

# Kinetic Theory for Geophysical Flows

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In collaboration with:

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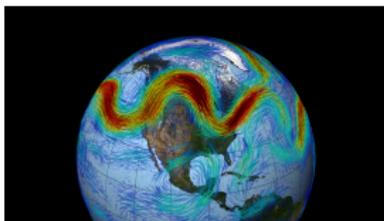
*G. Falkovich* Weizmann Institute, Israel

*A. Frishman* Princeton University, USA



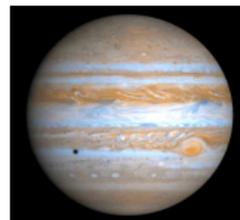
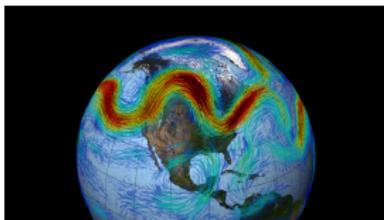
# Large-scale structures in geophysical flows

Geophysical flows self-organize into large-scale coherent structures, which play a major part in weather/climate.

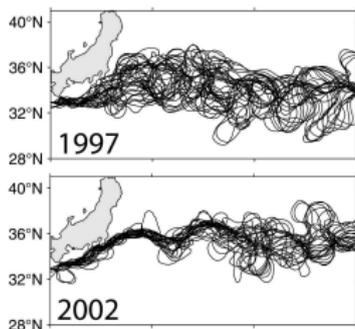


# Large-scale structures in geophysical flows

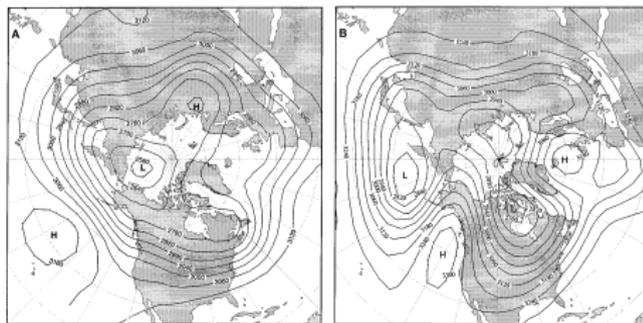
Geophysical flows self-organize into large-scale coherent structures, which play a major part in weather/climate.



They fluctuate and undergo abrupt transitions.



*Kuroshio path*<sup>1</sup>



*Zonal/blocked Jet Stream transition*<sup>2</sup>

<sup>1</sup>B Qiu and S. M. Chen (2005). *J. Phys. Oceanogr.*

<sup>2</sup>E. R. Weeks et al. (1997). *Science*

# Turbulent flows and degrees of freedom

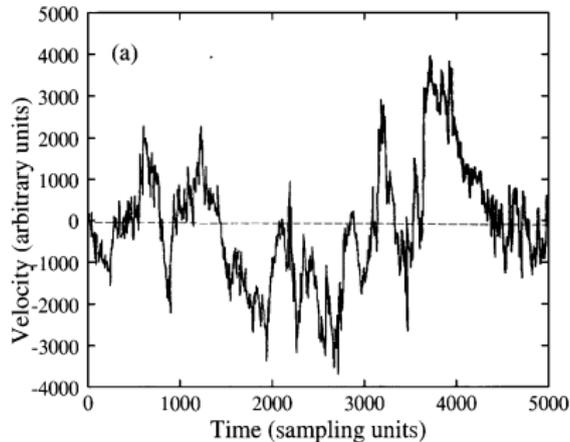
**Navier-Stokes equations:** nonlinear term couples wide range of scales.

$$\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla P + \nu \Delta \mathbf{u}, \quad \text{Re} = UL/\nu, \quad \#DOF \sim \text{Re}^{9/4} \approx 10^{20}$$

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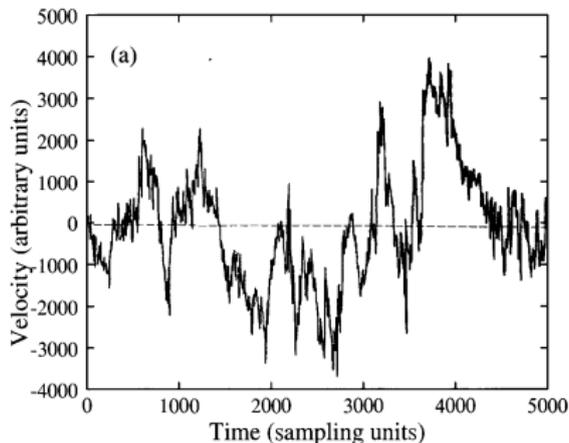


ONERA wind tunnel (Gagne, Hopfinger)

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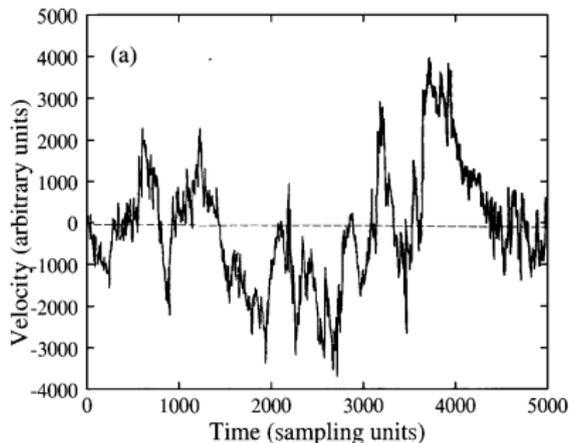
ONERA wind tunnel (Gagne, Hopfinger)

**Statistical Physics Approach**

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ONERA wind tunnel (Gagne, Hopfinger)

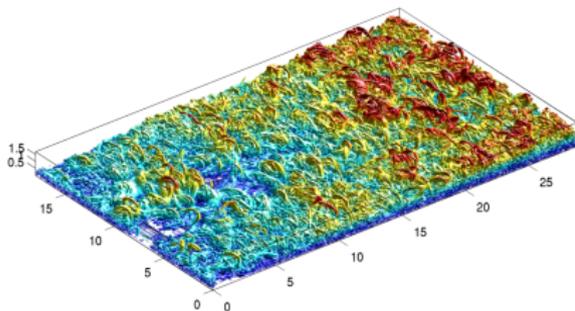
**Statistical Physics Approach**

*Difficult problem: closure !*

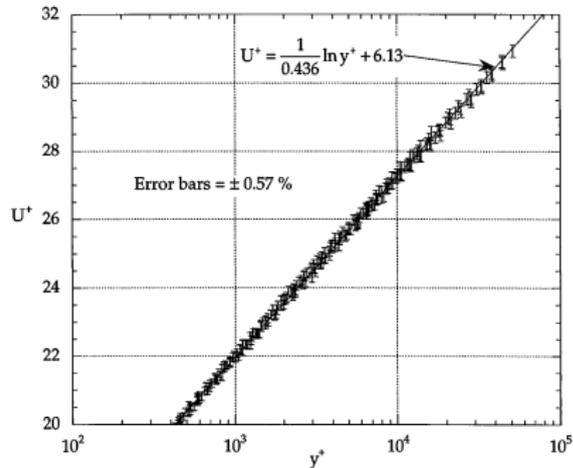
# Mean-flow/turbulence interactions

How to compute velocity profile in turbulent shear flows?

*There is no consistent theory of mean-flow/turbulence interactions.*



E.g. Turbulent Boundary Layer



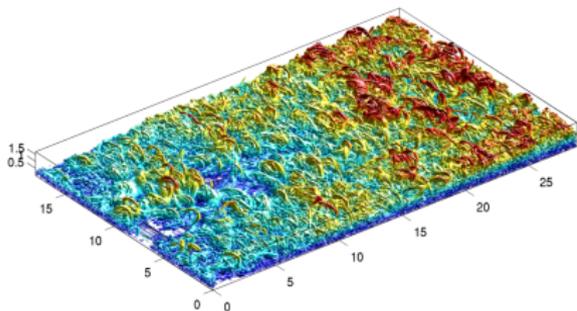
Experimental data for pipe flow<sup>3</sup>

<sup>3</sup>M. V. Zagarola and A. J. Smits (1997). *Phys. Rev. Lett.*

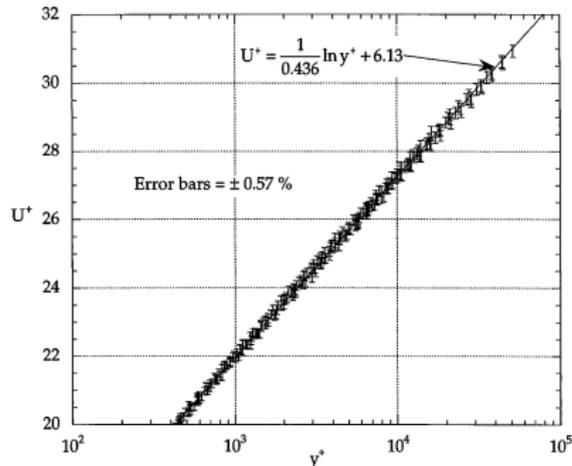
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Experimental data for pipe flow<sup>3</sup>

Conservation equations:  $\varepsilon \approx -\tau U'$  with  $\tau \equiv \langle uv \rangle \approx \text{const.}$

- ▶ 3D: The mean-flow feeds turbulence,  $\tau U' < 0$ . *Not closed!  $\varepsilon(y)$ ?*
- ▶ 2D: Turbulence feeds the mean-flow,  $\tau U' > 0$ .

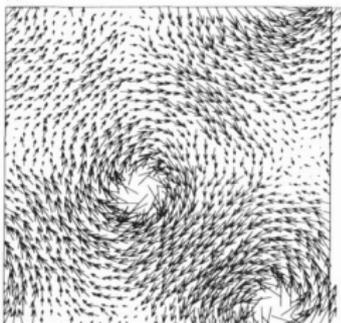
<sup>3</sup>M. V. Zagarola and A. J. Smits (1997). *Phys. Rev. Lett.*

# Finite-size effects in 2D turbulence: condensation

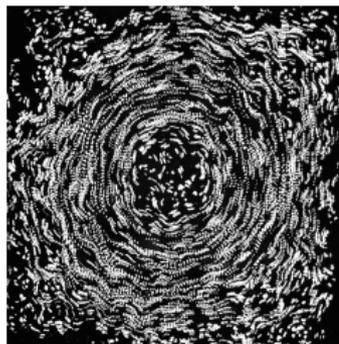
Simplest possible system: 2D Navier-Stokes with linear friction

$$\partial_t \omega + \mathbf{u} \cdot \nabla \omega = \nu \Delta \omega - \alpha \omega + f_\omega.$$

When  $\alpha$  is small enough, the inverse cascade reaches the box scale: "condensation".



DNS<sup>4</sup>



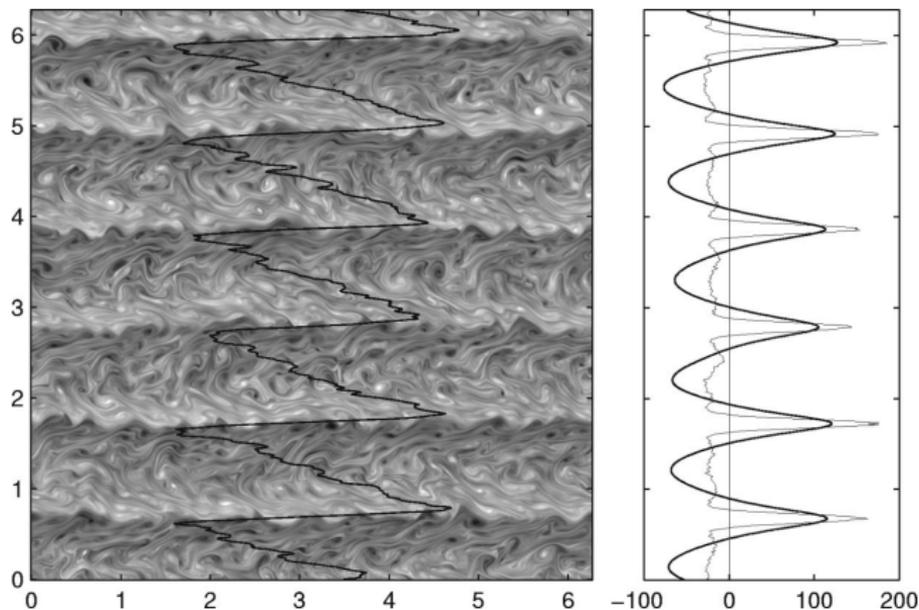
ANU 2D experiments  
(EM, Faraday waves)<sup>5</sup>

<sup>4</sup>L. M. Smith and V. Yakhot (1994). *J. Fluid Mech.*

<sup>5</sup>H Xia et al. (2009). *Phys. Fluids*; see also J. Sommeria (1986). *J. Fluid Mech.*

# Zonal Jets in Rotating Flows

Beta-plane turbulence:  $\partial_t \omega + \mathbf{u} \cdot \nabla \omega + \beta v = \nu \Delta \omega - \alpha \omega + f_\omega$ .



Potential vorticity and zonal velocity profile<sup>6</sup>.

<sup>6</sup>S. Danilov and D. Gurarie (2004). *Phys. Fluids*; see also P. B. Rhines (1975). *J. Fluid Mech.* M. E. Maltrud and G. K. Vallis (1991). *J. Fluid Mech.* B Galperin and S Sukoriansky (2001). *Phys. Fluids*,...

## Main Questions and Theoretical Tools

### Generic questions:

- ▶ Can we predict self-organization of geophysical flows into large scale coherent structures?
- ▶ Characterize the attractors of geophysical turbulence
- ▶ Study fluctuations around the mean state
- ▶ What aspects of abrupt transitions in turbulent flows are predictable?

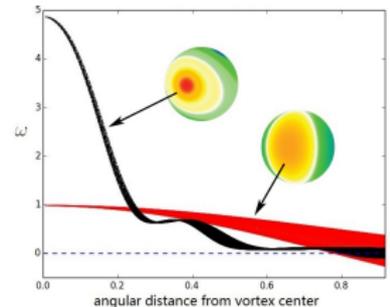
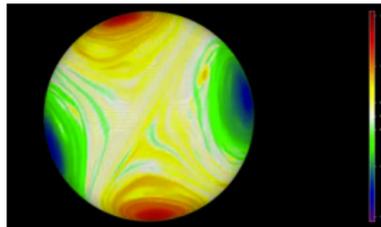
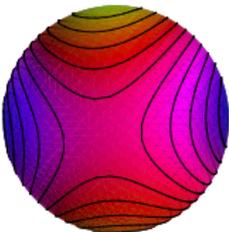
### Tools from Statistical Physics

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**Equilibrium States**<sup>7</sup> vs. numerical simulations<sup>8</sup>:



The set of MRS equilibria is huge: difficult to make quantitative predictions.

*Can we find a theory which provides quantitative predictions based on forcing/dissipation?*

<sup>7</sup>C. Herbert (2013). *J. Stat. Phys.*

<sup>8</sup>W. Qi and J. B. Marston (2014). *J. Stat. Mech.*

## Main Questions and Theoretical Tools

### Generic questions:

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### Tools from Non-Equilibrium Statistical Physics

- ▶ Kinetic Theory
- ▶ Large Deviation Theory

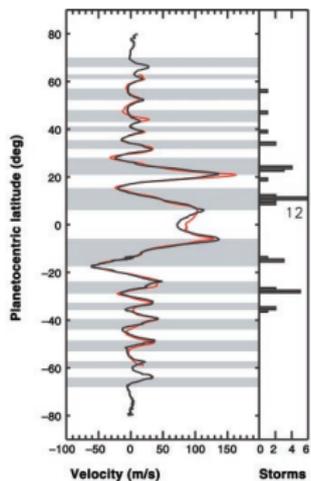
# Outline

- 1 Introduction
- 2 Kinetic Theory: timescale separation**
- 3 Testing Kinetic Theory: mean-flow and fluctuations
- 4 Conclusion

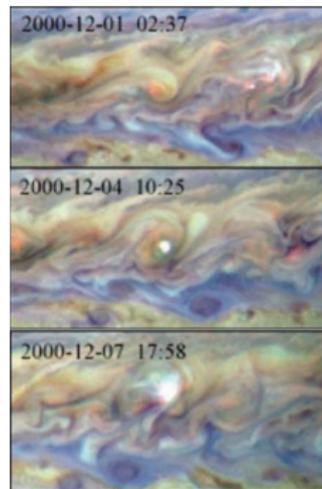
# Timescale separation in geophysical flows

In some flows, there is a *natural timescale separation*, usually associated to a broken symmetry of the Navier-Stokes equations.

E.g. Jupiter<sup>7</sup>:



Zonal wind measured by Voyager 2 (1979, red) and Cassini (2000, black).



Cassini

<sup>7</sup>C. C. Porco et al. (2003). *Science*.

Adiabatic elimination of fast variables<sup>8</sup> (*stochastic averaging*)

## Slow-fast SDE:

$$dX_t = f(X_t, Y_t)dt + \sqrt{2\epsilon}dW_t,$$

$$dY_t = \alpha^{-1}g(X_t, Y_t)dt + \sqrt{\alpha^{-1}}h(X_t, Y_t)dW_t.$$

- ▶ Joint PDF  $P(x, y; t)$ ; Fokker-Planck equation  $\partial_t P = (\alpha^{-1}L_0 + L_1)P$ .
- ▶ Stationary distribution for fast modes at fixed  $x$  and projection operator:

$$L_0 P_\infty^x(y) = 0, \quad \mathcal{P}\phi = P_\infty^x(y) \int dy \phi(x, y).$$

- ▶ Write  $P_s = \mathcal{P}P$ ,  $P_f = (1 - \mathcal{P})P$ . We have  
 $\partial_t P_s = \mathcal{P}(\alpha^{-1}L_0 + L_1)P = \mathcal{P}L_1 P$ .
- ▶ At lowest order,  $\partial_t P_s = \mathcal{P}L_1 P_s + O(\alpha)$  and  $P_s(x, y) = P_\infty^x(y)Q(x)$  with

$$\frac{\partial Q}{\partial t} = -\frac{\partial}{\partial x} [\mathbb{E}_\infty^x[f]Q(x)] + \epsilon \frac{\partial^2}{\partial x^2} Q + O(\alpha).$$

Finally, after adiabatic reduction:

$$dX_t = \mathbb{E}_\infty^{X_t}[f]dt + \sqrt{2\epsilon}dW_t.$$

<sup>8</sup>e.g. C. W. Gardiner (2009). *Handbook of Stochastic Methods for physics, chemistry, and the natural sciences*. 4th edition. Springer, Berlin.

# Adiabatic elimination of fast variables: zonal jets

## Reynolds decomposition for the zonal jets

$\omega = \bar{\omega} + \omega'$ , with  $\bar{\cdot}$  the projection on the (slow) zonal modes.  
Formally,

$$\begin{aligned}\partial_t \bar{\omega} + L_{\bar{\omega}}[\bar{\omega}] &= -\partial^i \overline{u'_i \omega'} + \bar{\eta}, \\ \partial_t \omega' + L'_{\bar{\omega}}[\omega'] &= -u'_i \partial^i \omega' + \partial^i \overline{u'_i \omega'} + \eta'.\end{aligned}$$

### Adiabatic reduction at lowest order<sup>9</sup>:

$$\begin{aligned}\partial_t \bar{\omega} + L_{\bar{\omega}}[\bar{\omega}] &= -\partial^i \mathbb{E}_{\bar{\omega}}[\overline{u'_i \omega'}] + \bar{\eta}, \\ \partial_t \omega' + L'_{\bar{\omega}}[\omega'] &= \eta'.\end{aligned}$$

- ▶ **No UV divergences**
- ▶ **Eddy-eddy interactions** do not contribute at leading order.

The fluctating vorticity field is an Ornstein-Uhlenbeck process characterized by the two-point correlation function  $g(\mathbf{r}_1, \mathbf{r}_2, t) = \mathbb{E}_{\bar{\omega}}[\omega'(\mathbf{r}_1, t)\omega'(\mathbf{r}_2, t)]$ , which satisfies the Lyapunov equation:

$$\partial_t g + L'^{(1)}_{\bar{\omega}} g + L'^{(2)}_{\bar{\omega}} g = C',$$

with  $C'(\mathbf{r}_1, \mathbf{r}_2, t) = \mathbb{E}[\eta'(\mathbf{r}_1, t)\eta'(\mathbf{r}_2, t)]$  the correlation matrix of the Gaussian white noise  $\eta'$ .

<sup>9</sup>F. Bouchet et al. (2013). *J. Stat. Phys.*

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## Numerical simulations in the quasi-linear framework<sup>9</sup>:

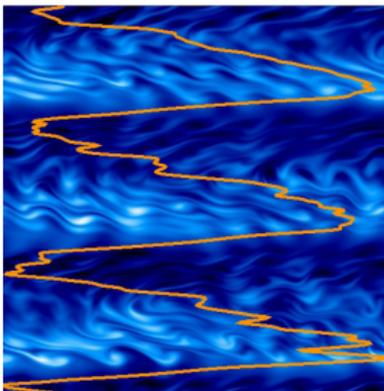
- ▶ Stochastic Structural Stability Theory<sup>10</sup>
- ▶ Cumulant Expansion “CE2”<sup>11</sup>

<sup>9</sup>T. Schneider and C. C. Walker (2006). *J. Atmos. Sci.* P. A. O’Gorman and T. Schneider (2007). *Geophys. Res. Lett.* K. Srinivasan and W. R. Young (2012). *J. Atmos. Sci.*

<sup>10</sup>B. F. Farrell and P. J. Ioannou (2003). *J. Atmos. Sci.* B. F. Farrell and P. J. Ioannou (2007). *J. Atmos. Sci.*

<sup>11</sup>S. M. Tobias and J. B. Marston (2013). *Phys. Rev. Lett.* J. B. Marston et al. (2016). *Phys. Rev. Lett.*

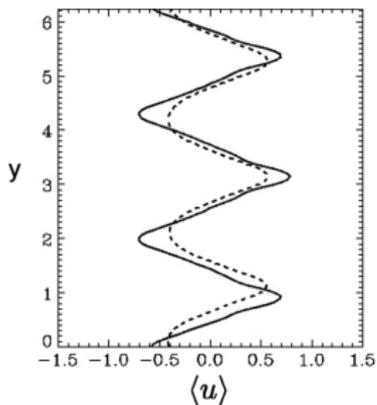
# Comparing QL numerical simulations with DNS



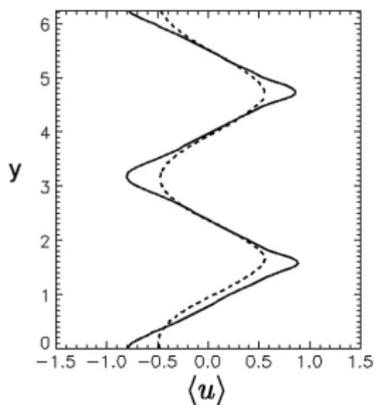
Cumulant Expansion CE2 for barotropic zonal jets<sup>12</sup>

$$R_\beta = \sqrt{\frac{U\beta^{1/5}}{2\varepsilon^{2/5}}}$$

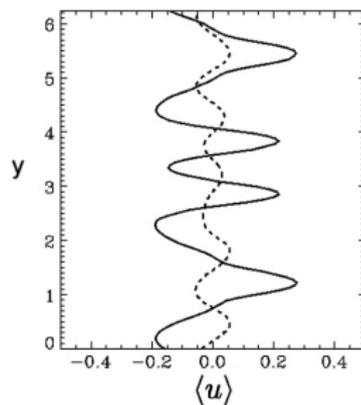
$R_\beta = 2.12$



$R_\beta = 1.98$



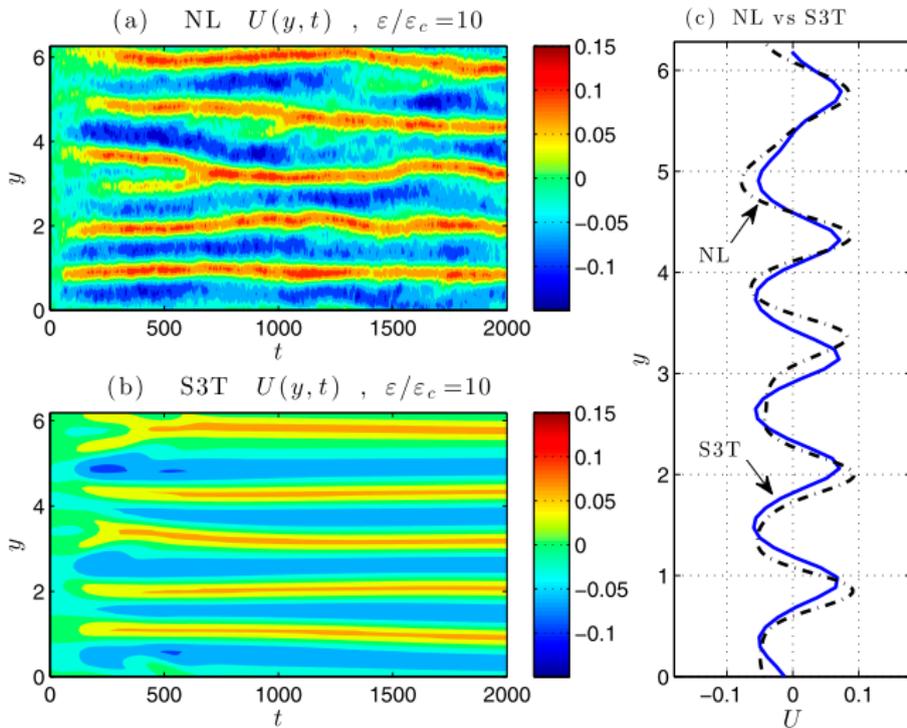
$R_\beta = 1.24$



<sup>12</sup>S. M. Tobias and J. B. Marston (2013). *Phys. Rev. Lett.*

# Comparing QL numerical simulations with DNS

## Stochastic Structural Stability Theory<sup>13</sup>



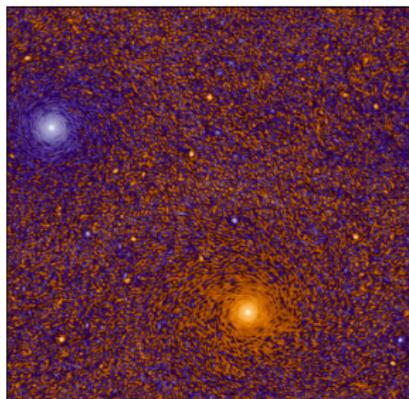
<sup>13</sup>N. C. Constantinou et al. (2014). *J. Atmos. Sci.*

# Outline

- 1 Introduction
- 2 Kinetic Theory: timescale separation
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## Explicit computations in the vortex condensate

Let us go back to 2D Navier-Stokes in a periodic square box with linear friction and small-scale random forcing:  $\partial_t \omega + \mathbf{u} \cdot \nabla \omega = \nu \Delta \omega - \alpha \omega + f_\omega$ .



DNS:  $1024^2$ ,  $k_f = 100$ , hyperviscosity,  $\alpha = 1.1 \times 10^{-4}$ .

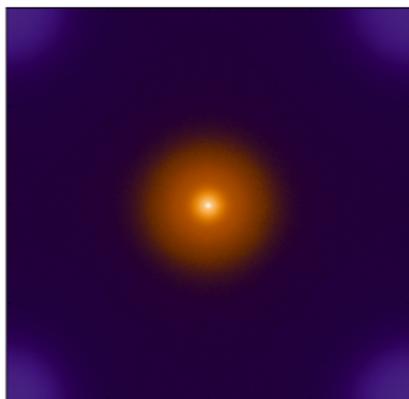
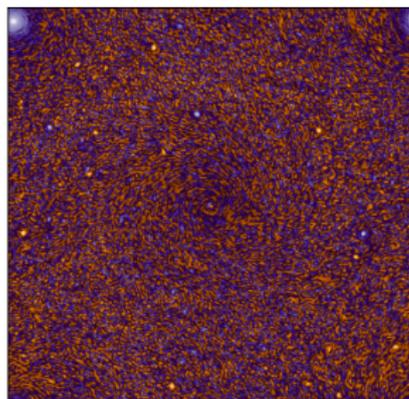
# Explicit computations in the vortex condensate

Reynolds decomposition:

$$\mathbf{v} = \mathbf{U} + \mathbf{u}, \text{ with } \mathbf{U} = U\mathbf{e}_\theta, \mathbf{u} = u\mathbf{e}_\theta + v\mathbf{e}_r \text{ and } \langle \mathbf{u} \rangle = 0,$$

$$\omega = \Omega + \omega', \text{ with } \langle \omega' \rangle = 0.$$

$$\partial_t \Omega + \mathbf{U} \cdot \nabla \Omega = -\alpha \Omega - \nabla \cdot \langle \mathbf{u} \omega' \rangle.$$

 $\Omega$  $\omega'$ 

DNS:  $1024^2$ ,  $k_f = 100$ , hyperviscosity,  $\alpha = 1.1 \times 10^{-4}$ .

## Explicit computations in the vortex condensate

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$$\omega = \Omega + \omega', \text{ with } \langle \omega' \rangle = 0.$$

$$\partial_t \Omega + \mathbf{U} \cdot \nabla \Omega = -\alpha \Omega - \nabla \cdot \langle \mathbf{u} \omega' \rangle.$$

### Timescale separation

*Perturbative expansion* of the equations of motion in  $\delta = \alpha L^{2/3} / \varepsilon^{1/3} \ll 1$  leads at first order to (Momentum and energy balance)<sup>14</sup>:

$$\begin{aligned} r^{-1} \partial_r (r^2 \langle uv \rangle) &= -\alpha r U, \\ r^{-1} \partial_r (r U \langle uv \rangle) + \alpha U^2 &= \varepsilon. \end{aligned}$$

Solution:

$$U = \sqrt{3\varepsilon/\alpha}, \quad \langle uv \rangle = -r \sqrt{\alpha\varepsilon/3}.$$

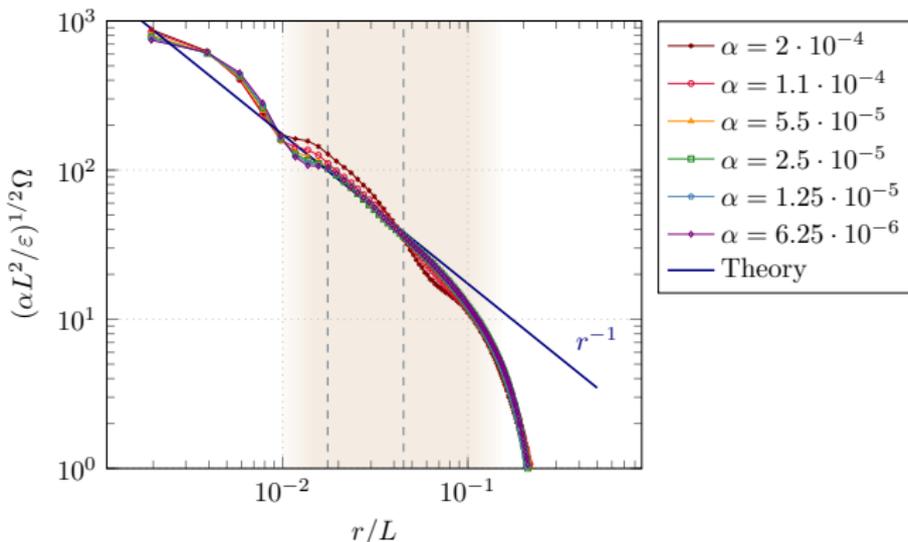
Therefore  $\Omega(r) = \sqrt{3\varepsilon/\alpha} r^{-1}$ .

Global energy balance neglecting small-scale dissipation yields  $U_{\text{rms}} = \sqrt{\varepsilon/\alpha}$ .

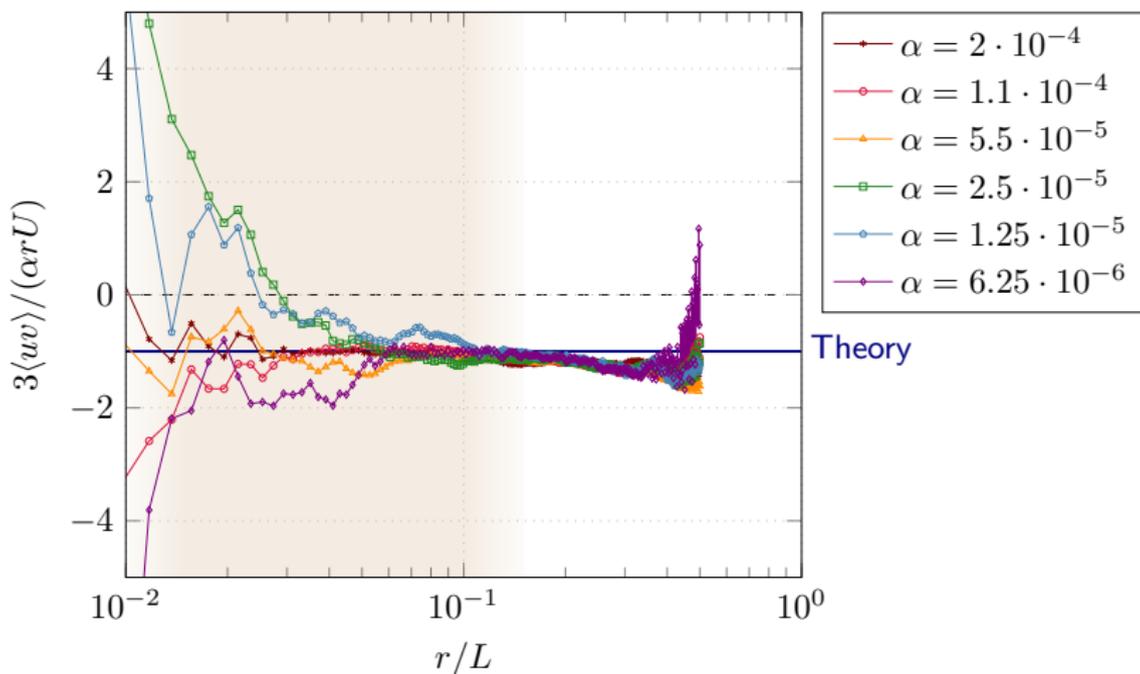
<sup>14</sup>J. Laurie et al. (2014). *Phys. Rev. Lett.*

# Explicit computation for the mean vorticity profile

Theoretical prediction:  $\Omega(r) = \sqrt{3\varepsilon/\alpha}r^{-1} = \sqrt{3}(\varepsilon L)^{1/3}\delta^{-1/2}r^{-1}$ .

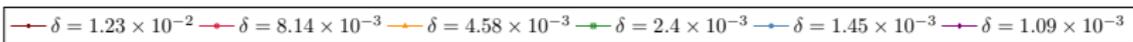
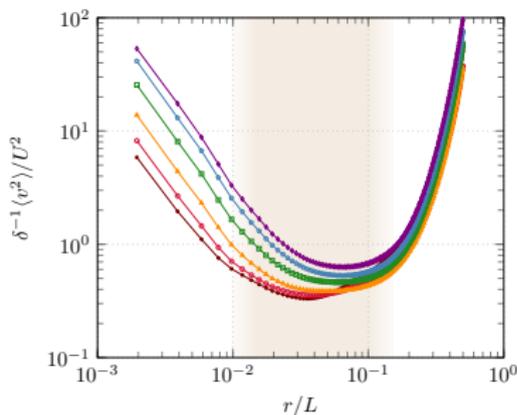
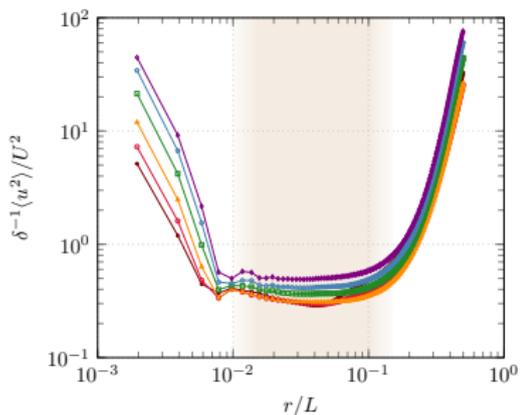


Our DNS ( $512^2$  and  $1024^2$ ) support the  $\alpha$ -scaling on a wide range of  $\alpha$ , and seem compatible with the  $r$ -scaling.

Explicit computation for the Reynolds tensor<sup>15</sup> $\langle uv \rangle / U^2 = O(\delta^{3/2})$  and not sign definite.DNS:  $512^2$ ,  $k_F = 100$ , hyperviscosity,  $\sim 300000$  turnover times.<sup>15</sup>A. Frishman and C. Herbert (submitted). *Phys. Rev. Lett.*

# Fluctuations: turbulent energy profile

DNS:  $512^2$ ,  $k_F = 100$ , hyperviscosity.



In the region of interest:

- ▶  $\langle u^2 + v^2 \rangle \ll U^2$  confirmed.
- ▶ Weak dependence on  $r$ .

*Turbulent energy profile not given by leading order energy/momentum balance.*

# Fluctuations: turbulent energy profile

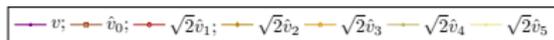
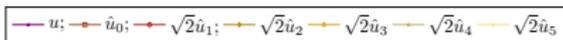
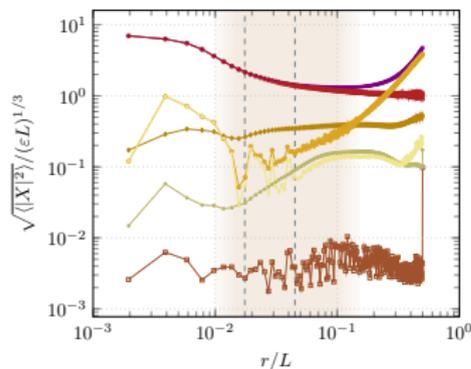
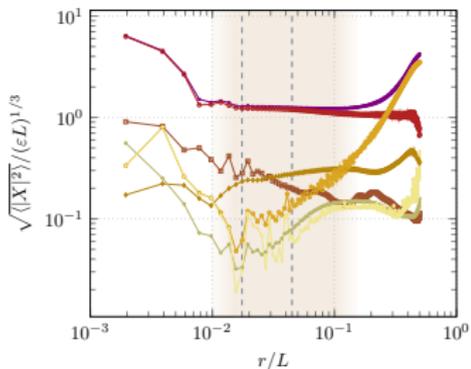
*Quasi-linear dynamics:*

$$\partial_t \omega + L_U[\omega] = \eta, \quad L_U = \frac{U}{r} \partial_\theta - \frac{\Omega'}{r} \partial_\theta \Delta^{-1} + \alpha \text{Id} - \nu \Delta.$$

Two-point correlation function:  $g(\mathbf{r}_1, \mathbf{r}_2) = \langle \omega(\mathbf{r}_1) \omega(\mathbf{r}_2) \rangle$ .

Lyapunov equation:  $\partial_t g + [L_U^{(1)} + L_U^{(2)}]g = C, \quad C = \langle \eta(\mathbf{r}_1) \eta(\mathbf{r}_2) \rangle$ .

Decompose into azimuthal harmonics  $u(r, \theta) = \sum_m \hat{u}_m(r) e^{im\theta}$



*Harmonics  $m = 1$  dominates in the region of interest.*

Here  $\alpha = 6.25 \times 10^{-6}$ , but this holds for all the runs.

## Fluctuations: turbulent energy profile

Behavior of harmonics explained by *zero modes of advection by the mean-flow*<sup>16</sup>

$$\langle |\hat{u}_1|^2 \rangle = A_1 k_f^{-4/9} \delta^{-1/3} (\varepsilon L)^{2/3} + \dots = A_1 (\varepsilon R_u)^{2/3} + \dots,$$

$$\langle |\hat{v}_1|^2 \rangle = (\varepsilon L)^{2/3} [A_1 k_f^{-4/9} \delta^{-1/3} + A_2 k_f^{-4/3} \delta^{-1} (\ell_f/r)^2] + \dots,$$

$$\text{Im} \langle \hat{u}_1 \hat{v}_1^* \rangle = A_1 k_f^{-4/9} \delta^{-1/3} (\varepsilon L)^{2/3} + \dots,$$

with  $A_1, A_2 = O(1)$ .

<sup>16</sup>A. Frishman and C. Herbert (submitted). *Phys. Rev. Lett.*

# Fluctuations: turbulent energy profile

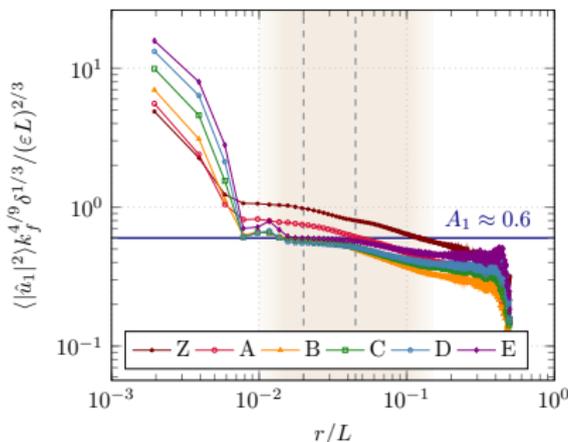
Behavior of harmonics explained by *zero modes of advection by the mean-flow*<sup>16</sup>

$$\langle |\hat{u}_1|^2 \rangle = A_1 k_f^{-4/9} \delta^{-1/3} (\varepsilon L)^{2/3} + \dots = A_1 (\varepsilon R_u)^{2/3} + \dots,$$

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# Fluctuations: turbulent energy profile

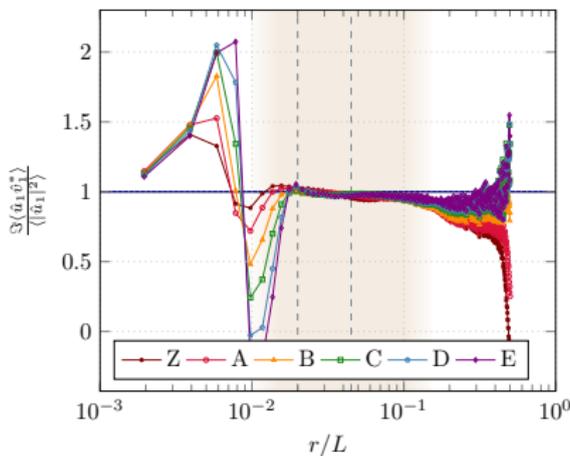
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# Fluctuations: turbulent energy profile

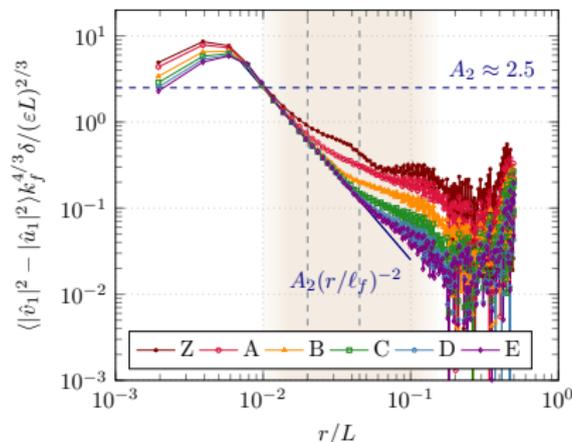
Behavior of harmonics explained by *zero modes of advection by the mean-flow*<sup>16</sup>

$$\langle |\hat{u}_1|^2 \rangle = A_1 k_f^{-4/9} \delta^{-1/3} (\varepsilon L)^{2/3} + \dots = A_1 (\varepsilon R_u)^{2/3} + \dots,$$

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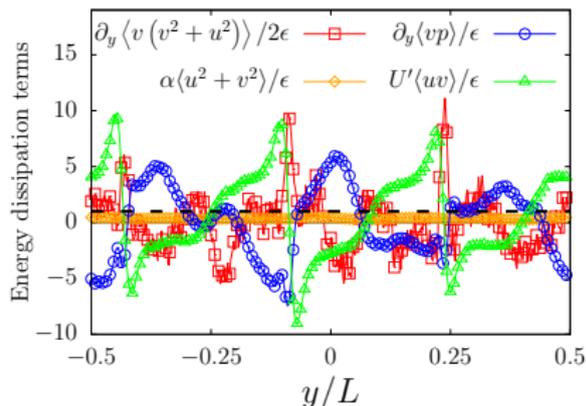
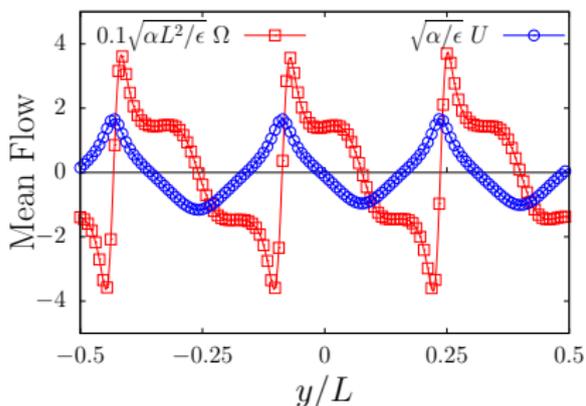
with  $A_1, A_2 = O(1)$ .



<sup>16</sup>A. Frishman and C. Herbert (submitted). *Phys. Rev. Lett.*

# Prospect: Zonal jet profile and fluctuations

Is the jet profile determined by the forcing-advection balance  $U'\langle uv \rangle = \epsilon$ ?<sup>17</sup>



Simulations by Jason Laurie (Aston University, UK)

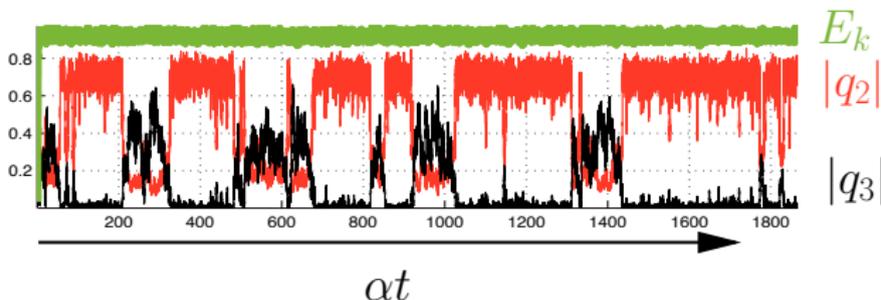
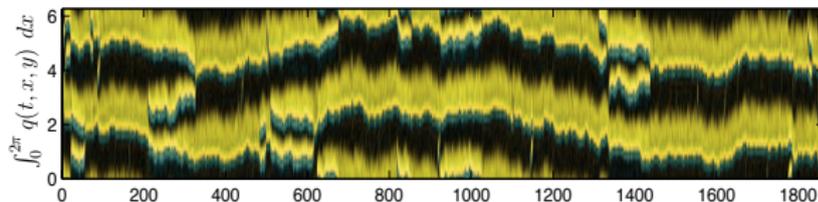
<sup>17</sup>E. Woillez and F. Bouchet (2017). *EPL*.

# Prospect: Bistability in zonal jets

Zonal jets in the stochastic barotropic vorticity equation:

$$\partial_t \omega + \mathbf{u} \cdot \nabla \omega + \beta v = -\alpha \omega + \nu \Delta \omega + f_\omega$$

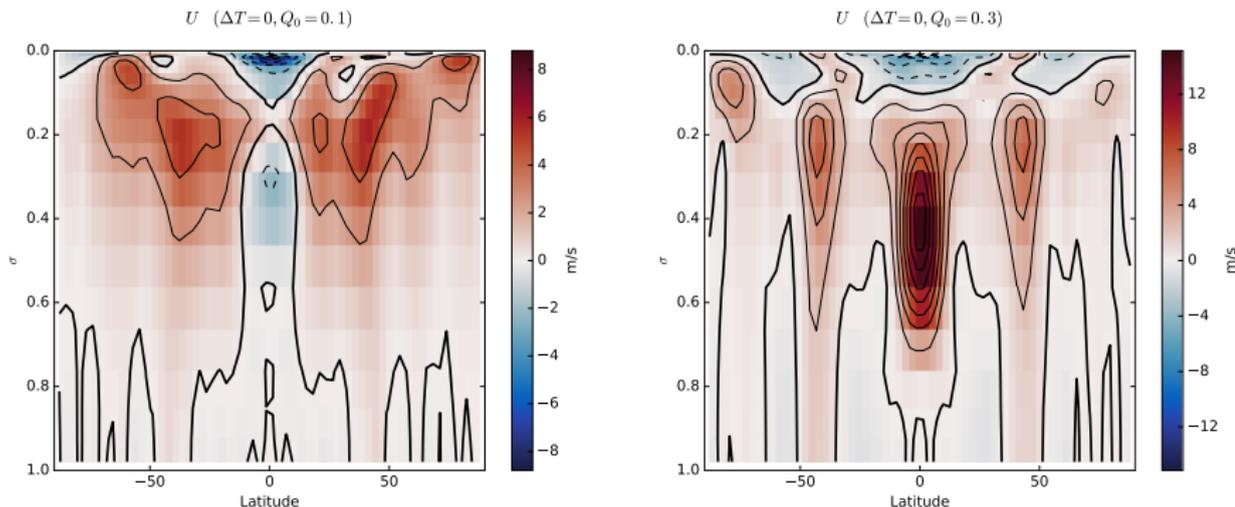
$$\beta = 0.555 \beta_{\text{mid}} \quad \alpha = 1.5 \cdot 10^{-3}$$



Simulations by Eric Simonnet (Inphyni, Nice).

# Prospect: Bistability in zonal jets

Zonal wind in idealized GCM (T. Schneider, Caltech):



- ▶ Held-Suarez forcing ( $T_{eq} = 315$  K):

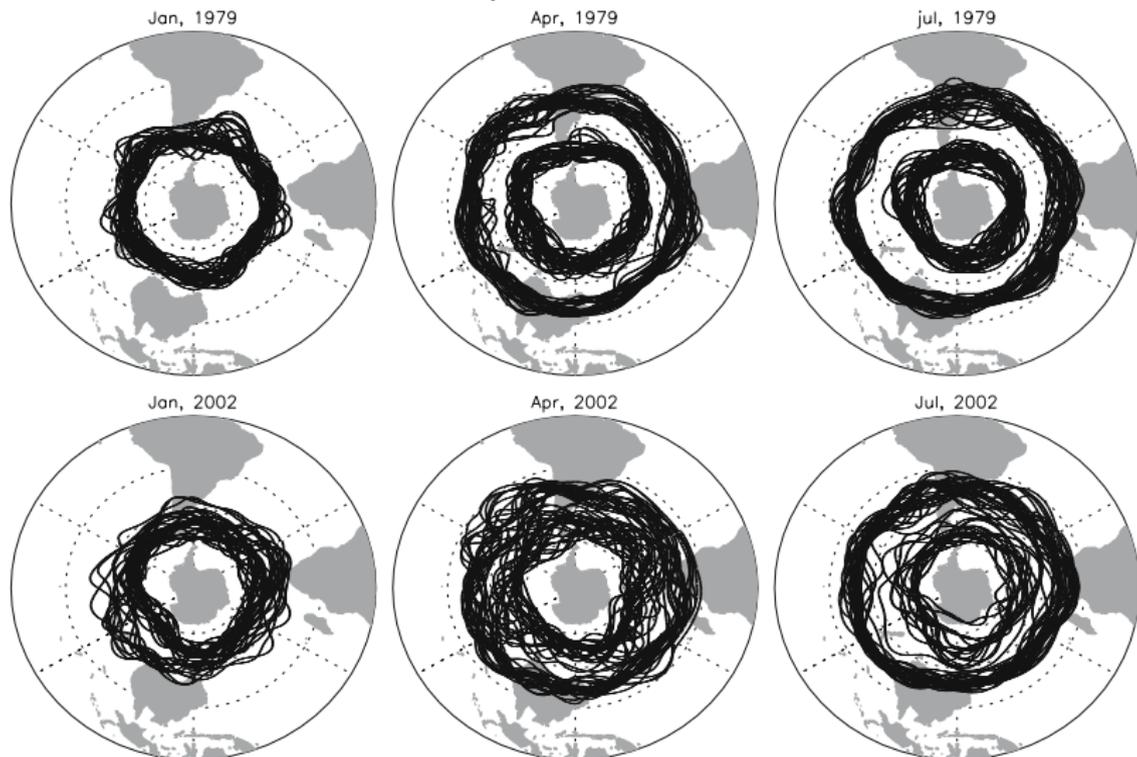
$$T_*(p, \phi) = [T_{eq} - \Delta T \sin^2 \phi - (\Delta\theta)_z \ln(p/p_0) \cos^2 \phi] (p/p_0)^{R/c_p}.$$

- ▶ Equatorial forcing:

$$Q(p, \phi, \lambda) = Q_0 \cos(k\lambda) \exp\left(-\frac{\phi^2}{\Delta\phi^2}\right) \sin\left(\pi \frac{p - p_t}{p_b - p_t}\right).$$

# Prospect: Bistability in zonal jets

## Jet Stream in the Southern Hemisphere<sup>18</sup>



<sup>18</sup>D. Gallego et al. (2005). *Clim. Dyn.*

# Summary

## Kinetic theory

- ▶ Asymptotic statistical closure of the Navier-Stokes equations based on timescale separation
- ▶ Reproduces emergence of large-scale coherent structures
- ▶ Allows explicit computations of mean-flow and fluctuations profiles in idealized context, in agreement with DNS

## Prospects

- ▶ Reduced model to study slow dynamics of eddy-driven jets: attractors, fluctuations,...
- ▶ Combined with large deviation theory, study abrupt transitions in jet dynamics



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# Can the mean-flow be described by an equilibrium state?

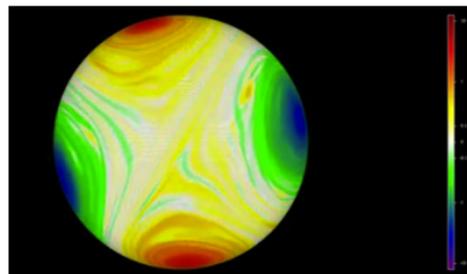
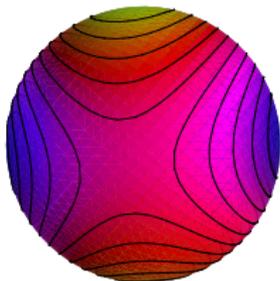
**2D Euler equations:**  $\partial_t \omega + \mathbf{u} \cdot \nabla \omega = 0$ .

Unlike 3D HIT, they have non-trivial equilibrium states<sup>19</sup>.

**Invariants:**  $\int s(\omega(\mathbf{r})) d\mathbf{r}$ .

At equilibrium, all the energy is in the mean-flow; no fluctuations.

Ex: flow on a sphere



Theoretical Equilibrium: Quadrupole<sup>20</sup>

DNS Final State<sup>21</sup>

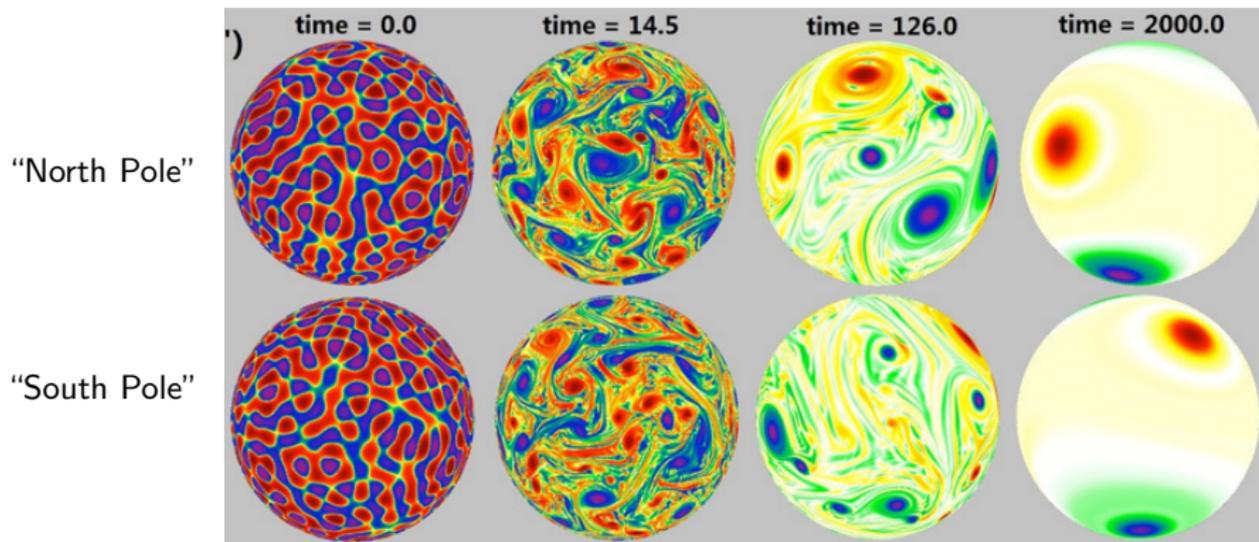
<sup>19</sup>L. Onsager (1949). *Il Nuovo Cimento*; R. H. Kraichnan (1967). *Phys. Fluids*; R. Robert and J. Sommeria (1991). *J. Fluid Mech.*

<sup>20</sup>C. Herbert (2013). *J. Stat. Phys.*

<sup>21</sup>W. Qi and J. B. Marston (2014). *J. Stat. Mech.*

# The effect of rotation (DNS results<sup>22</sup>)

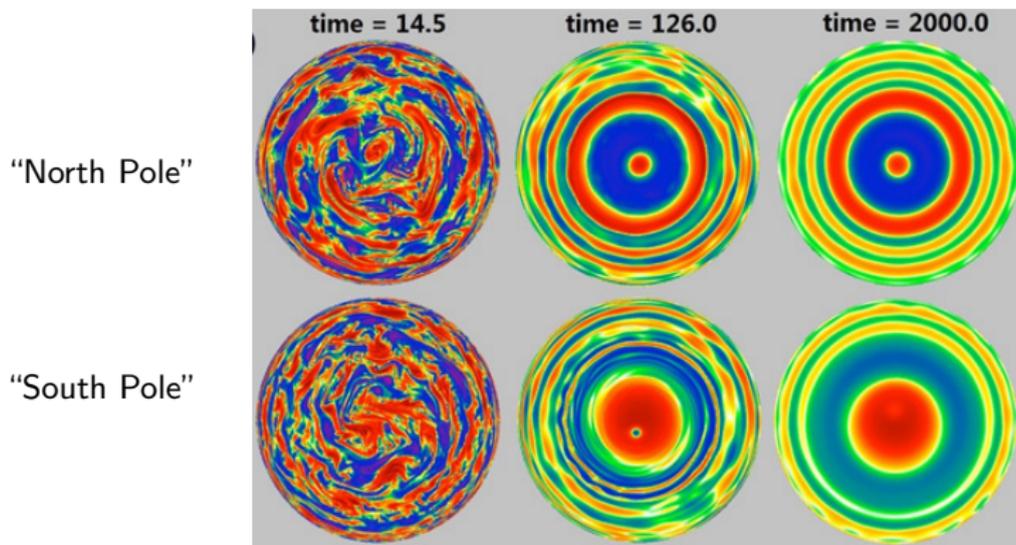
When the Rossby waves are sufficiently slow, the system relaxes towards its equilibrium state.



<sup>22</sup>W. Qi and J. B. Marston (2014). *J. Stat. Mech.*

## The effect of rotation (DNS results<sup>22</sup>)

For faster rotation rates, Rossby waves arrest the cascade at the Rhines scale and lead to the emergence of zonal flows.



<sup>22</sup>W. Qi and J. B. Marston (2014). *J. Stat. Mech.*

# The closure problem

## Reynolds decomposition:

$$u_i = \bar{u}_i + u'_i,$$

where  $\bar{\cdot}$  is a projection operator.

The Navier-Stokes equations become:

$$\begin{aligned}\partial_t \bar{u}_i + \bar{u}_j \partial^j \bar{u}_i &= -\partial_i \bar{P} + \nu \partial_j \partial^j \bar{u}_i - \partial^j \overline{u'_i u'_j}, \\ \partial_t u'_i + \bar{u}_j \partial^j u'_i + u'_j \partial^j \bar{u}_i &= -\partial_i P' + \nu \partial_j \partial^j u'_i - \partial^j u'_i u'_j + \partial^j \overline{u'_i u'_j}.\end{aligned}$$

The major difficulty is to compute the Reynolds stress tensor  $-\partial^j \overline{u'_i u'_j}$ .

## Modeling approaches:

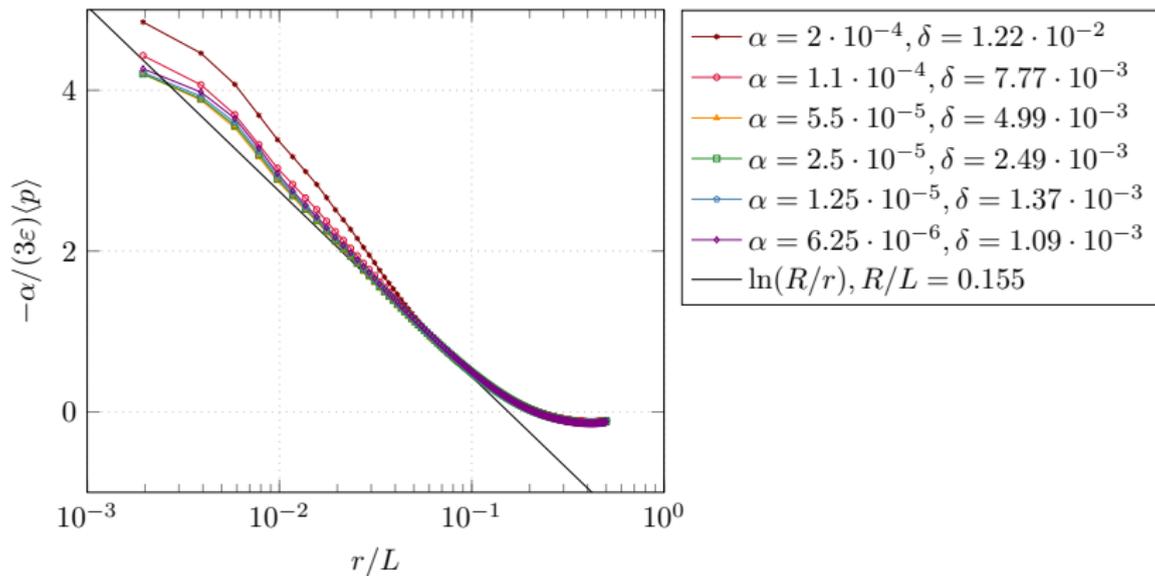
- ▶ *Large Eddy Simulations*: spatial filtering

$$\bar{u}_i(\mathbf{x}, t) = \int G(\mathbf{x} - \mathbf{y}) u_i(\mathbf{y}, t) d\mathbf{y}$$

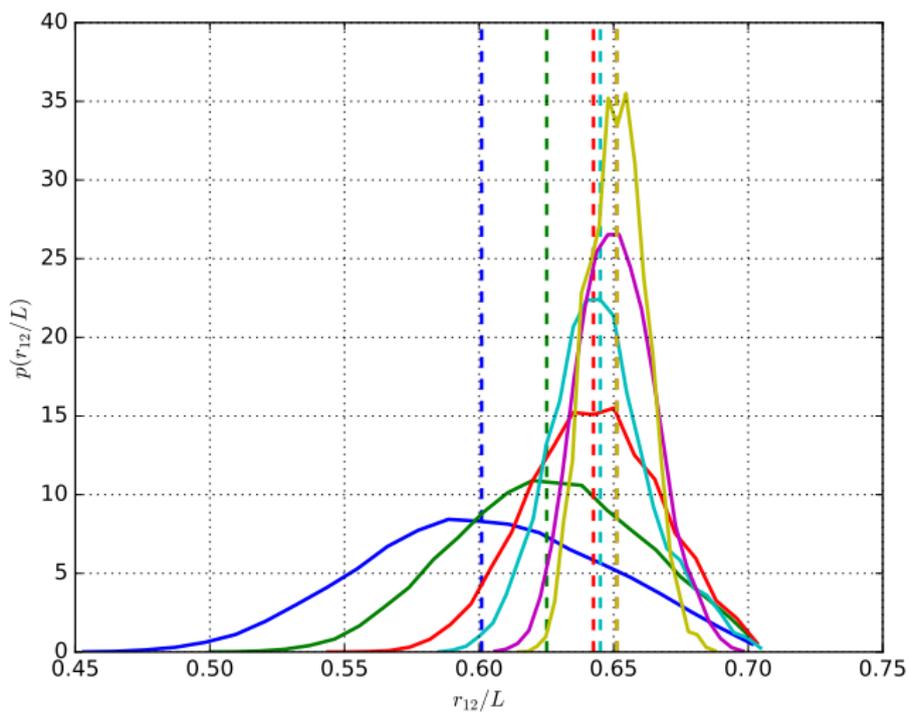
- ▶ *Reynolds Average Navier-Stokes*: time filtering

# Mean pressure profile (DNS)

DNS:  $512^2$ ,  $k_F = 100$ , hyperviscosity,  $\sim 300000$  turnover times.



## Intervortex distance

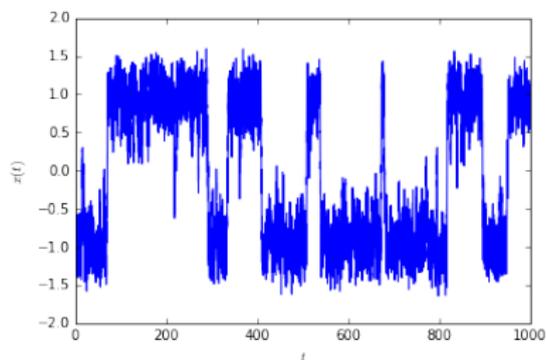
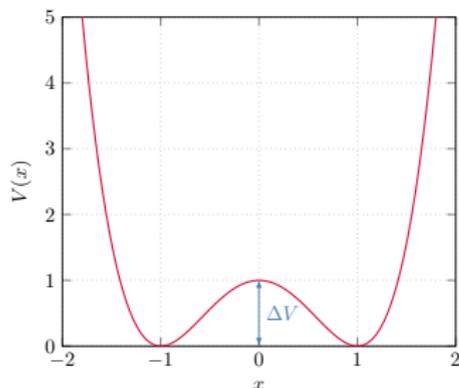


*Simple stochastic process reproducing the dynamics of the inter-vortex distance?*

Theoretical framework for noise induced transitions: the Kramers problem<sup>23</sup>

Overdamped Langevin dynamics:

$$\dot{x} = -V'(x) + \sqrt{2\epsilon}\eta, \quad V(x) = (x^2 - 1)^2, \quad \mathbb{E}[\eta(t)\eta(t')] = \delta(t - t').$$

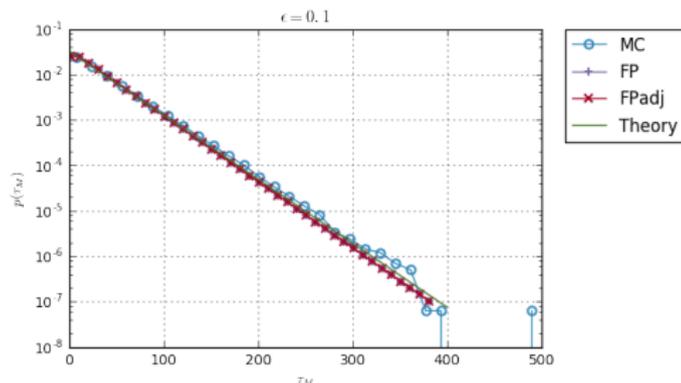
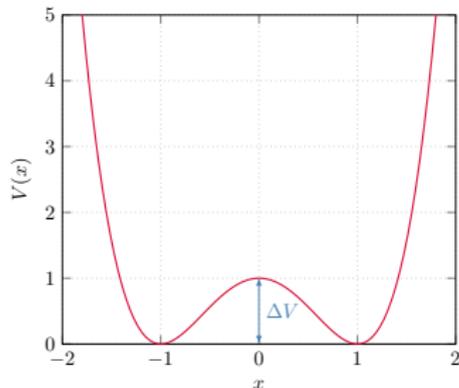


<sup>23</sup>H. A. Kramers (1940). *Physica*.

Theoretical framework for noise induced transitions: the Kramers problem<sup>23</sup>

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### Transition probability

In the weak noise limit ( $\epsilon \rightarrow 0$ ), transition times form a Poisson point process with transition rate  $\lambda = \tau^{-1} e^{-\Delta V/\epsilon}$ .

This is a large deviation result.

<sup>23</sup>H. A. Kramers (1940). *Physica*.

Theoretical framework for noise induced transitions: the Kramers problem<sup>23</sup>

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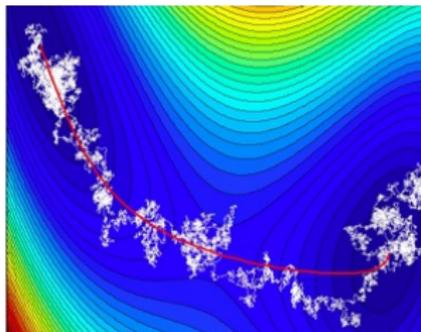
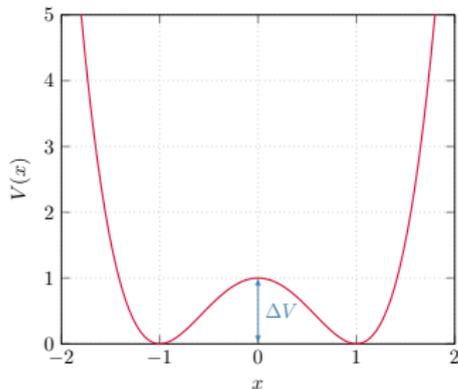


Fig. E. Vanden-Eijnden (Courant)

## Instantons

### Path integral formalism

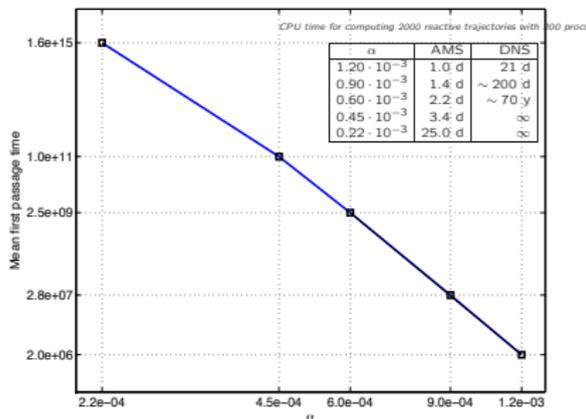
$$\mathbb{E}[\mathcal{O}] = \int \mathcal{D}[x] \mathcal{O}[x] \exp(-\mathcal{A}[x]/\epsilon), \quad \text{Action: } \mathcal{A}[x] = \frac{1}{4} \int dt (\dot{x} + V'(x))^2.$$

**Instanton:** most probable path:  $\min_x \{\mathcal{A}[x] | x(-T) = -1, x(T) = 1\}$ .

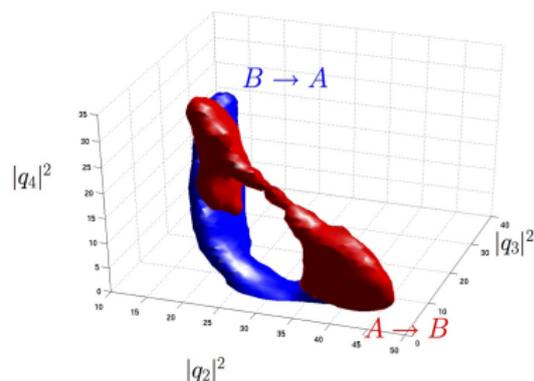
<sup>23</sup>H. A. Kramers (1940). *Physica*.

# Arrhenius law and Instantons in jet transitions

Numerical algorithms to compute large deviations: dynamics biased in a controlled way<sup>24</sup>.



Arrhenius law



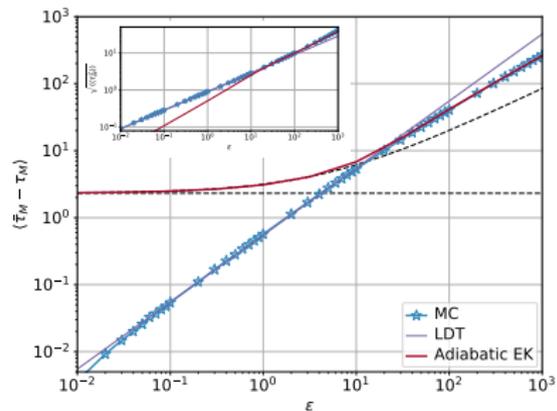
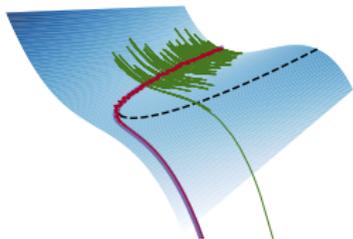
“Instantons”

Jet transition simulations with rare event algorithm (AMS) by Eric Simonnet (Inphyni).

<sup>24</sup> C Giardinà et al. (2011). *J. Stat. Phys.* F. Cérou and A. Guyader (2007). *Stoch. Anal. Appl.*

Escape in stochastic saddle-node bifurcation<sup>25</sup>

$$dx_t = (x_t^2 + t)dt + \sqrt{2\epsilon}dW_t, \quad \tau_M = \inf\{t \geq t_0, x_t \geq M\}$$



Competition between deterministic and stochastic effects.

<sup>25</sup>C. Herbert and F. Bouchet (2017). *Phys. Rev. E*.