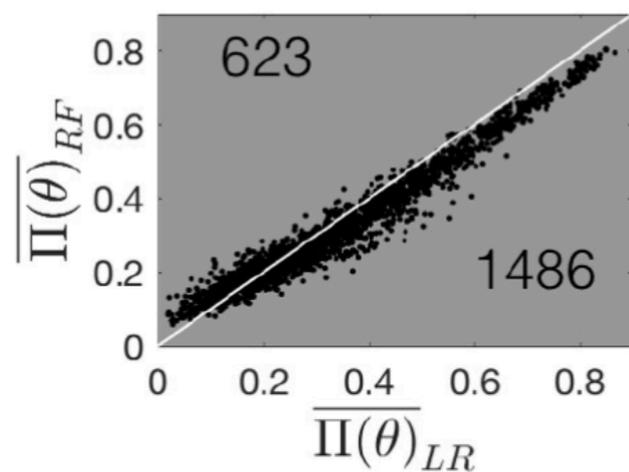
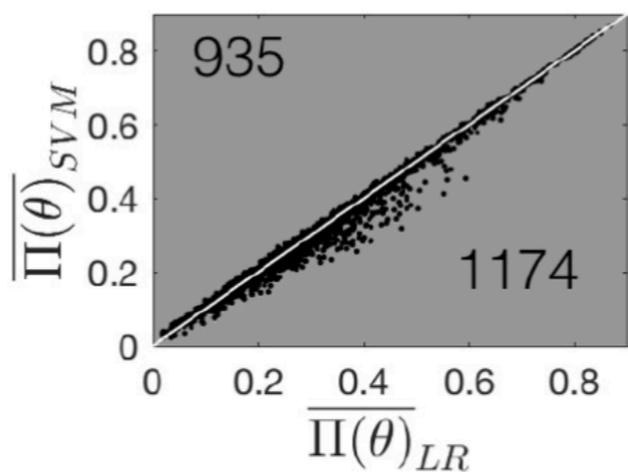
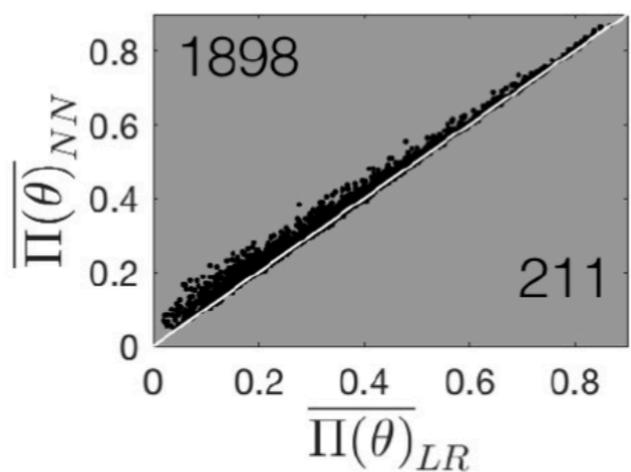


TF: nonlinear (NN, SVM, RF)

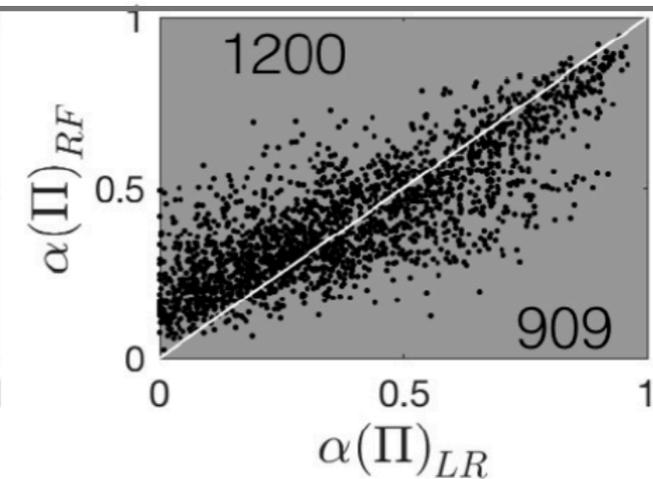
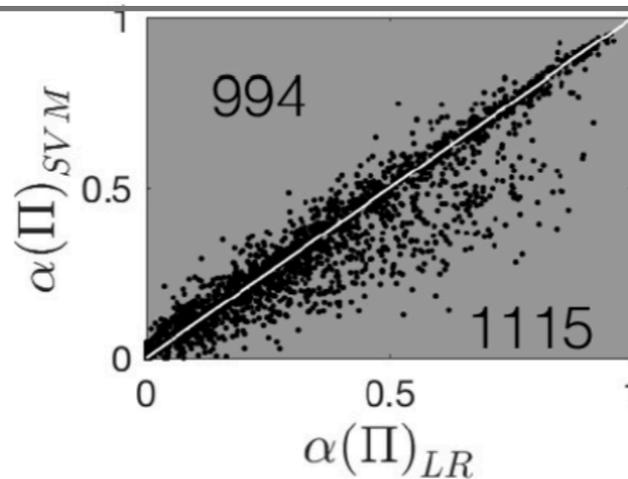
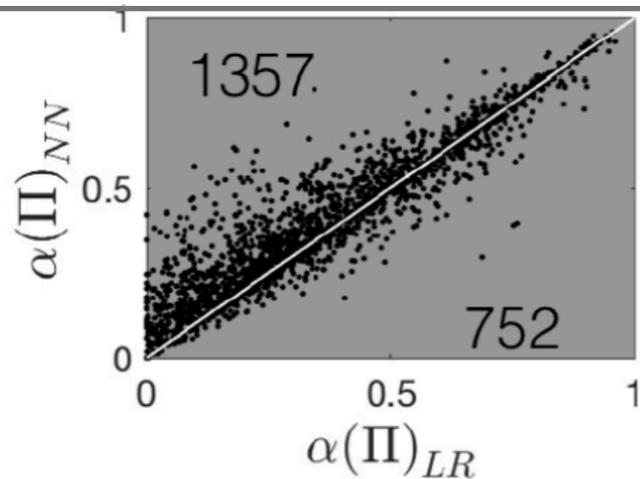
**Daily
Winter**



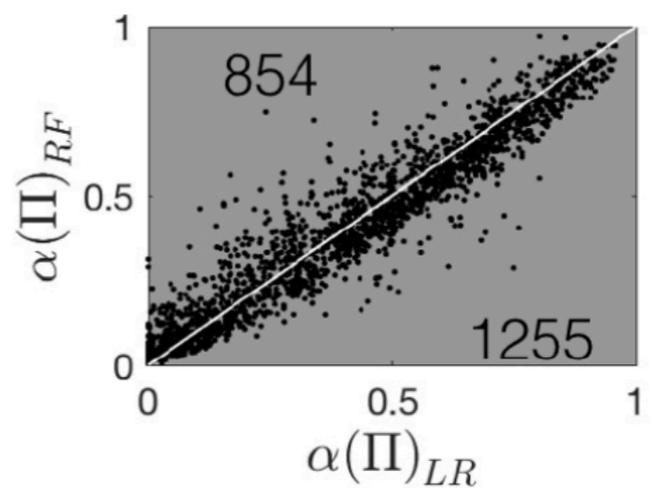
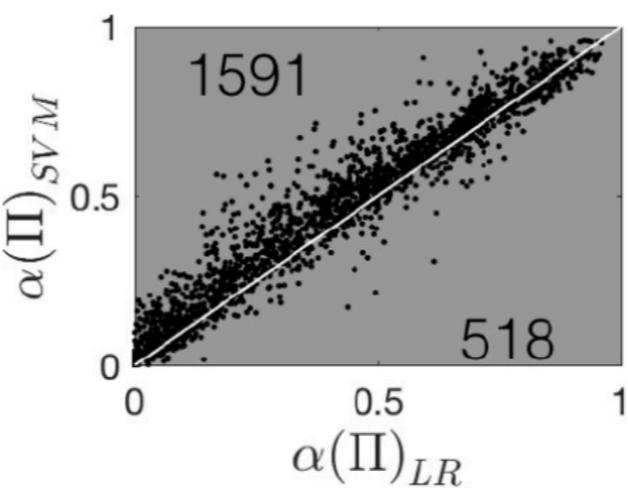
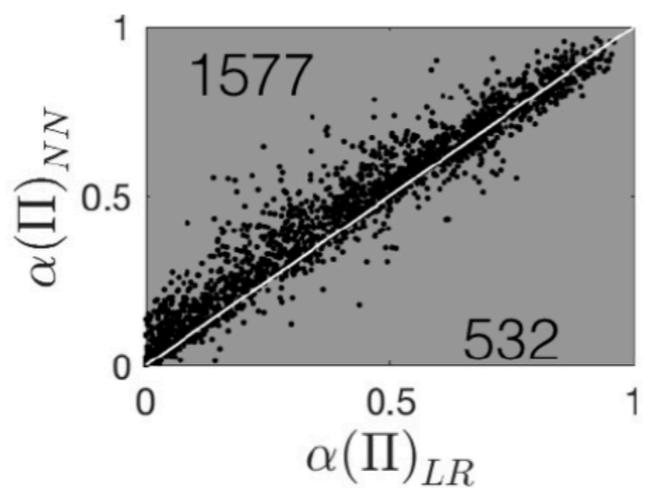
**Monthly
Winter**



**Monthly
Winter**

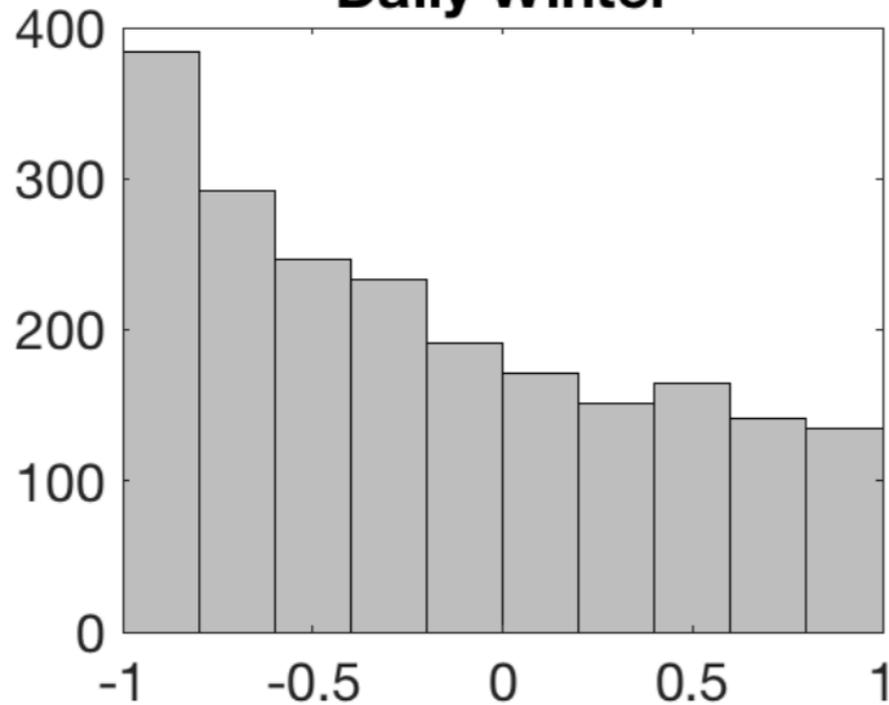


**Daily
Winter**



Daily Winter

$\rho(\sigma, \beta)$



$$\beta = \frac{\Pi(\theta)_{BestNL}}{\Pi(\theta)_{LR}}$$

Systematic improvement by NL
no substantial improvement by NL

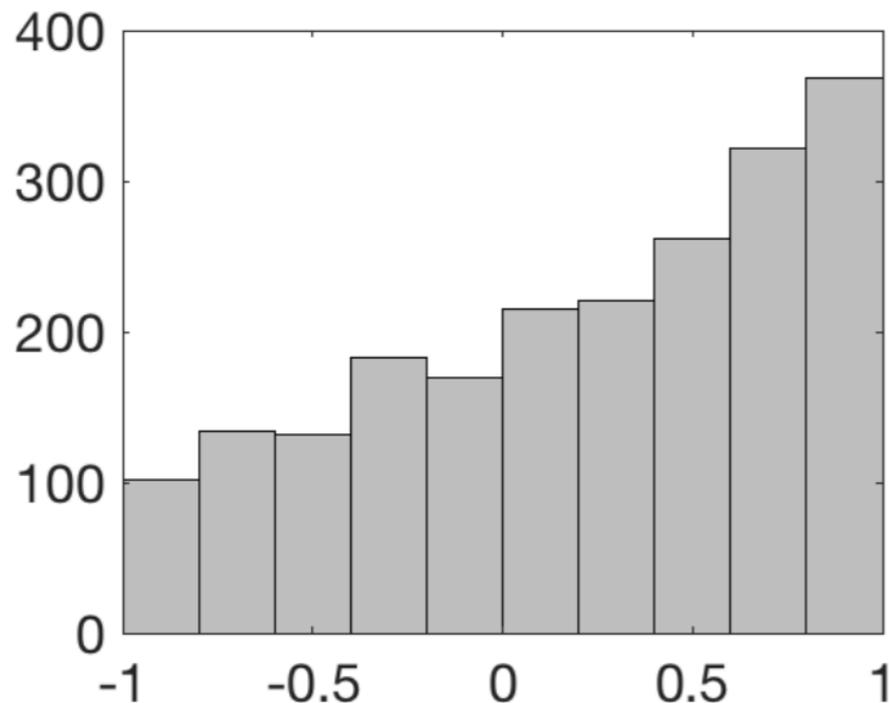
e.g.,

$$\Pi_{NL} = 0.1$$

$$\beta = 10$$

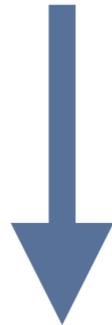
$$\Pi_{LR} = 0.01$$

$\rho(kurt, \beta)$



No compelling evidence to suggest that strong nonlinear relationships exist between large-scale predictors and surface wind components

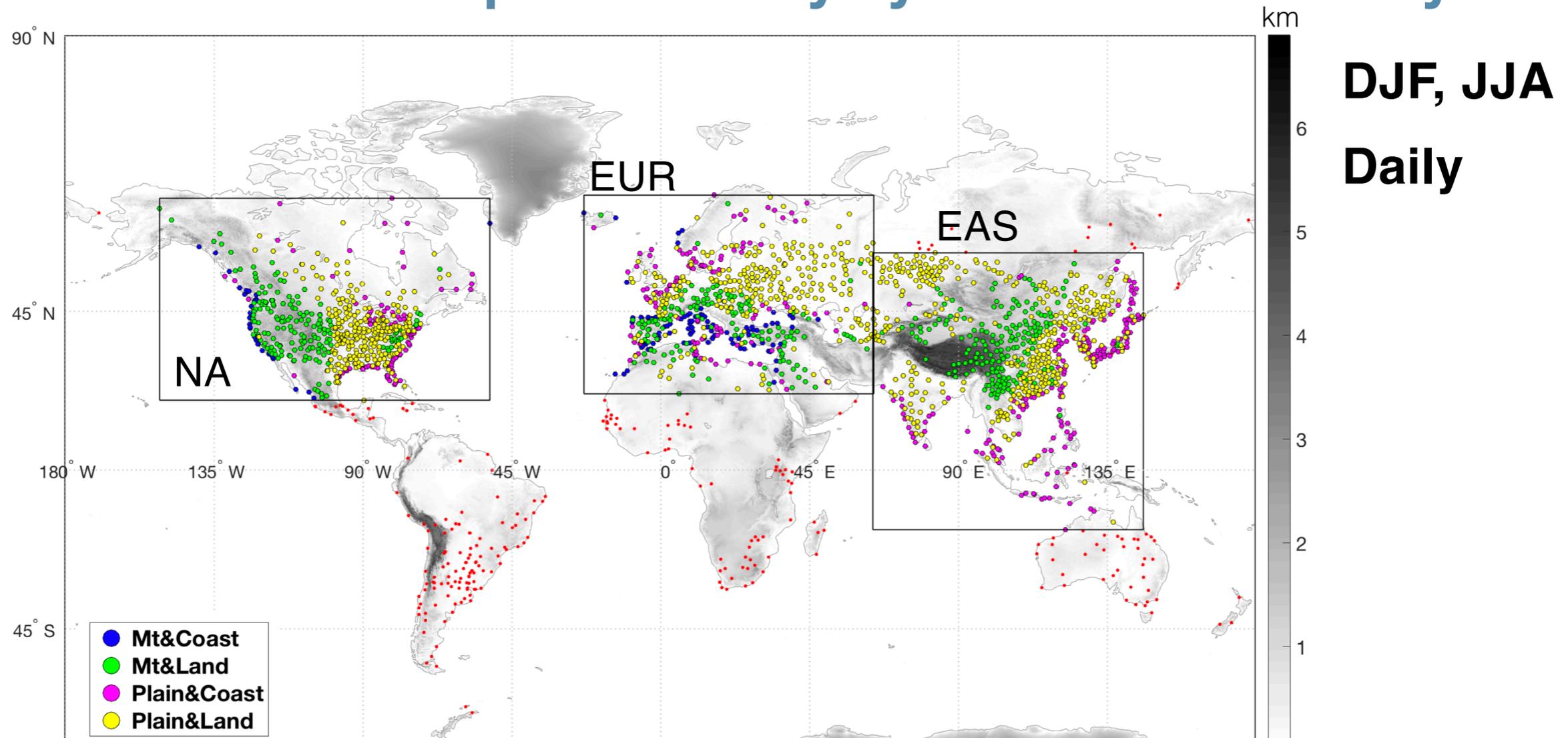
predictive anisotropy (contributing factors)



- wrong functional form of TF (i.e. linear TF)? **X**
- physical factors?

Can predictive anisotropy be explained by some unknown physical factors?

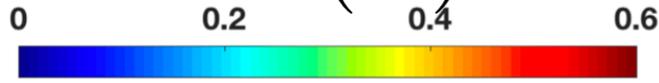
Simulation of linear predictability by RCMs and Reanalysis



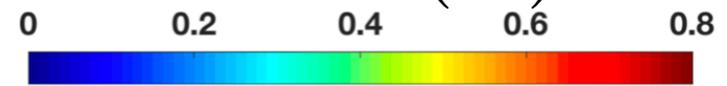
Labels	Driving reanalysis	Regional models	Modeling Group	Project	References
NA1	ECMWF-ERAINT	Canadian Regional Climate Model 4 (CanRCM4)	Canadian Centre for Climate Modelling and Analysis (CCCma)	CORDEX	Scinocca et al., 2016
NA2	NCEP2	Weather Research & Forecasting Model (WRF)	Pacific Northwest National Lab, US	NARCCAP	Mearns et al., 2007, updated 2014
EUR1	ECMWF-ERAINT	Rosby Centre regional atmospheric model (RCA4)	Swedish Meteorological and Hydrological Institute (SMHI)	CORDEX	Strandberg et al., 2015
EUR2	ECMWF-ERAINT	Canadian Regional Climate Model 4 (CanRCM4)	Canadian Centre for Climate Modelling and Analysis (CCCma)	CORDEX	Scinocca et al., 2016
EAS1	ECMWF-ERAINT	HadGEM3-RA	National Institute of Meteorological Research (NIMR), Korea	CORDEX	Davies et al., 2005
EAS2	NCEP2	Weather Research & Forecasting Model (WRF)	Seoul National University (SNU), Korea	CORDEX	Skamarock et al., 2005

Obs

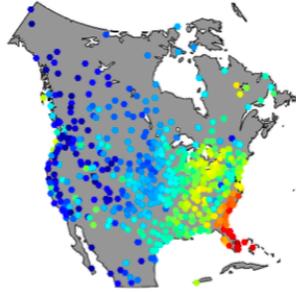
$min(\Pi)$



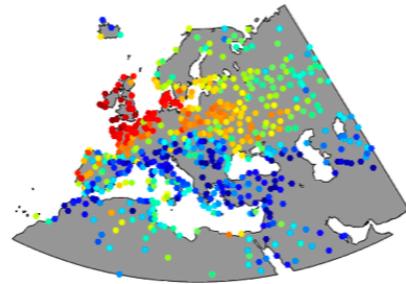
$max(\Pi)$



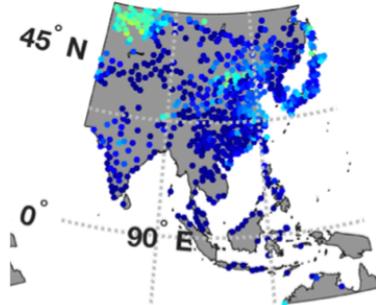
DJF min(Π)



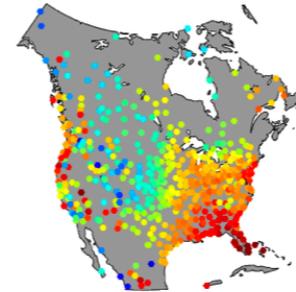
DJF min(Π)



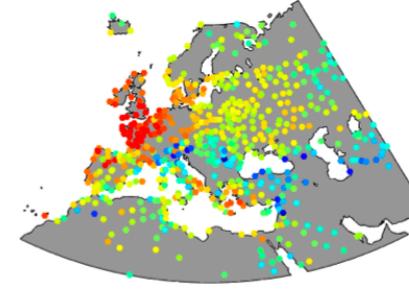
DJF min(Π)



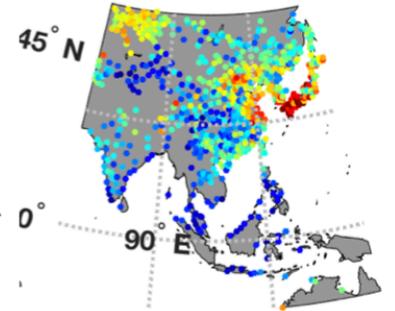
DJF max(Π)



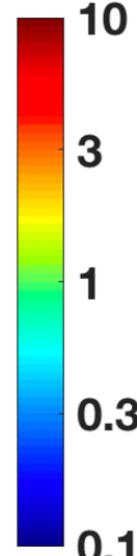
DJF max(Π)



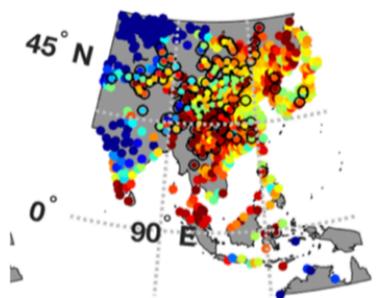
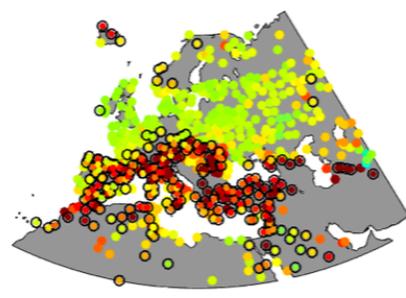
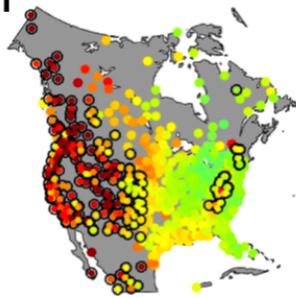
DJF max(Π)



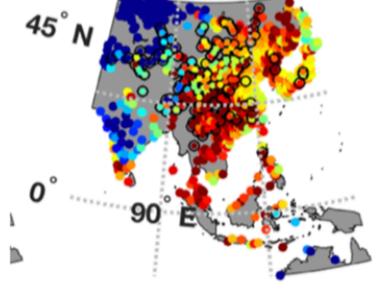
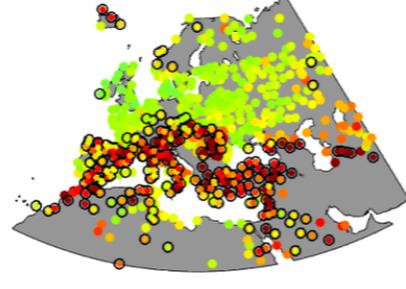
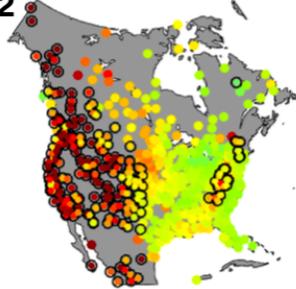
M/O



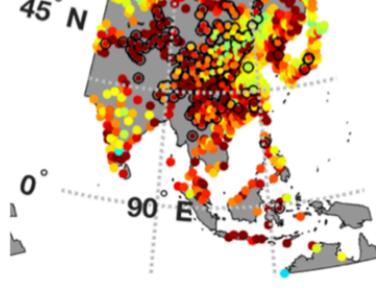
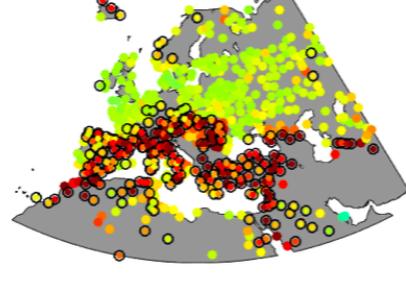
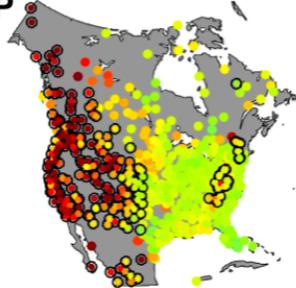
RCM1



RCM2



NCEP



North America

Europe

East Asia

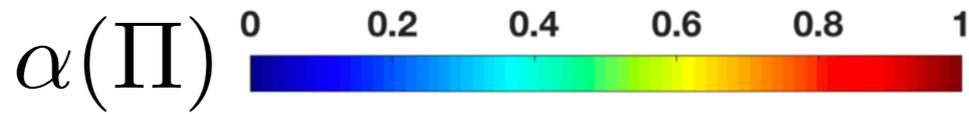
North America

Europe

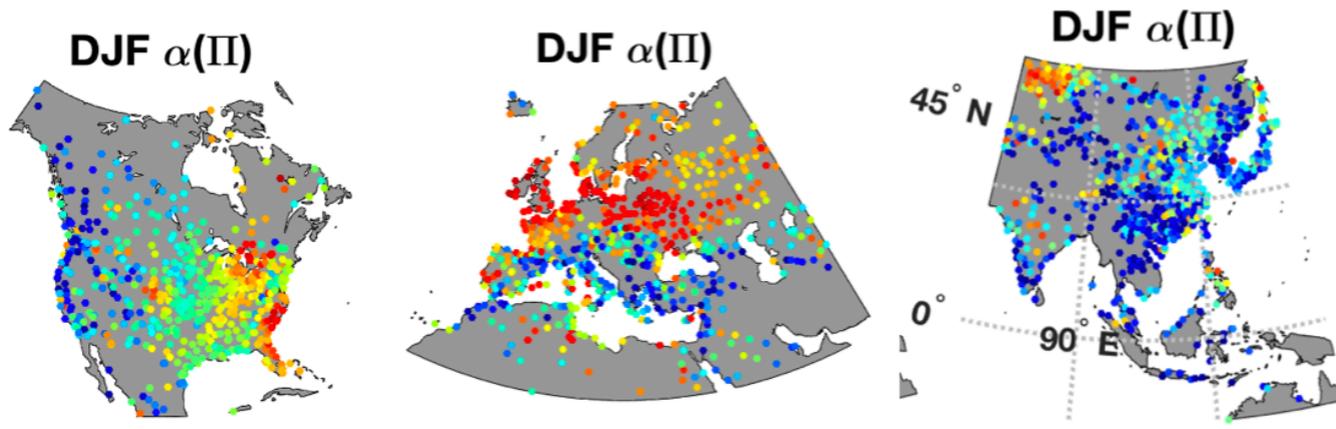
East Asia

○ Mt stations

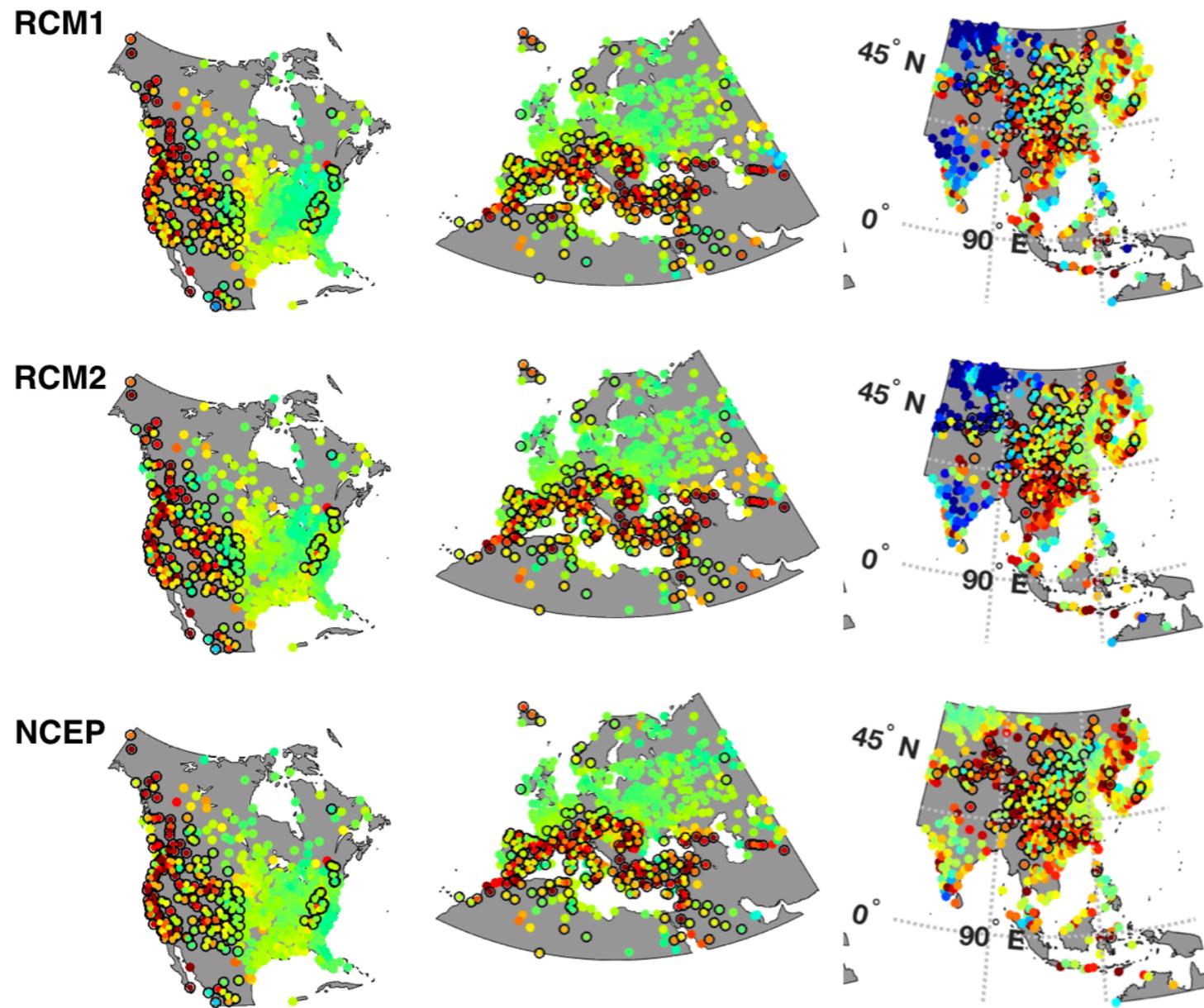
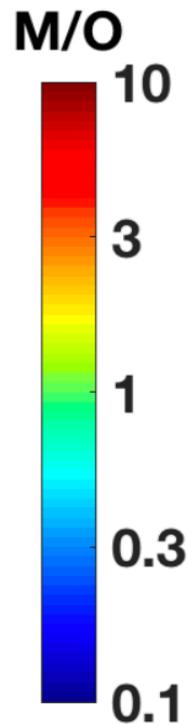
RCMs cannot capture small-scale physical processes (e.g. associated with surface heterogeneity)



Obs



Simulated predictive anisotropy is substantially weakened in mountainous region.



Small-scale physical processes (not captured by RCMs) contribute to predictive anisotropy

North America Europe East Asia

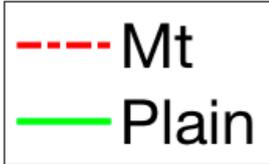
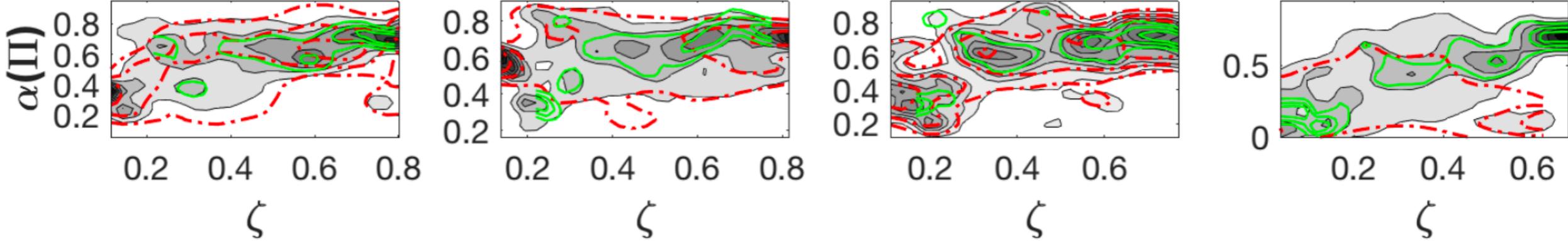
North America, (Europe, East Asia)

RCM1

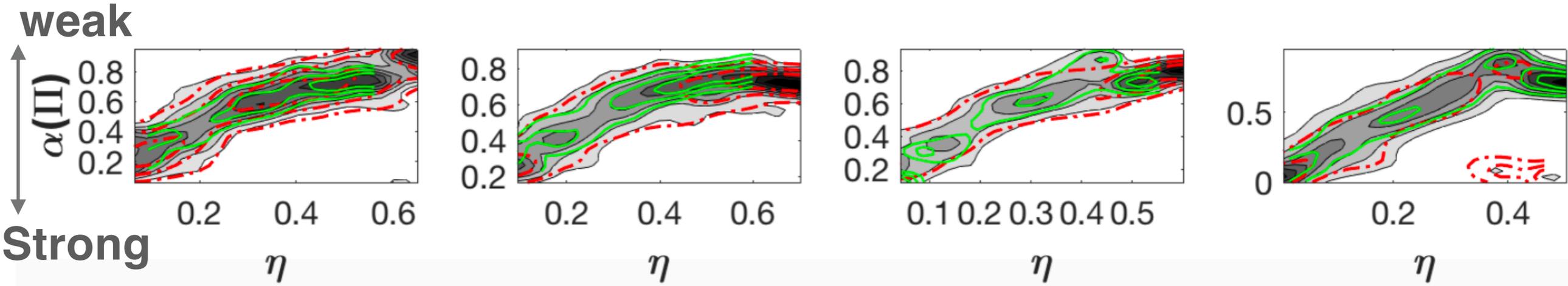
RCM2

Reanalysis

Obs



SNR $|\hat{e}_{max}(\Pi)$ large-scale physical processes



SNR $|\hat{e}_{min}(\Pi)$ small-scale physical processes

Predictive anisotropy can be explained by small-scale physical processes.

Conclusions

- * Predictive anisotropy is a **common** feature
- * Surface wind components are **better predicted** along the directions characterized by **more variable** and near **Gaussian** fluctuations.
- * **Poor predictability** is often found in **topographic complex** regions (e.g. mountainous regions), along the directions of **weak and non Gaussian** fluctuations of surface wind components.
- * **No** concrete evidence to show that the relationships between free-tropospheric predictors and surface wind components are **nonlinear**.
- * **Small scale physical processes** (not captured by RCMs) contribute to **predictive anisotropy**.
- * **Future study is needed to identify physical processes responsible for predictive anisotropy.**